# Mantle Convection: Current Status and Future Challenges

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## Outline

### 1) Introduction.

What is mantle convection and why it is interesting/important.

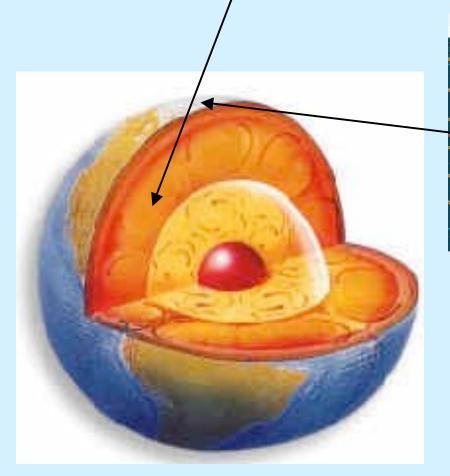
#### 2) Current Status.

- Numerical methods & codes.
- Example problems.
- 3) Unique challenges.
- Multi-scale physics and non-linear and highly variable rheology.
- Ongoing and future developments

# What is mantle convection?

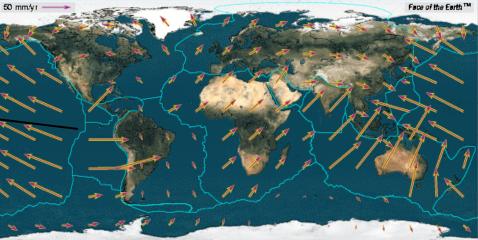
- <u>Objectives:</u> Understand the long-term heat and mass *transfer* and structure *formation* in the mantle and their *implications* for surface geological, geochemical and geophysical observables and for evolution of the Earth and other planets (terrestrial planets and icy satellites.
- <u>Physical basis:</u> Treat the mantle as viscous flow. Conservation of *MASS*, *ENERGY*, and *MOMENTUM* + *VISCOUS* rheological equation.

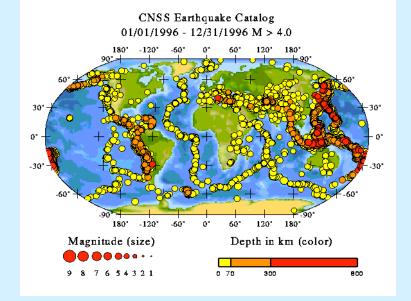
## The Mantle – The Earth's "Engine"



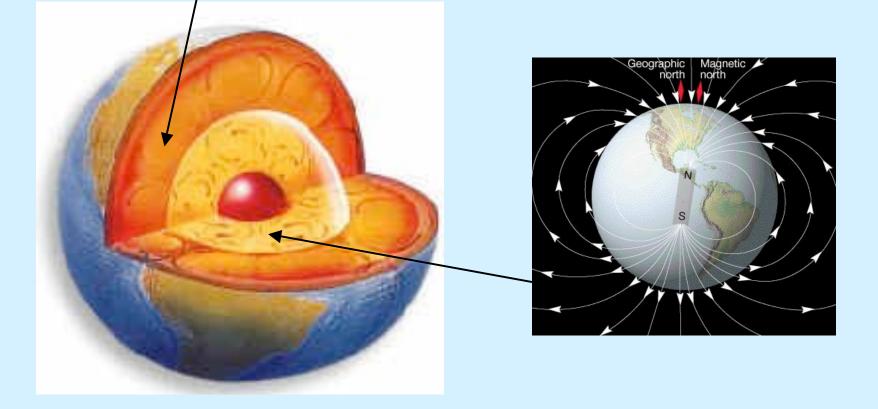
To remove radiogenic heating & primordial heating

Present-day surface plate motion

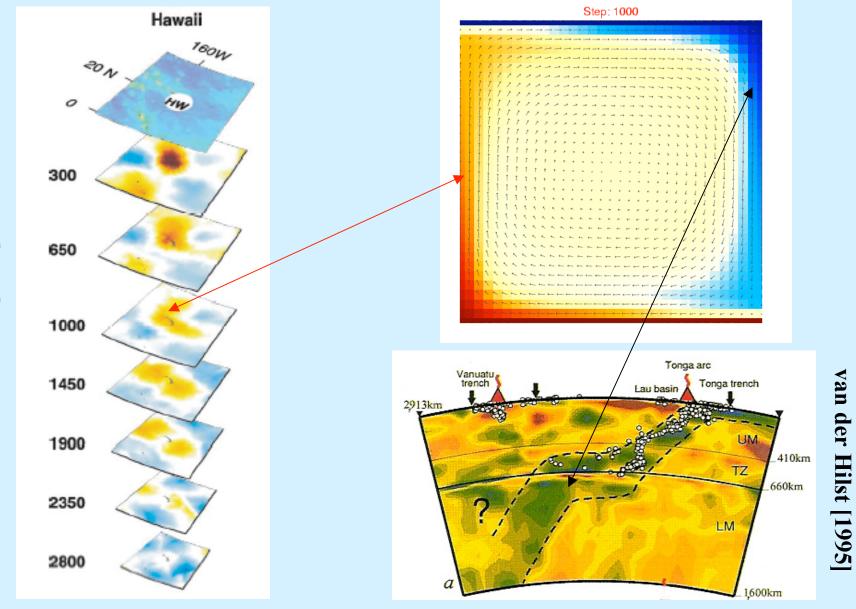






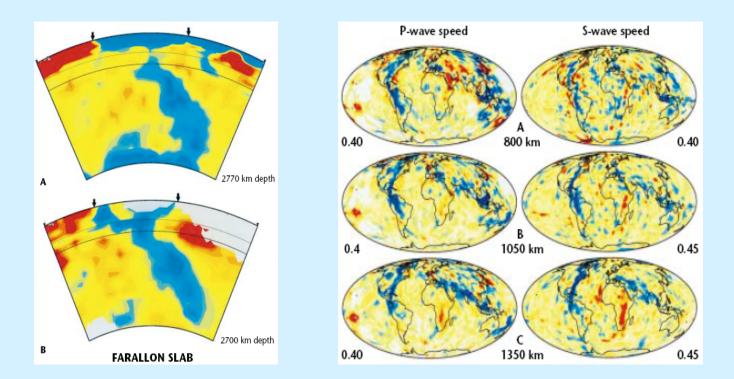


### A Simple Picture of Mantle Convection: Boundary Layers



Montelli et al. [2004]

## More seismic images of mantle convection

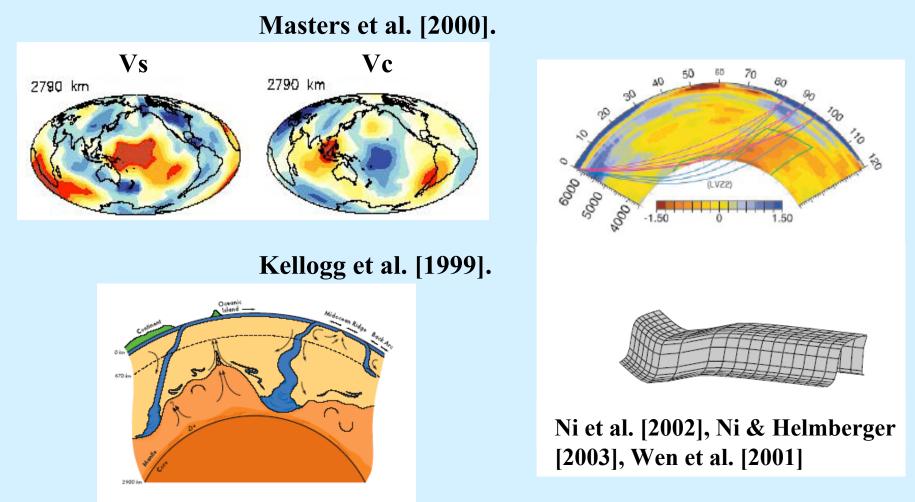


Grand, van der Hilst, & Widiyantoro [1997]

van der Hilst et al. [1992; 1997], Grand [1994], and early models by Dziewonski's group, Masters' group, and Romanowicz's group.

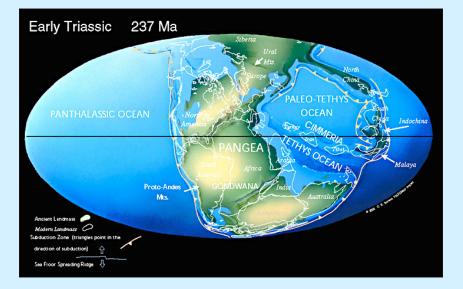
#### **Compositionally Distinct Mantle Reservoirs**

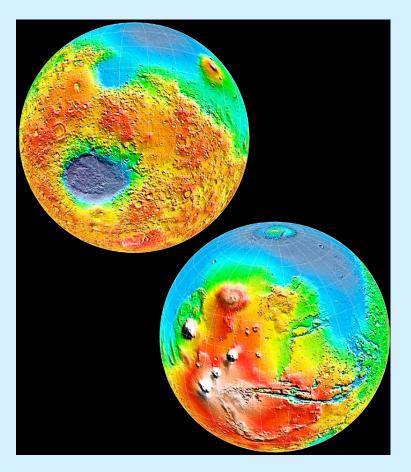
A lot of geochemical evidences.



Important implications for the cooling of the core and geodynamo!

### Back in time or out into the space ...





## Governing Equations (a simplified form)

Mass conservation:

$$\nabla \cdot \boldsymbol{u} = 0,$$

Momentum conservation:

$$-\nabla p + \nabla \cdot [\eta (\nabla \boldsymbol{u} + \nabla^T \boldsymbol{u})] - \delta \rho g_0 \boldsymbol{e}_r = 0,$$

**Energy conservation:** 

$$\rho_0 C_p \left( \frac{\partial T}{\partial t} + \boldsymbol{u} \cdot \nabla T \right) = k \nabla^2 T + \rho_0 H,$$

Stokes' flow

**Equation of state:** 

$$\delta \rho = -\rho_0 \alpha (T - T_0) + \Delta \rho C,$$

**Composition conservation:** 

$$\frac{\partial C}{\partial t} + \boldsymbol{u} \cdot \nabla C = \boldsymbol{0},$$

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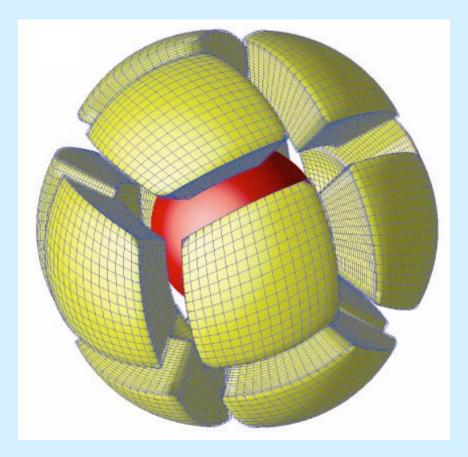
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## **Various Methods**

- 1) <u>Spectral methods</u> [Glatzmaier et al., 1990; Zhang & Christensen, 1993; Zhang & Yuen, 1995].
- 2) *Finite volume methods* [Ratcliff et al., 1997; Harder & Hansen, 2005; Tackley, 2008].
- 3) *Finite elements* [Baumgardner, 1985; King et al., 1990; Moresi & Gurnis, 1996; Zhong et al., 2000].

#### Some General Features in CitcomS



Zhong, Zuber, Moresi & Gurnis [2000]

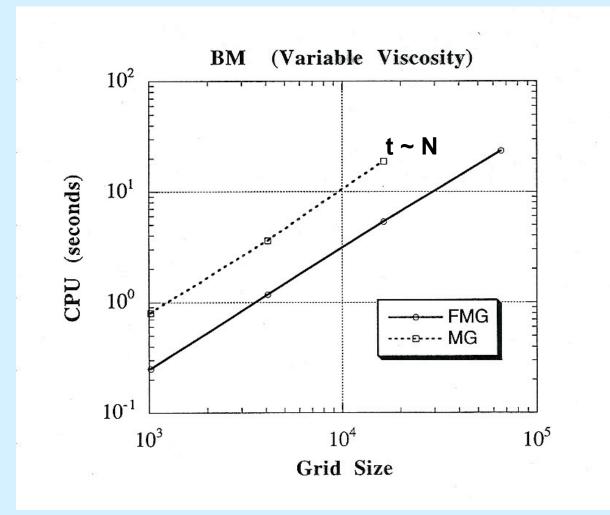
Stokes flow solver: Uzawa Derived from the 3D algorithm with two level Cartesian code Citcom iterations [Ramage and Moresi & Gurnis, 1996]. Wathen, 1994].

Grid: First divide a spherical Energy equation solver: shell into 12 caps, and then Streamline Upwind Petroy further divide each cap into Galerkin formulation elements with roughly [Brooks, 1981] uniform size. Each cap has

the same number of elements. All the equations are

explicitly written in spherical Elements: Brick element (8 geometry and then coded up yelocity nodes and constant in stiffness matrix, force,... pressure per element).

## **Geometric Multigrid for the Inner Loop Solver in the Uzawa**



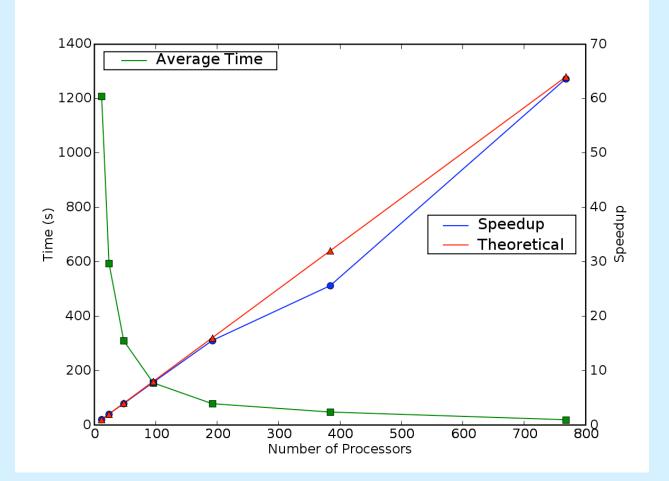
## **CitcomS' Parallel Efficiency**

#### **CPU Time vs Number of Cores on TACC's Ranger**

Number of cores	total time (sec)	iterations	time per v- iteration (sec)	Efficiency (%)
12(1x1x1)	69.8	112(118)	0.59	100
24(1x1x2)	64.1	95(103)	0.62	95
48(2x2x1)	53.7	73(78)	0.69	86
96(2x2x2)	53.9	74(79)	0.68	87
192(4x4x1)	47.2	55(63)	0.75	79
384(4x4x2)	46.8	55(58)	0.81	73
768(4x4x4)	52.4	58(61)	0.86	69
1536(8x8x2)	58.2	60(70)	0.83	71
3072(8x8x8)	59.1	54(57)	1.04	57

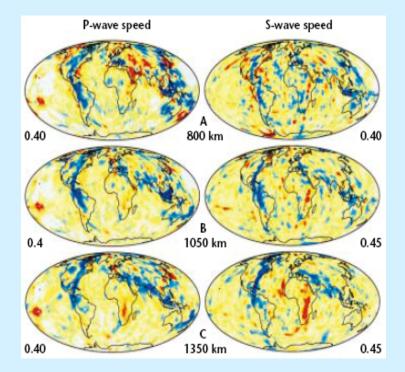
Zhong et al. [2008, G<sup>3</sup>, in press]

#### CitcomS' Parallel Efficiency on BlueGen (only conj-grad)

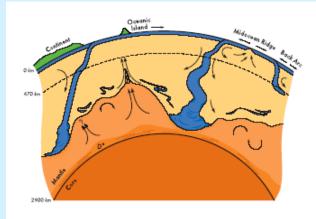


**Provided by Scott King** 

## More seismic images of mantle convection



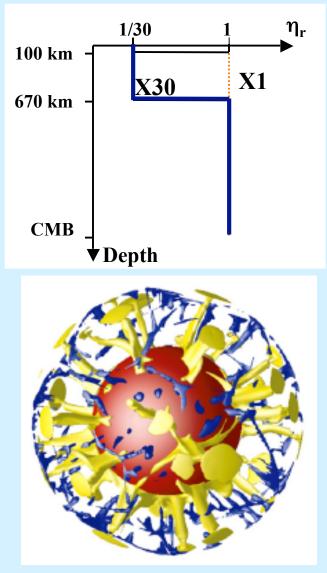
Kellogg et al. [1999].

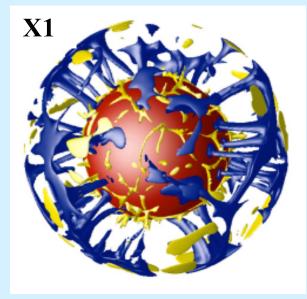


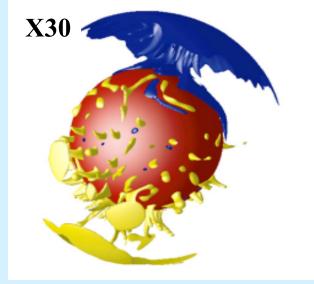
Grand, van der Hilst, & Widiyantoro [1997]

#### **Example 1: Structure Formation [Zhong et al., 2007]**

 $\eta = \eta_r \exp[E(0.5-T)]$ 

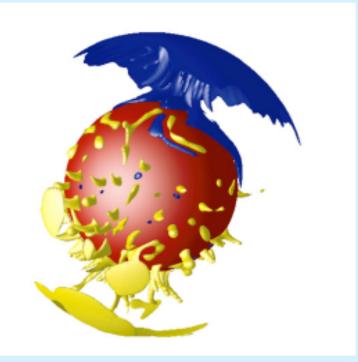


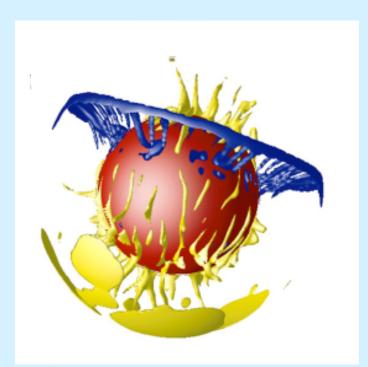




### Effects of a supercontinent on mantle structure

Add a supercontinent

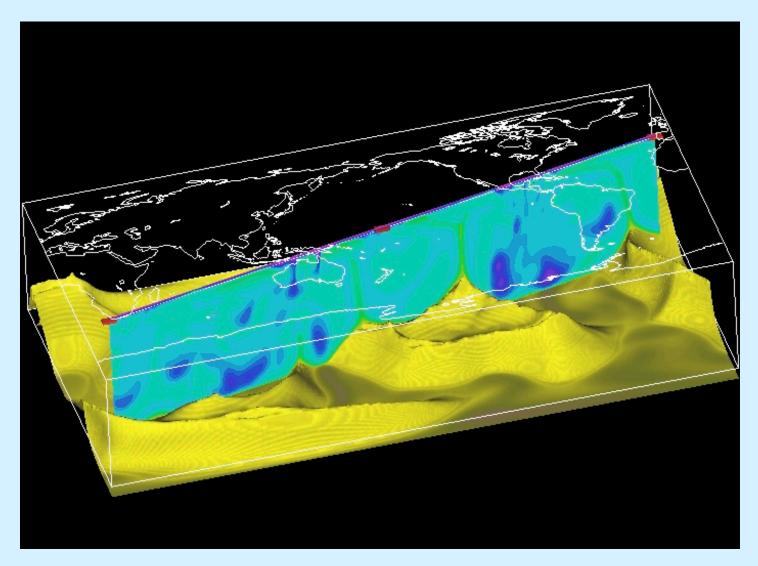




## Time evolution of mantle structure

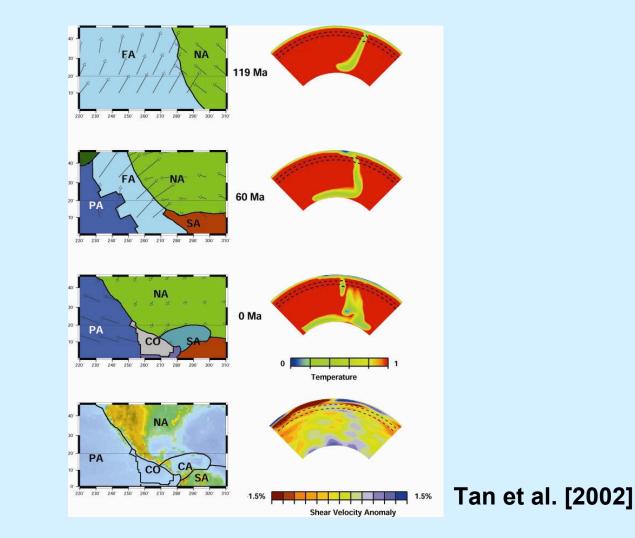


#### Mantle Convection with Distinct Mantle Reservoirs

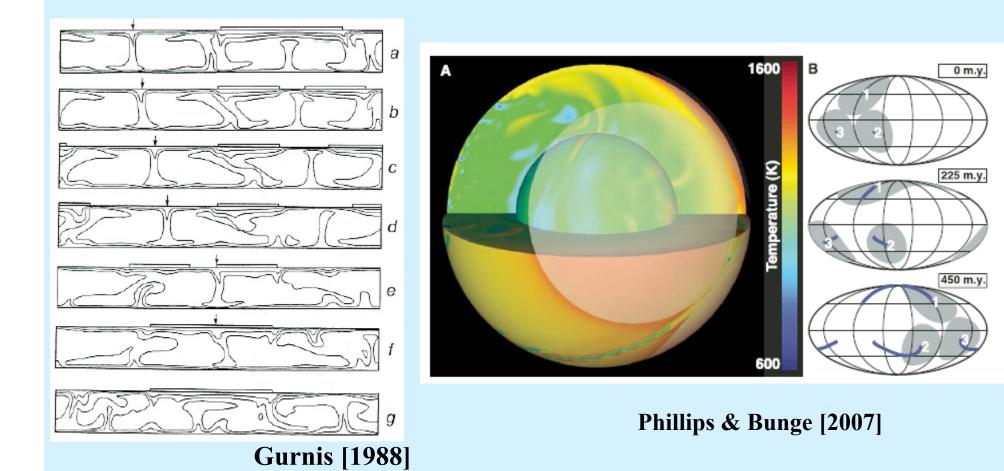


McNamara and Zhong [2005]

## **Example: Modeling the evolution of Farallon Subduction**



#### **Example: Time evolution of mantle structure and** continents

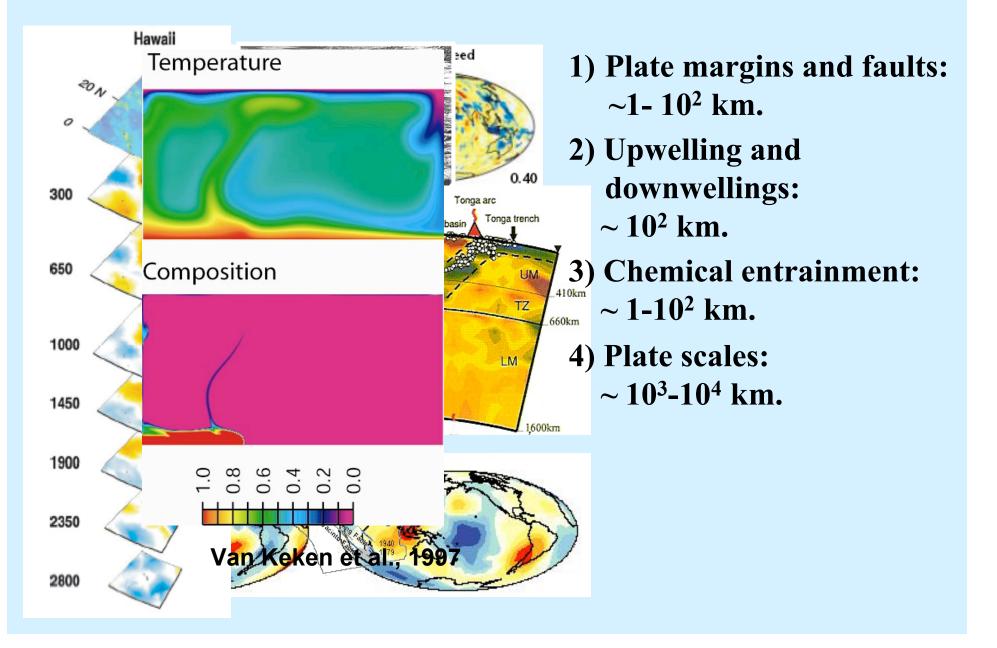


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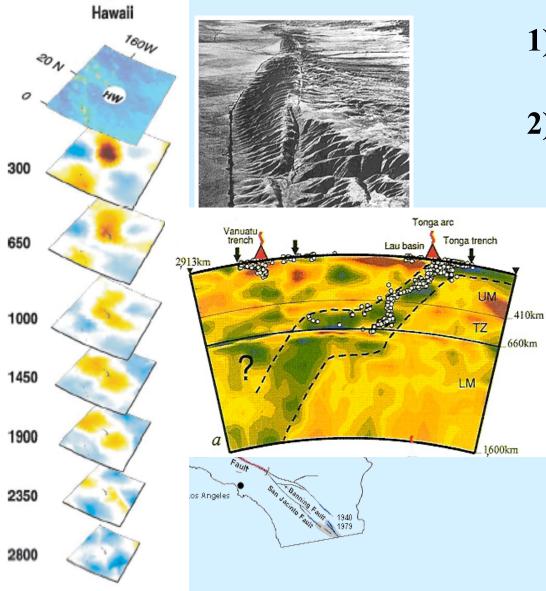
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## Multi-scale (from $10^{\circ}$ km to $10^{4}$ km) Physics



## Nonlinear and Highly Variable Rheology



- 1) Faulting and Plastic deformation: nonlinear
- 2) Highly temperaturedependent viscosity for silicate mantle (up to 3 orders of magnitude variations in plumes and
  https://www.ellings.

## **Ongoing and Future Developments**

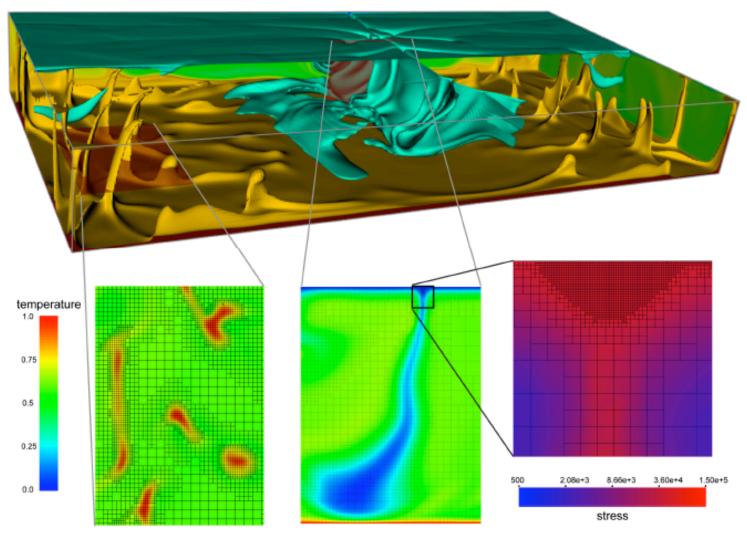
• Better treatment of multi-scale processes (both physics and numerics).

*i) Lithospheric plates and plate boundaries. ii) Melting.* 

iii) Thermo-chemical convection and entrainment.

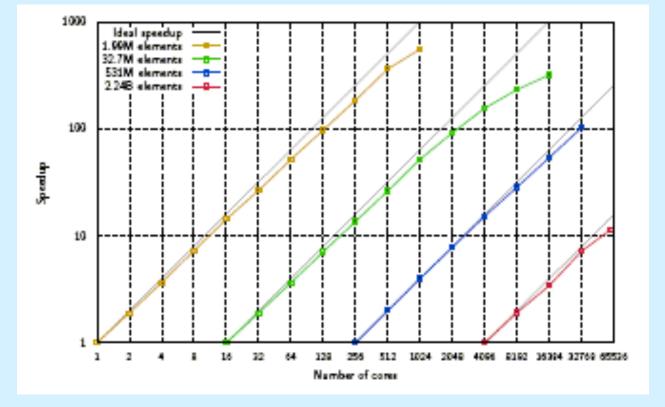
- Adaptive mesh refinement (AMR) [Davies & Davies, 2008 a 2D study; Bangerth's deal-II ].
- Efficient AMR on ~60K cores [Burstedde et al., 2008 – a collaborative effort by UT-Austin, Caltech, & UC-Boulder]

#### A new code: Rhea (octree-based AMR & massively parallel; local resolution to 1.5 km!)

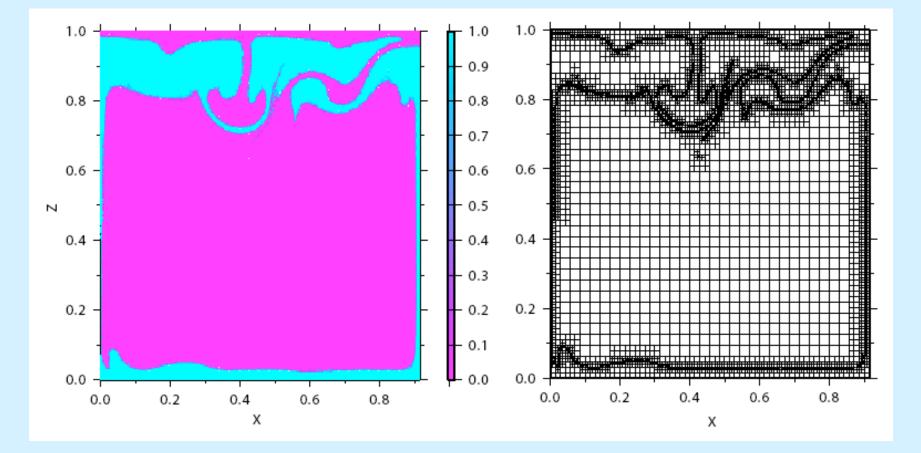


Burstedde et al. [2008]

#### Rhea's Parallel Efficiency (fixed problem sizes)



## Applying octree-based AMR to 2-D thermochemical convection



Leng & Zhong [2008 Fall AGU]