

Slow slip events simulated using a friction law with a velocity-weakening to -strengthening transition

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Allan Rubin

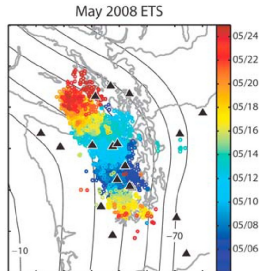
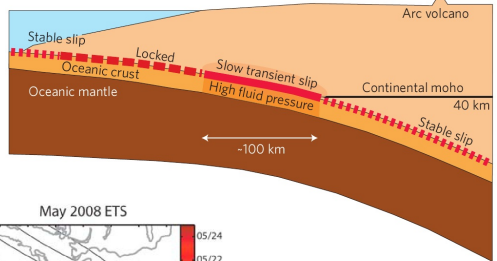
Princeton University

June 22, 2012

Slow slip events (in Cascadia)

- Every 12-16 months
- 2-3 centimeters of slip
- Last 3-5 weeks
- Slip rates 10-100 times plate rate
- Accompanied by tremor: lots of very small earthquakes on the plate interface
- Propagate several hundred kilometers along strike

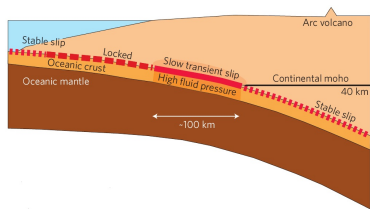
Peng and Gomberg, 2011



3,677 epicenters, 227 hours

Wech et. al., 2009

Outline



Peng and Gomberg, 2011

Which friction law is appropriate?

- Standard velocity-weakening friction with tuned size
- Dilatancy
- Velocity-weakening to -strengthening friction
- All can produce episodic slow slip
- All require tuning
- But can they reproduce all of the observed features of slow slip?

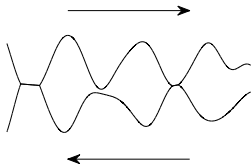
Here: Try to understand slow slip events simulated with a friction law with a velocity-weakening to -strengthening transition.

- Model motivation and setup
- Model predictions
 - Steady propagation
 - Stress drops
 - Gradual modulation due to tidal forcing
 - Back-propagating fronts

Rate and state friction

frictional strength = $f(\text{slip rate } V, \text{ state } \theta)$

$$\mu(V, \theta) = \text{constant} + a \ln\left(\frac{V}{V^*}\right) + b \ln\left(\frac{V_c \theta}{D_c}\right)$$



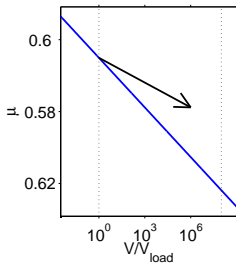
$a \ln\left(\frac{V}{V^*}\right)$: direct effect

- accounts for adhesion at individual asperities
- **larger for steady slip at higher slip rates**

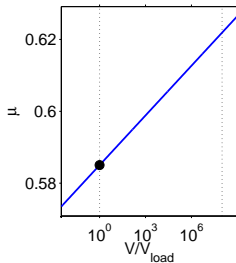
$b \ln\left(\frac{V_c \theta}{D_c}\right)$: evolution effect

- accounts for number and size of asperities
- at larger slip rates, longer asperity lifetime, larger contact area
- **smaller for steady slip at higher slip rates**

$a < b$, evolution effect dominates, earthquakes



$a > b$, direct effect dominates, steady sliding

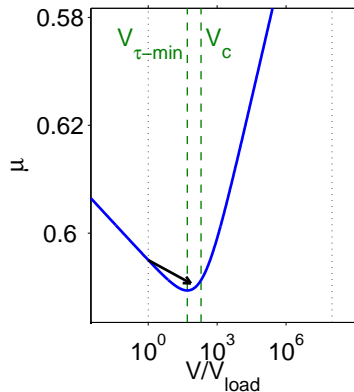
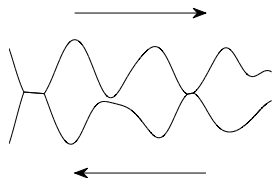


One constitutive law for episodic slow slip

frictional strength = $f(\text{slip rate } V, \text{ state } \theta)$

$$\tau(V, \theta) = \text{constant} + a\sigma \ln\left(\frac{V}{V^*}\right) + b\sigma \ln\left(\frac{V_c\theta}{D_c} + 1\right)$$

In theory and experiments, this law results from a lower limit on the size of asperities



Need equations for state evolution

aging law

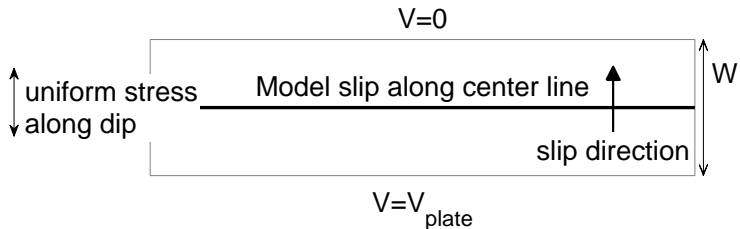
$$\frac{d\theta}{dt} = 1 - \frac{V\theta}{D_c}$$

slip law

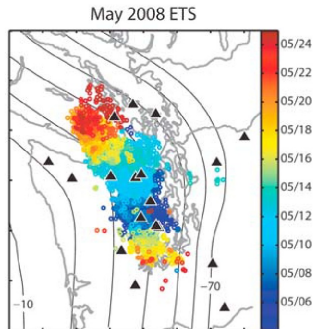
$$\frac{d\theta}{dt} = -\frac{V\theta}{D_c} \ln\left(\frac{V\theta}{D_c}\right)$$

velocity-weakening at low slip rates, velocity-strengthening at high slip rates

Model geometry



- Slow slip events often extend farther along strike than along dip
- Use one point per along-strike distance
- Assume stress is uniform along dip

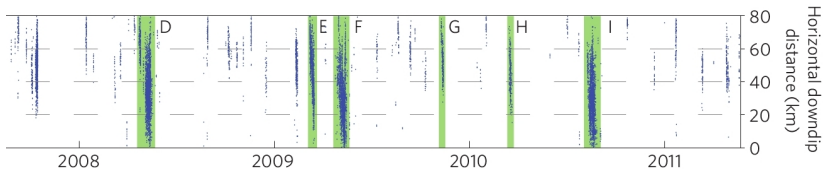
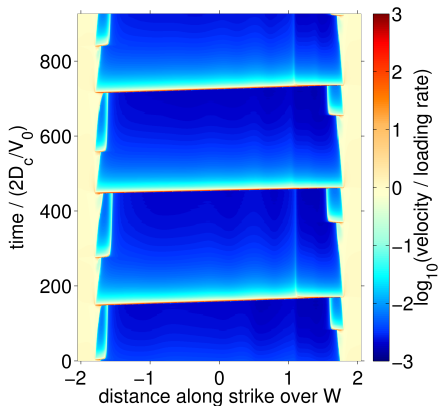


3,677 epicenters, 227 hours

Simulated slow slip events

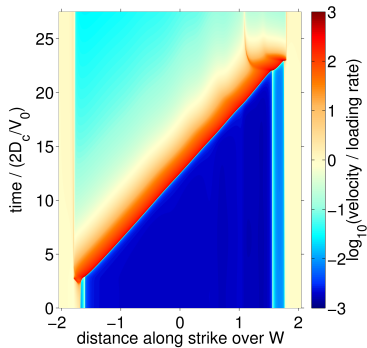
Model does produce episodic slow slip events

- Slip rates around V_c , 100 times plate rate
- Episodic large events with a number of small events between



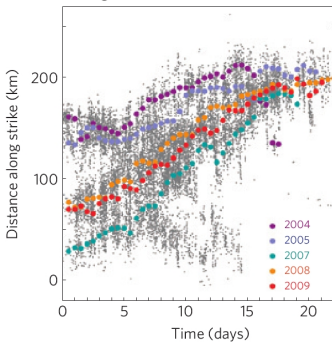
Steady propagation

Events propagate steadily
along strike

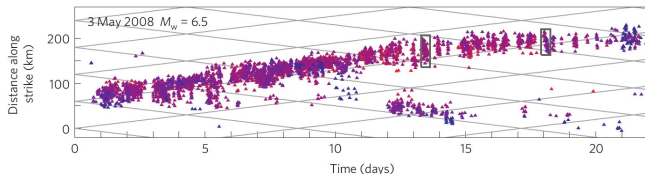


Tremor observations

Along-strike evolution of five ETS events

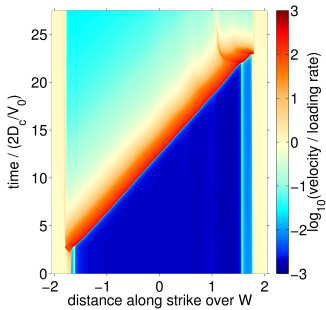


Houston et. al., 2011

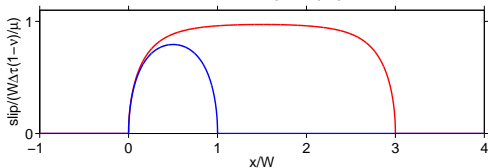


Steady propagation

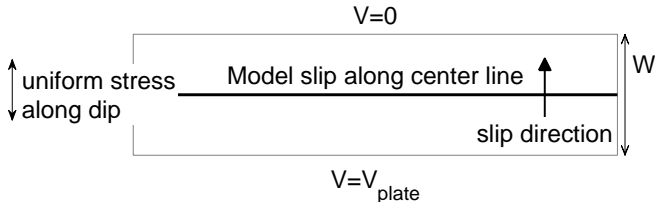
Events propagate steadily along strike



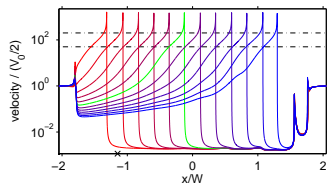
Uniform stress drop slip profiles



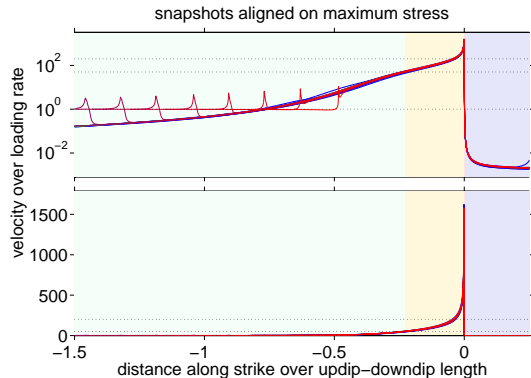
With the strip model geometry, points don't feel slip at locations much more than W away. This makes a steady solution possible.



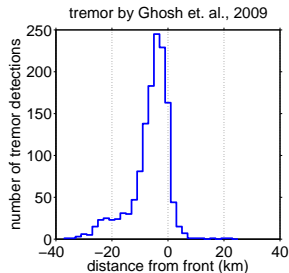
Steady velocity and stress profiles



- Most of the slip accumulates in a region behind the front somewhat smaller than the updip-downdip length W
- Velocity profiles resemble asymmetric tremor density

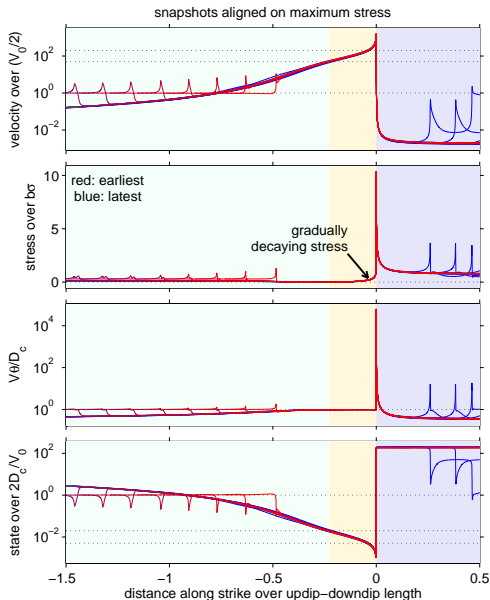
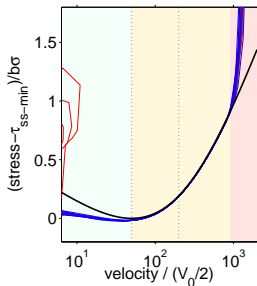


2 days of tremor stacked behind a tremor-defined front



Steady velocity and stress profiles

- purple: unbroken, well-healed
- pink: strongly slip-weakening region, above steady state
- yellow: gradually decaying stress, near steady state
 - Specific to this friction law
 - Know stress-velocity relation in this region
- green: low-stress, below steady state

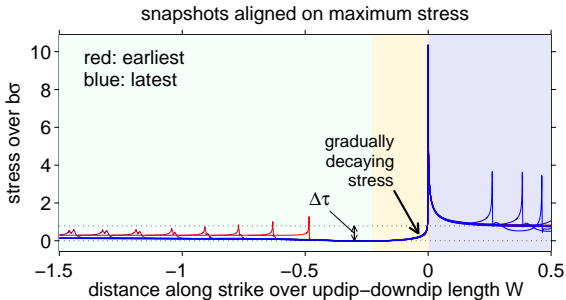
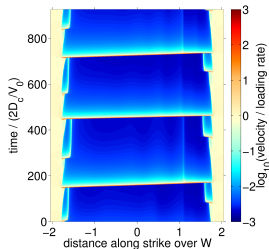


Conclusions to this point

- Simulations exhibit many small events as well as periodic large events
- Large events propagate steadily “along strike” with the strip model

Recurrence intervals: energy balance

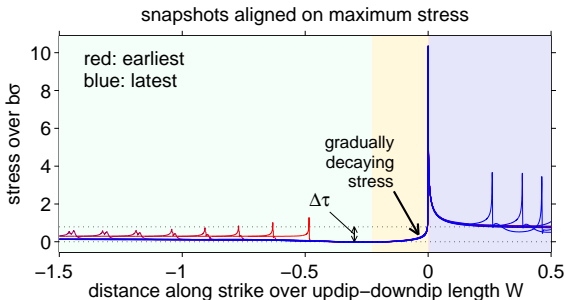
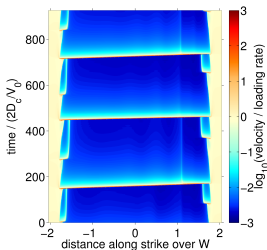
- controlled by the ability of events to propagate, not by nucleation.



strain energy release rate $G =$ energy dissipated by friction (fracture energy G_c)

Recurrence intervals: energy balance

- controlled by the ability of events to propagate, not by nucleation.



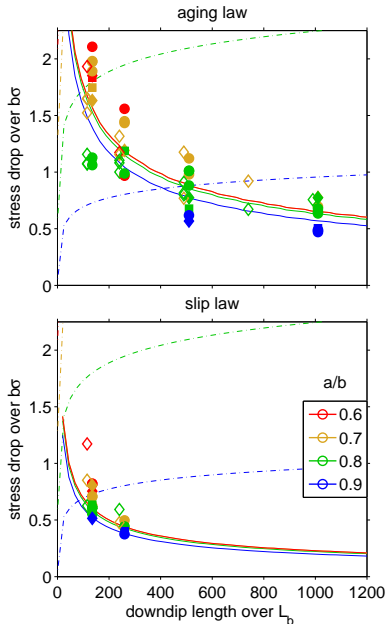
strain energy release rate $G =$ energy dissipated by friction (fracture energy G_c)
 constant \times stress drop \times slip \approx near-tip contribution + gradual decay contribution

$$\frac{\psi}{2} \Delta\tau \left(\Delta\tau \frac{W}{\mu} \right) \approx b\sigma D_c \times \text{evolution-law-dependent constant}$$

$$\Delta\tau \approx \psi' b\sigma \sqrt{\frac{L_b}{W}} \quad \text{where } L_b = D_c \frac{\mu}{b\sigma}$$

$b =$ evolution effect coefficient, $\psi =$ geometric factor,
 $D_c =$ slip distance for state evolution, $\mu =$ shear modulus

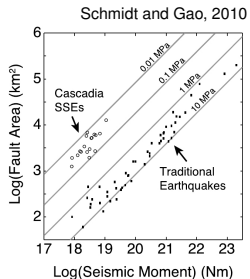
Recurrence intervals: simulation results



Observed along-dip length W : 50-100 km

Observed stress drops: 10-100 kPa

$$\text{slip} \times \mu / (1 - \nu) / W = 2 \text{ cm} \times 40 \text{ GPa} / 50 \text{ km} = 15 \text{ kPa}$$



$$\Delta\tau \approx \psi' b\sigma \sqrt{\frac{L_b}{W}} \quad \text{where } L_b = D_c \frac{\mu}{b\sigma}$$

$$b\sigma D_c \approx \frac{\Delta\tau^2 W}{\mu\psi'^2}$$

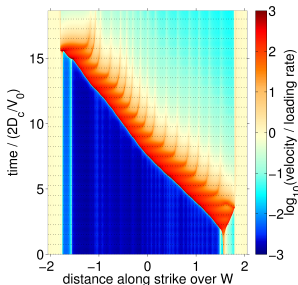
Required model parameters:

- Slip law: $D_c b\sigma \approx 3 - 200 \text{ Pa}\cdot\text{m}$
- Aging law: $D_c b\sigma \approx 0.3 - 20 \text{ Pa}\cdot\text{m}$
- For the slip law, $b = 0.01$
- For $D_c = 10 \text{ } \mu\text{m}$: $\sigma \approx 30 \text{ MPa} - 1 \text{ GPa}$
- For $D_c = 100 \text{ } \mu\text{m}$: $\sigma \approx 3 \text{ MPa} - 200 \text{ MPa}$
- For $D_c = 1 \text{ mm}$: $\sigma \approx 0.3 \text{ MPa} - 20 \text{ MPa}$

Conclusions to this point

- Simulations exhibit many small events as well as periodic large events
- Large events propagate steadily “along strike” with the strip model
- Stress drops controlled by ability to propagate long distances along strike
 - Understood with an energy balance approach

Effect of a tidal forcing

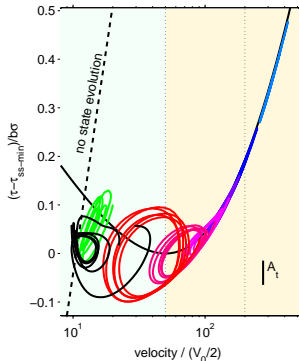
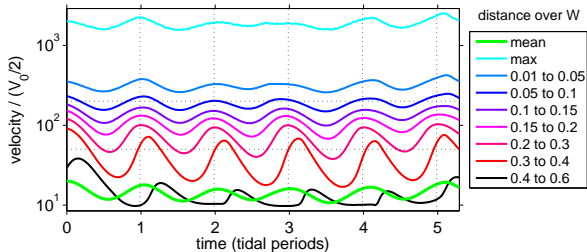


Tides introduce a gradual modulation of the slip rate.

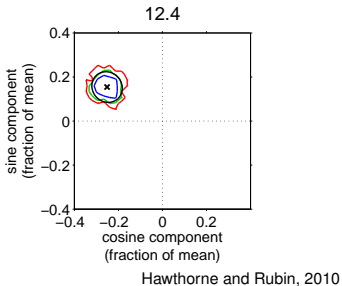
Larger tidal stress = faster slip, with some complications from the friction law

- More modulation in shallower portions of the steady state curve, at lower slip rates
- More modulation when there is enough slip in each period for state evolution, at higher slip rates

slip rate in bins behind the front



Effect of a tidal forcing



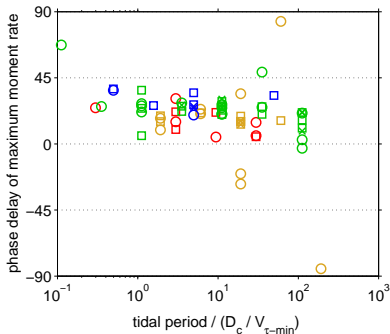
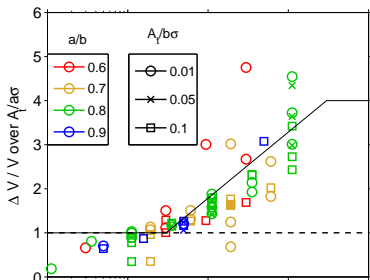
- Modulation of slip in Cascadia A_t : $\pm 30\%$
- Tidal stress on the interface: ± 0.5 to 1.5 kPa

$$\frac{\Delta V}{V} = 1-4 \frac{A_t}{a\sigma}$$

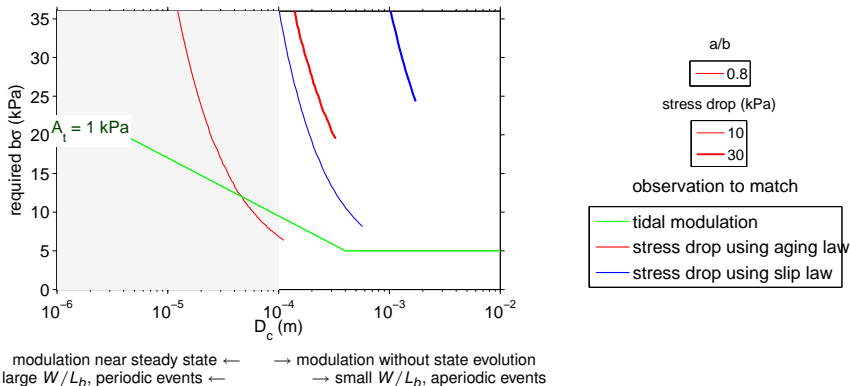
$$0.3 = 1-4 \frac{1 \text{ kPa}}{a\sigma}$$

Implies $a\sigma \approx 1 - 20$ kPa

For $a \approx 0.01$, $\sigma \approx 0.1 - 2$ MPa



Matching tidal forcing and recurrence interval simultaneously?

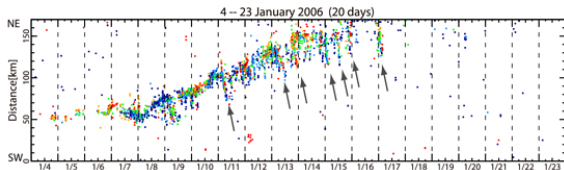
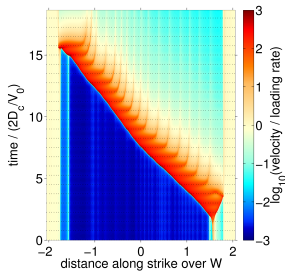


- Possible to match 10 kPa stress drop and 30% modulation with the aging law
- But not 30 kPa stress drop and 30% modulation
- More difficult to match the observations with the slip law
 - Slip law preferred by experiments for frictional energy dissipation estimates—for our stress drop estimates
- Lines terminate when they leave the regime that allows for periodic steadily propagating events

Conclusions to this point

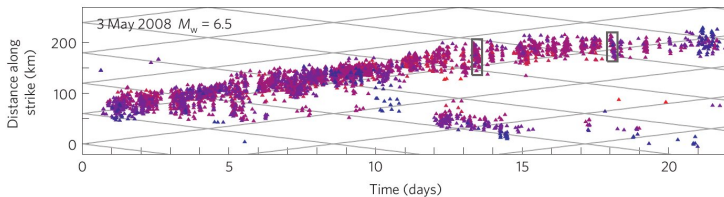
- Simulations exhibit many small events as well as periodic large events
- Large events propagate steadily “along strike” with the strip model
- Stress drops controlled by ability to propagate long distances along strike
 - Understood with an energy balance approach
- Tidal forcing results in a gradual modulation of the slip rate behind the propagating front
- Difficult to match the observed stress drops and tidal modulation with the slip law

Back-Propagating Fronts



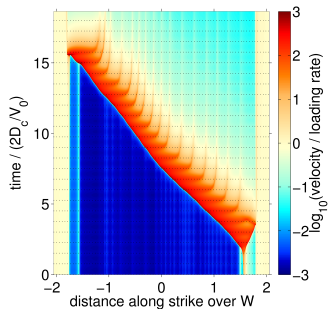
Obara et. al., 2012

Regions of high slip rate propagating back through the region that has already slipped



Houston et. al., 2011

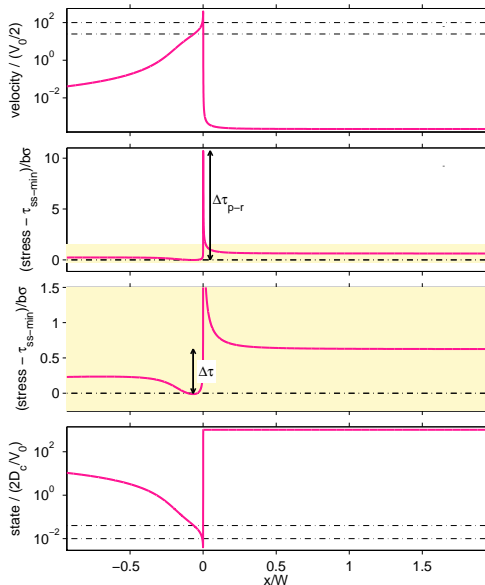
Back-Propagating Fronts



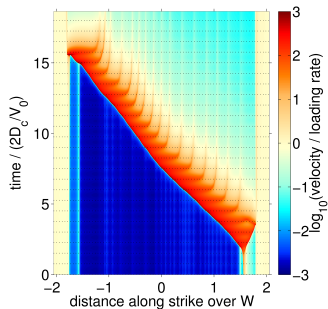
Driven by available stress drop

- From tidal forcing
- From naturally arising stress recovery

Similar to forward-propagating fronts, but with smaller stress drop and smaller initial state



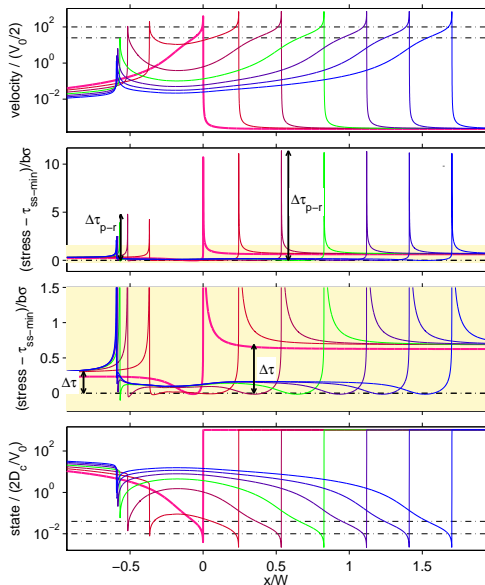
Back-Propagating Fronts



Driven by available stress drop

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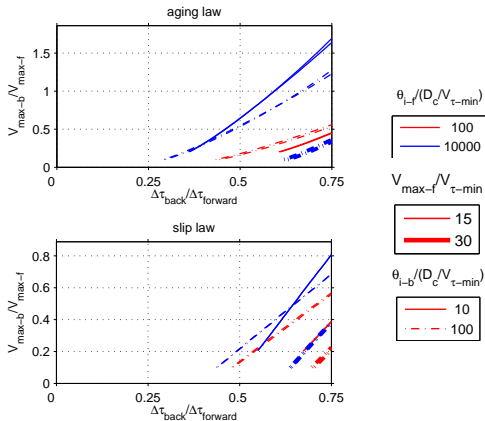
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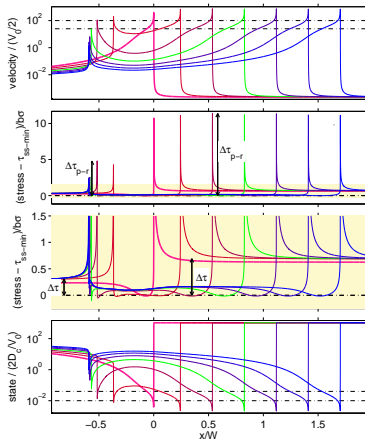
Back-Propagating Fronts: Maximum Slip Rates

strain energy release rate $G =$ energy dissipated by friction (fracture energy G_c)
 constant \times stress drop \times slip \approx near-tip contribution + gradual decay contribution

depends on $\Delta\tau$ not V_{\max} or θ_i depends on V_{\max}, θ_i , not $\Delta\tau$ depends on V_{\max} , not θ_i or $\Delta\tau$



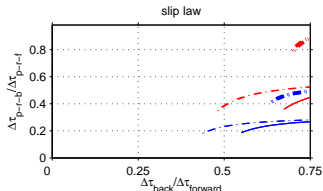
Back-propagating fronts have smaller $\Delta\tau$,
 but also smaller θ_i



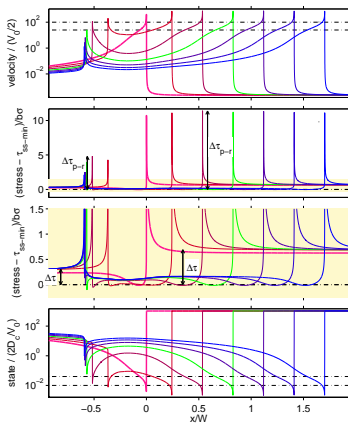
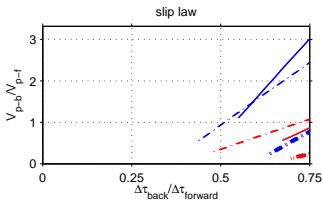
Back-Propagating Fronts: Propagation Rates

$$V_{\text{prop}} = (0.5 \text{ to } 0.65) V_{\text{max}} \frac{\text{shear modulus } \mu}{\Delta\tau_{p-r}}$$

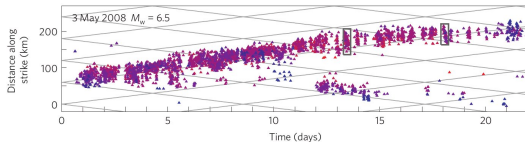
$\Delta\tau_{p-r}$ is smaller for the back-propagating fronts:



But not by enough to match the observations:



Observed propagation around 10 times forward rate.



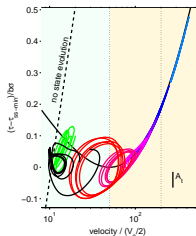
Conclusions 1

- Simulations exhibit many small events as well as periodic large events
- Large events propagate steadily “along strike” with the strip model
- Stress drops controlled by ability to propagate long distances along strike
 - Understood with an energy balance approach
- Tidal forcing results in a gradual modulation of the slip rate behind the propagating front
- Difficult to match the observed stress drops and tidal modulation with the slip law
- Heterogeneity in moment rate smaller than in observed events
- **Back-propagating fronts arise**
 - **Can understand their maximum and propagation velocities with an energy balance approach**
 - **Propagate too slowly to match the observed fronts**

Conclusions 2: Shortcomings of the model

Can we throw out this constitutive law as an explanation for slow slip?

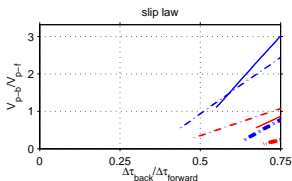
Problem 1: Matching amplitude of modulation and stress drops



Modulation depends strongly on steady state stress: choose a different curve?

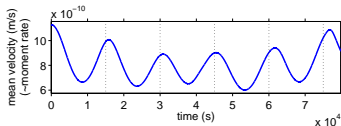
- Probably acceptable for experiments
- This curve has theoretical support
- A flatter curve could mean large changes in propagation velocity

Problem 2: Lack of large variation in moment rate, slow back-propagating fronts

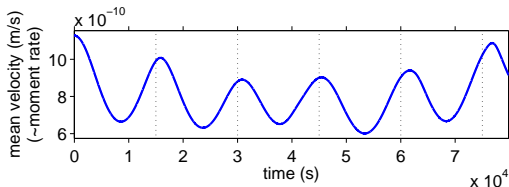


Include heterogeneity in friction parameters, normal stress?

- Need to preserve constant along-strike propagation
- Would likely influence tidal modulation
- Fast back-propagating fronts require stress drops comparable to the overall stress drop

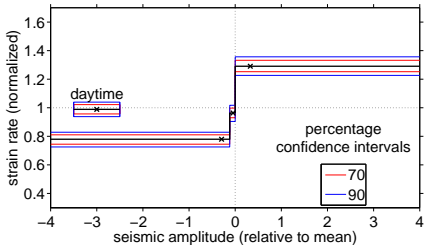


Matching heterogeneity



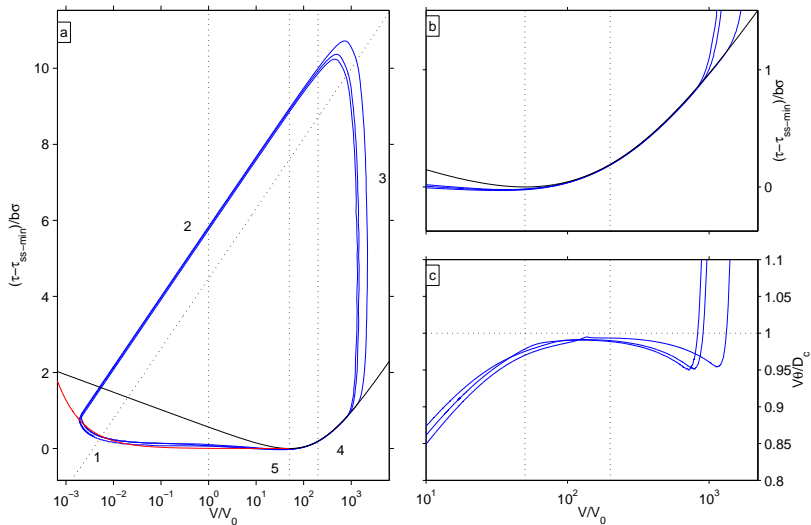
Models don't have enough variability on non-tidal timescales.

- Need heterogeneity?
- Or the friction law is incorrect?

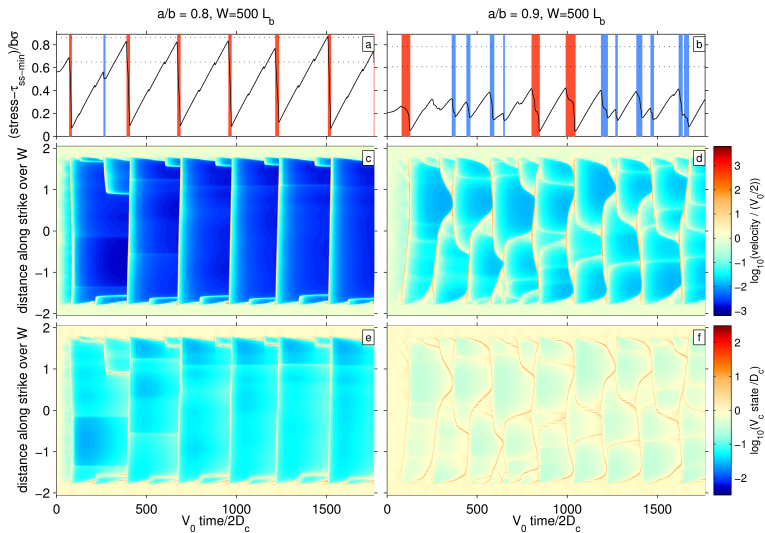


- at least $\pm 30\%$ variation in moment rate at a range of periods shorter than 4 hours
- from correlations between tremor and strain rate

Evolution of stress and velocity



Evolution of stress and velocity



Effect of a tidal forcing

- More modulation in shallower portions of the steady state curve, at lower slip rates
- More modulation when there is enough slip in each period for state evolution, at higher slip rates

