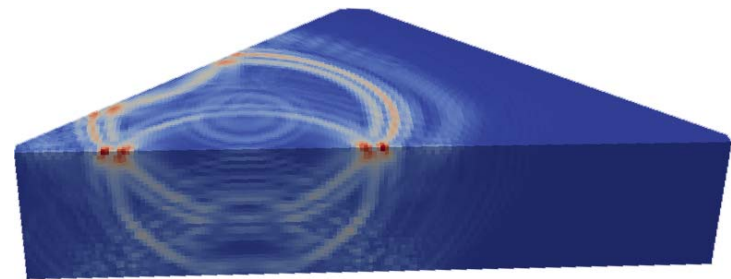
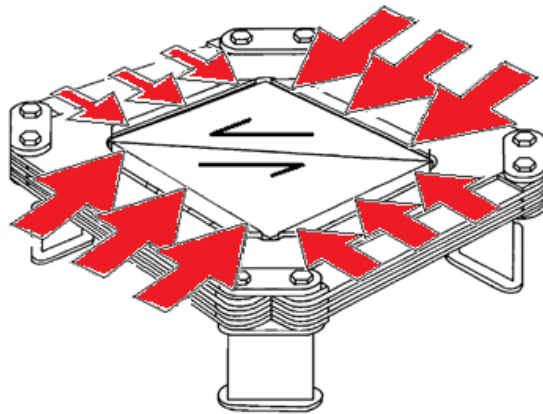


The nucleation of a laboratory earthquake: Implications for foreshocks and minimum earthquake size

Greg McLaskey

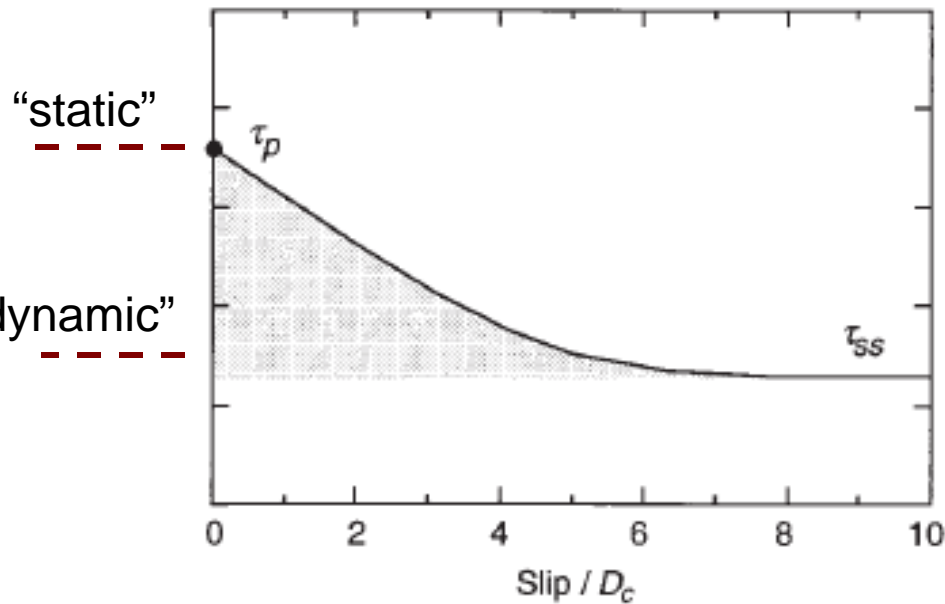
USGS Menlo Park

June 22, 2012

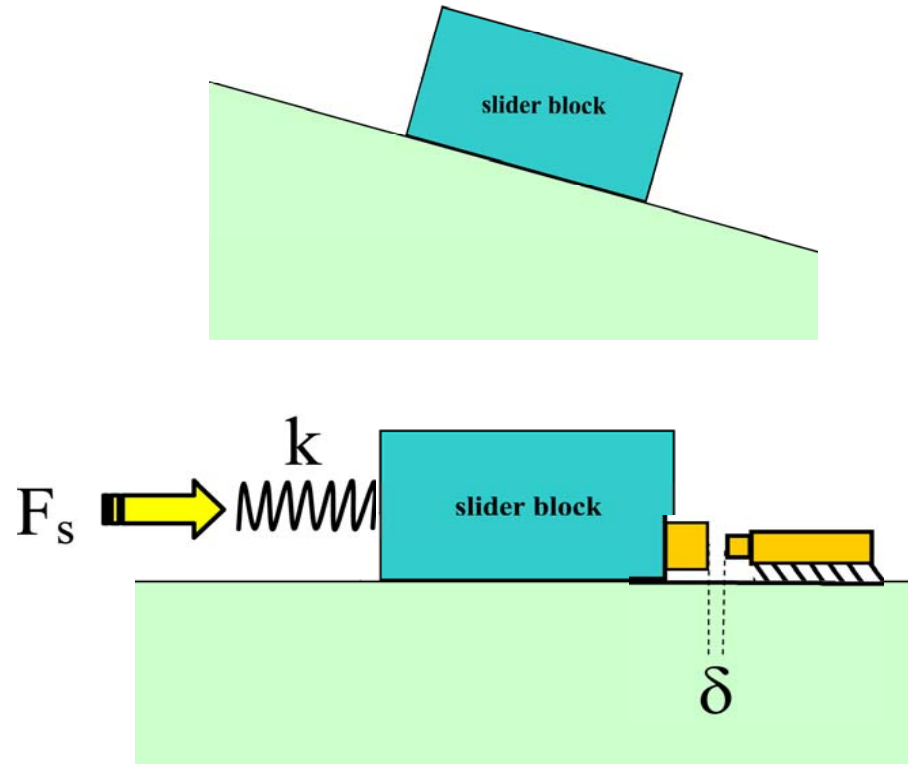


Brian Kilgore, Nick Beeler, Dave Lockner

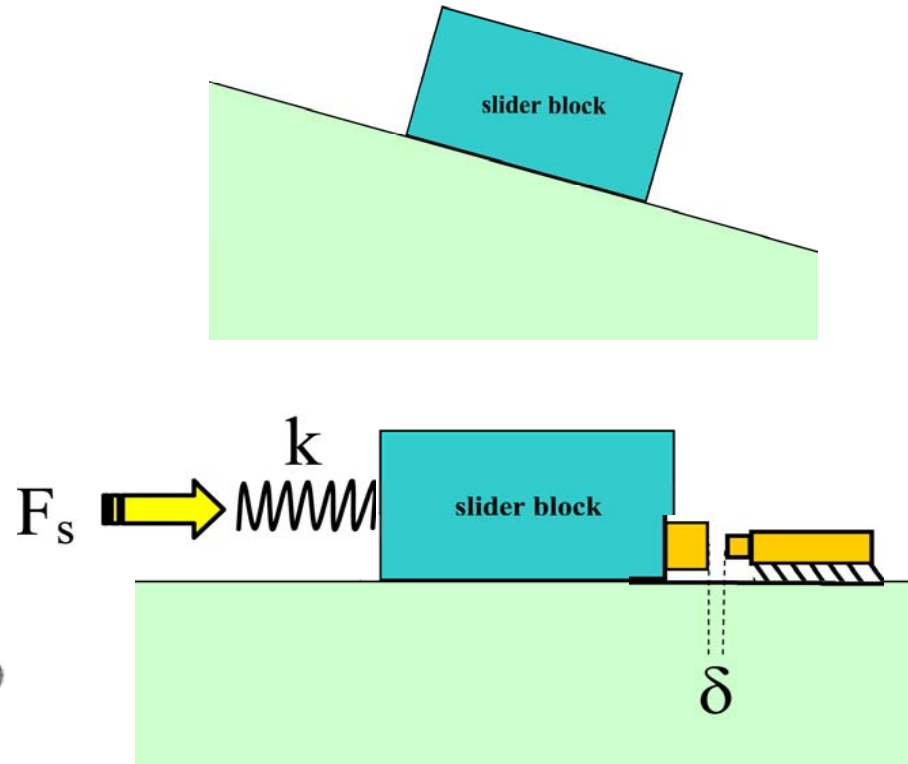
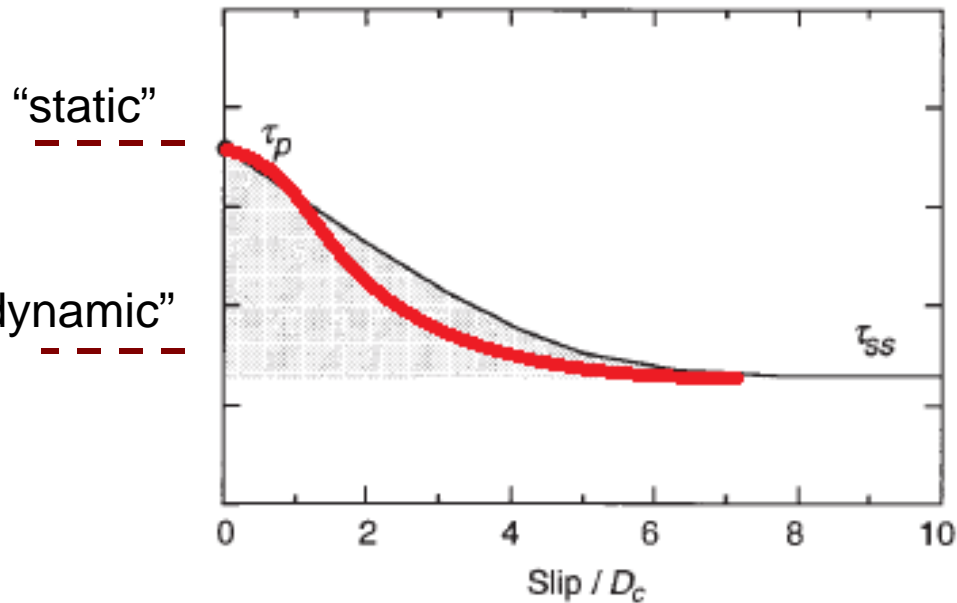
Slip weakening and stability



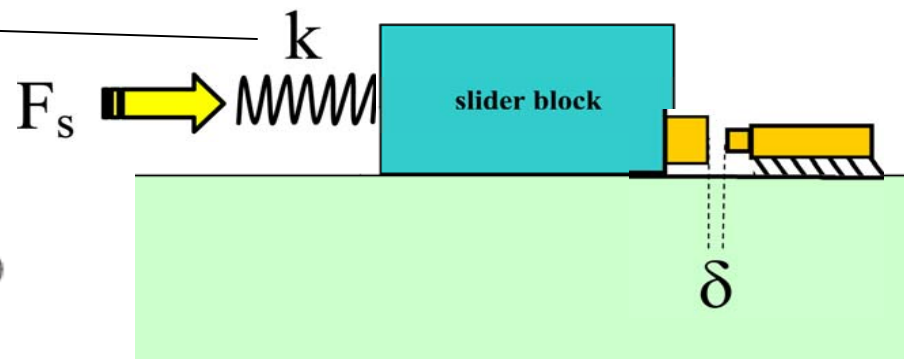
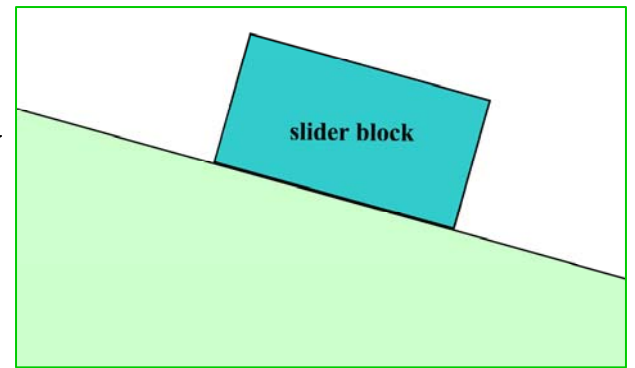
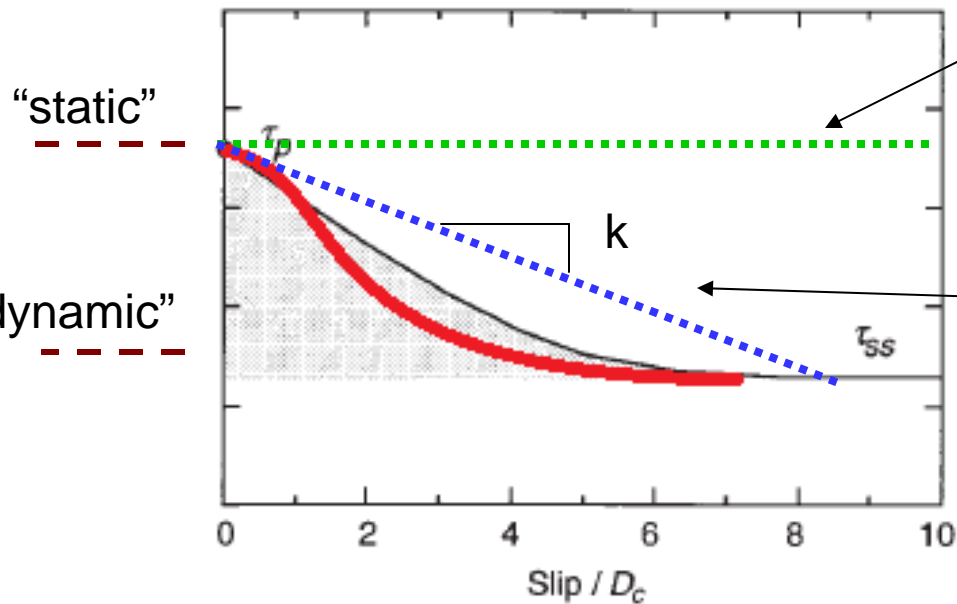
Dieterich and Kilgore 1996



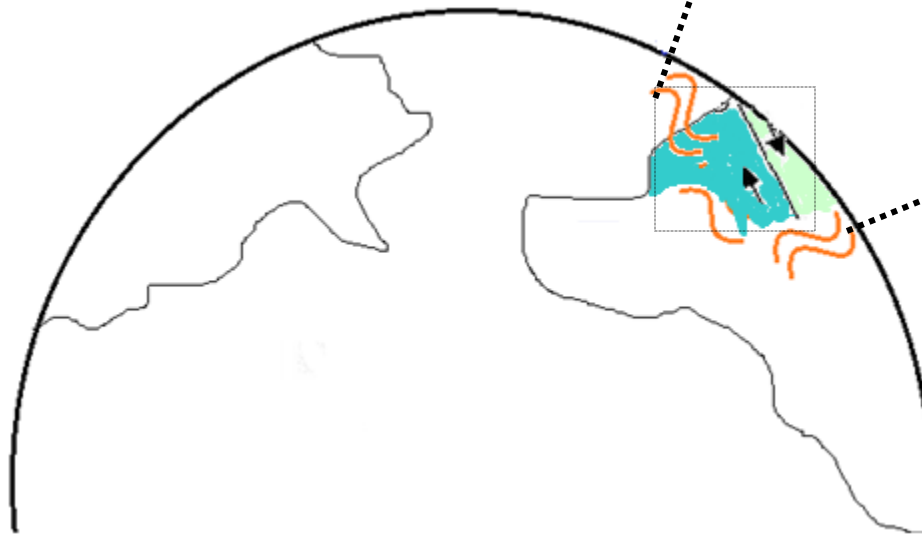
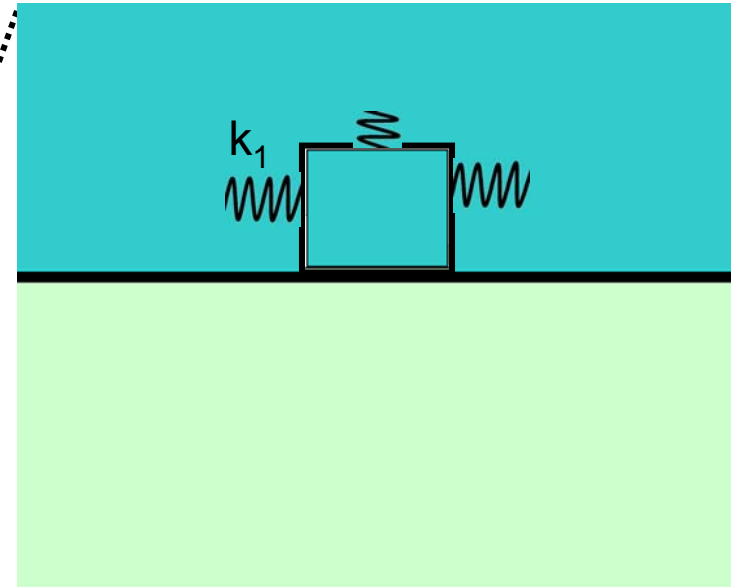
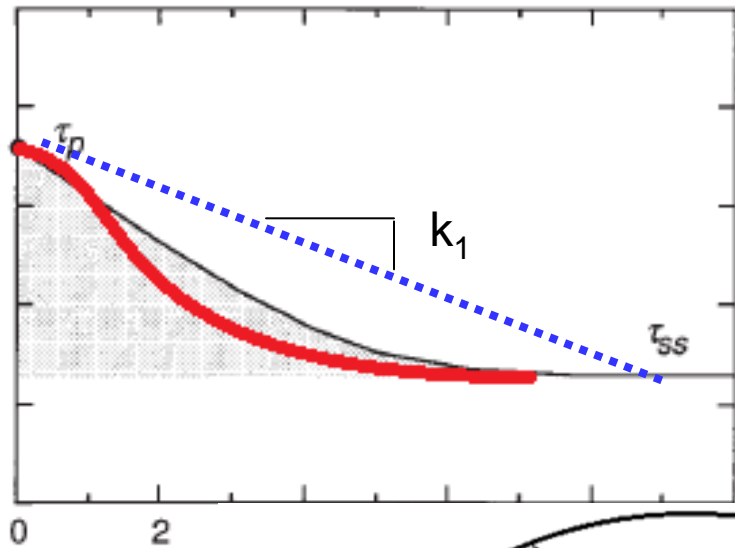
Slip weakening and stability



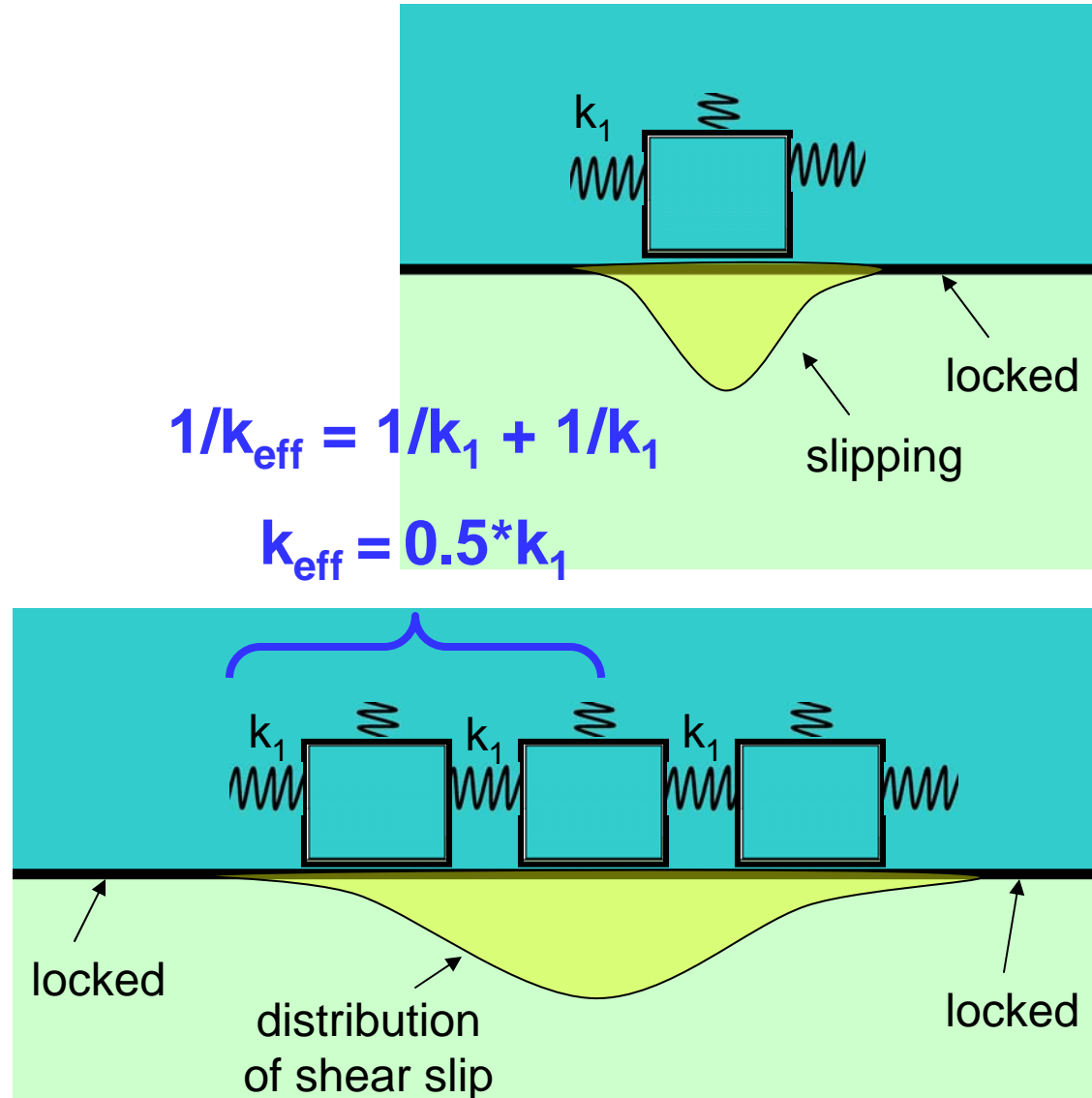
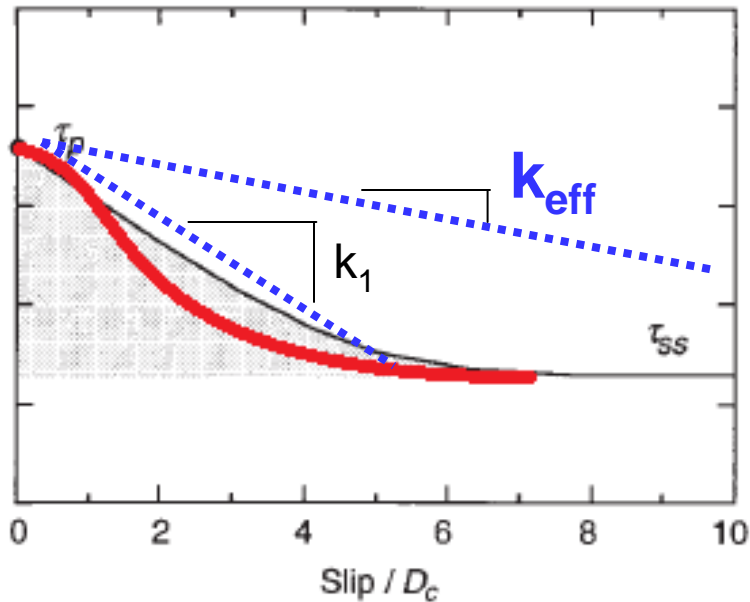
Slip weakening vs. unloading curve



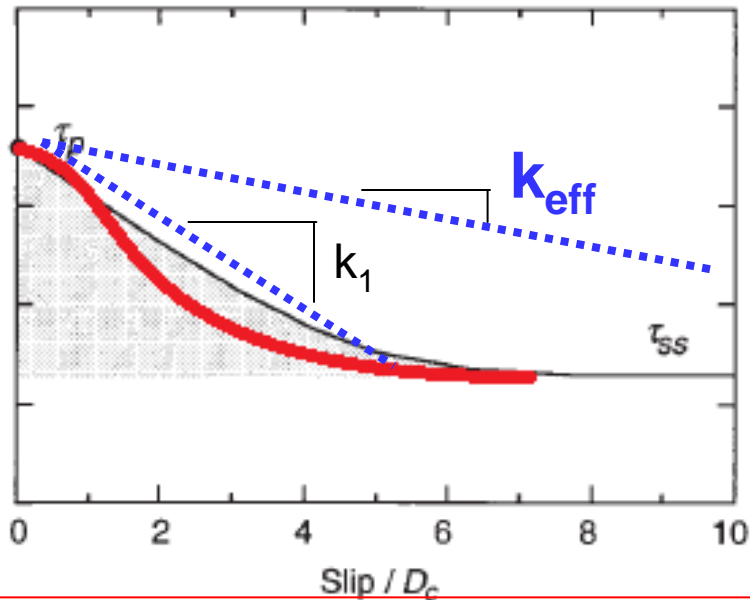
Who cares about slider blocks?



Stability of a slipping region in an elastic space



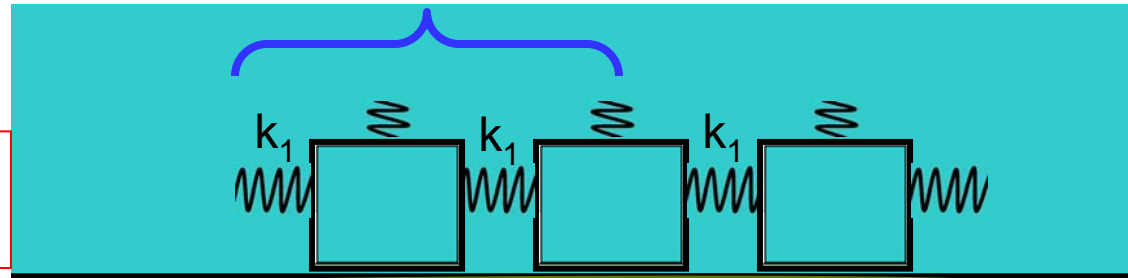
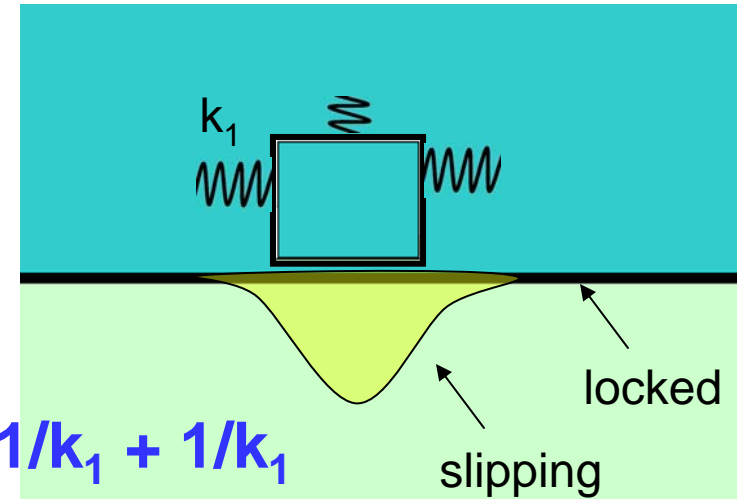
Stability of a slipping region in an elastic space



Assumption: fault strength is slip weakening with no discontinuities.

$$1/k_{eff} = 1/k_1 + 1/k_1$$

$$k_{eff} = 0.5 * k_1$$



- 1) there must be a minimum size of slipping patch below which slip is stable
- 2) even for patches that are large enough to host unstable slip, slip accelerates

Motivating questions

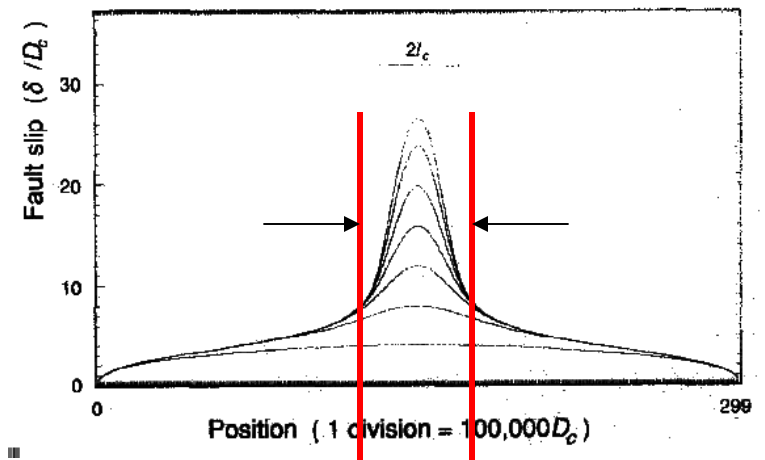
- Better understand this transition from aseismic to seismic
 - More effective use of initial P-waves can be more efficiently utilized for earthquake early warning systems
- Better physical interpretation of microseismicity
 - Intelligently interpret observations of small earthquakes for short-term earthquake forecasting
 - Are small earthquakes indicators of the nucleation of a larger and potentially-damaging earthquake?

Earthquake nucleation models

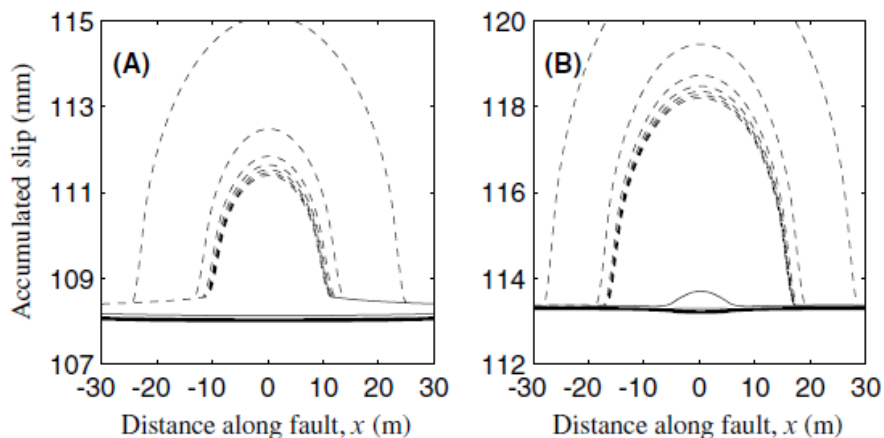
Rapid slip is ALWAYS preceded by slow and accelerating slip in a nucleation zone

This nucleation zone is $\sim 1\text{m} - 10\text{m}$?
depending on parameters

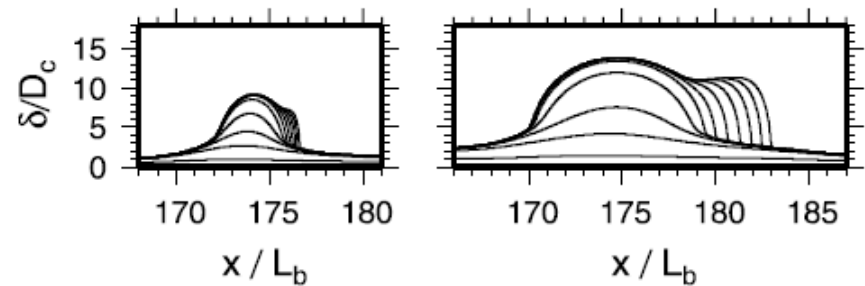
Dieterich 1992



Kaneko and Lapusta 2008

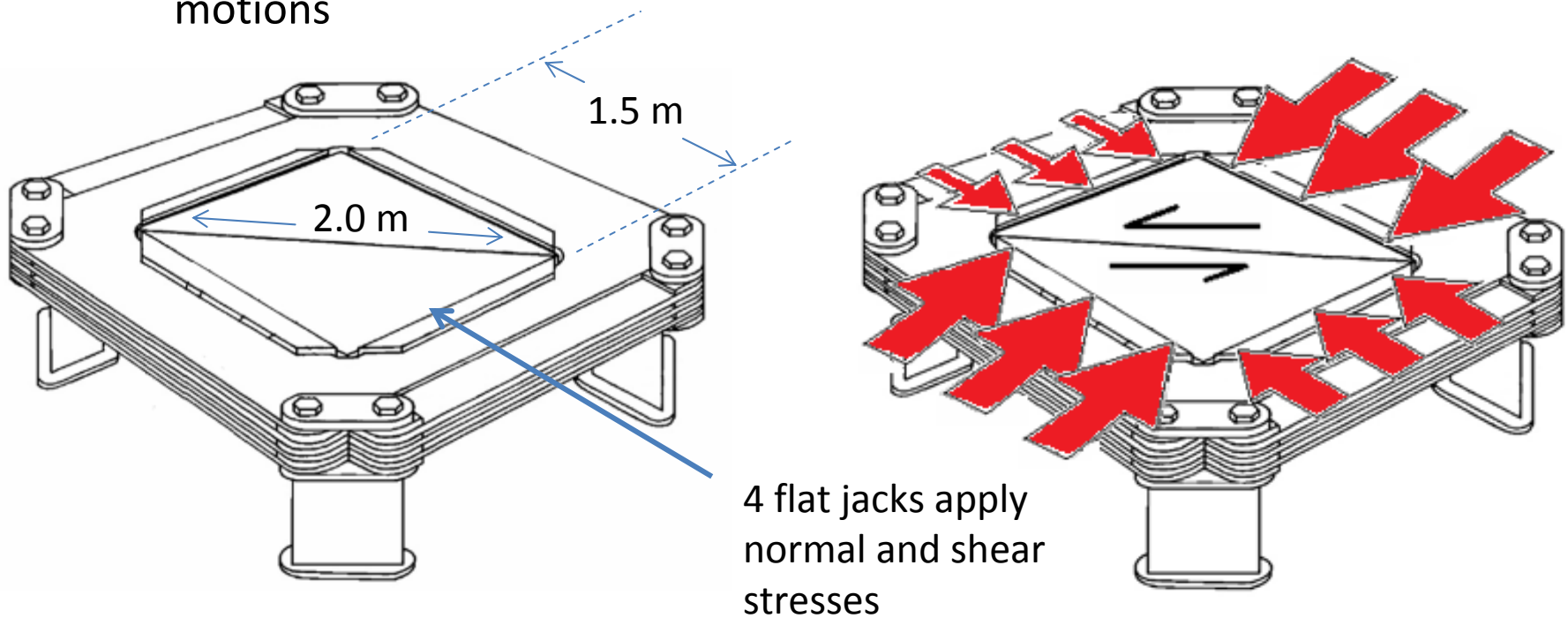


Ampuero and Rubin 2008

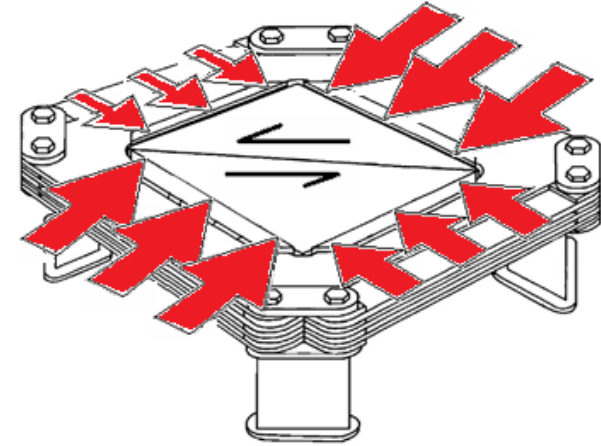


Laboratory Earthquakes

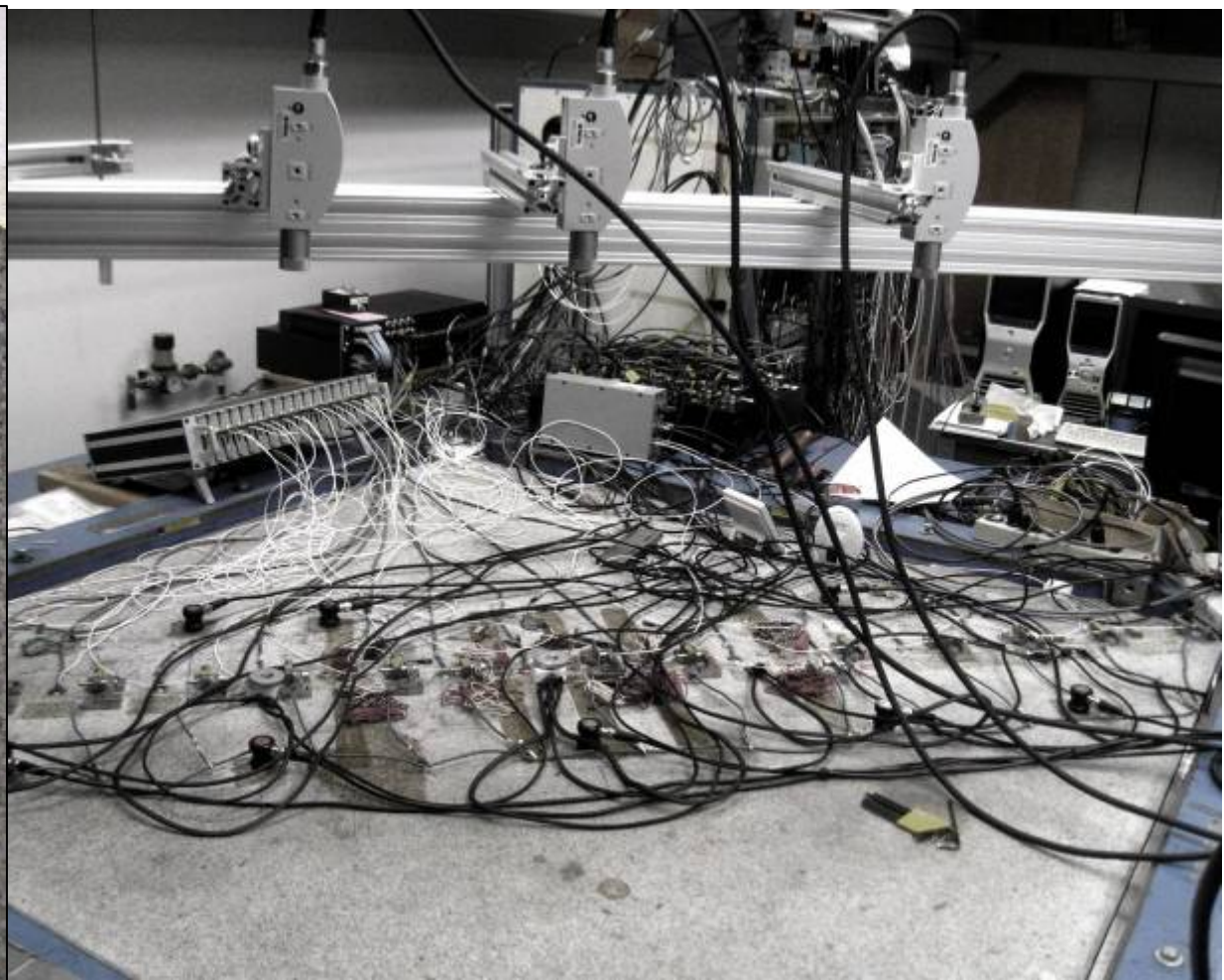
- 2 meter biaxial press
- Fault is large enough that one part of it can slip while another part remains locked
 - We can nucleate earthquakes, measure local stress state along the fault, observe dynamic rupture, and measure high frequency ground motions



2 meter biaxial press

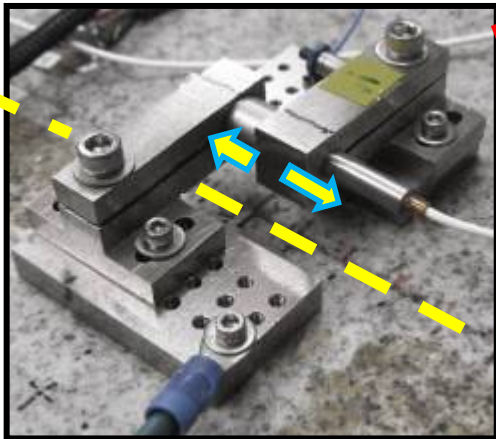


Laboratory Fault:
Sierra White Granite



2 meter biaxial press: instrumentation

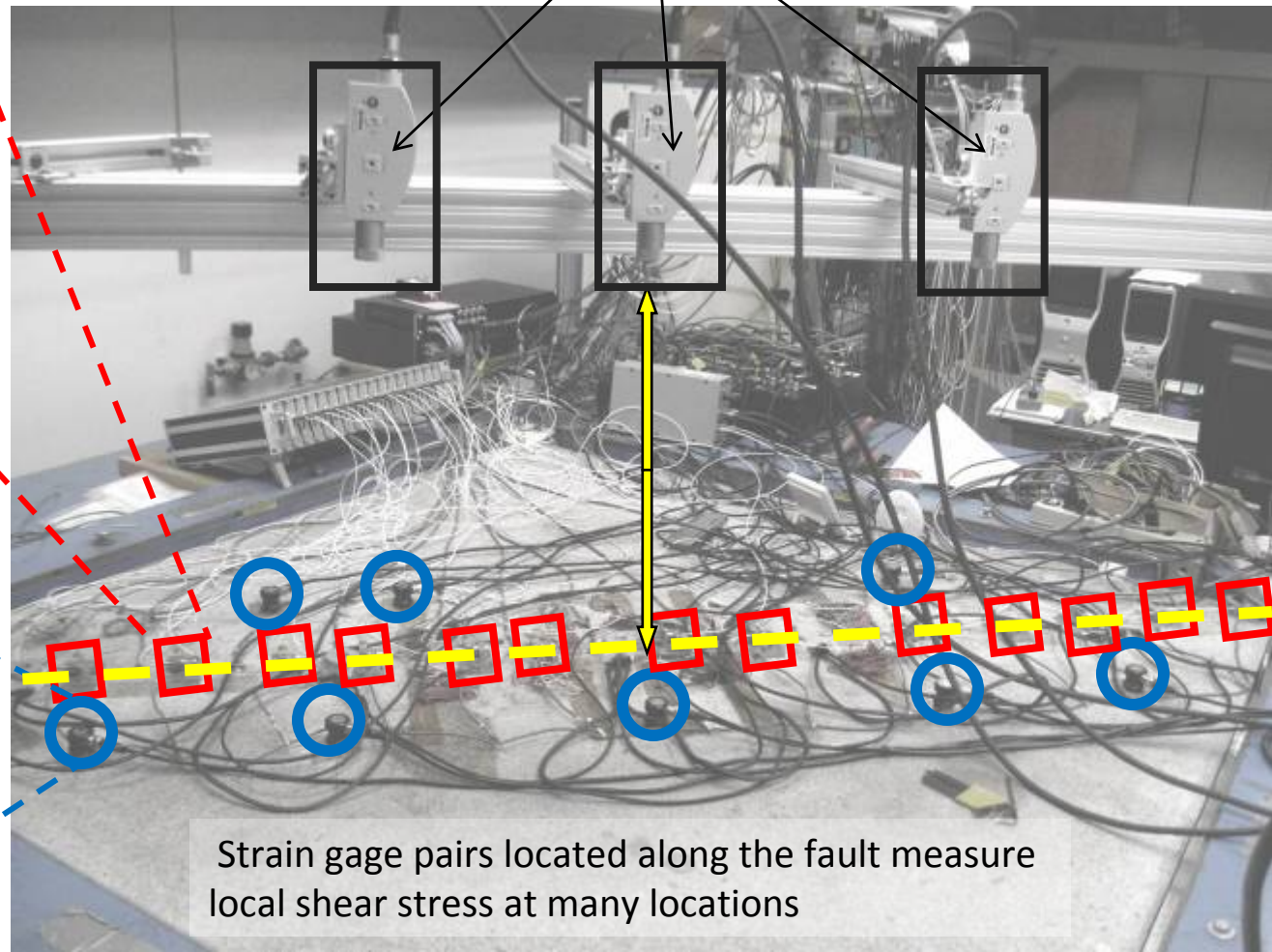
Capacitive slip sensor
(measures fault slip)



Piezoelectric sensor
(detects surface normal motion)

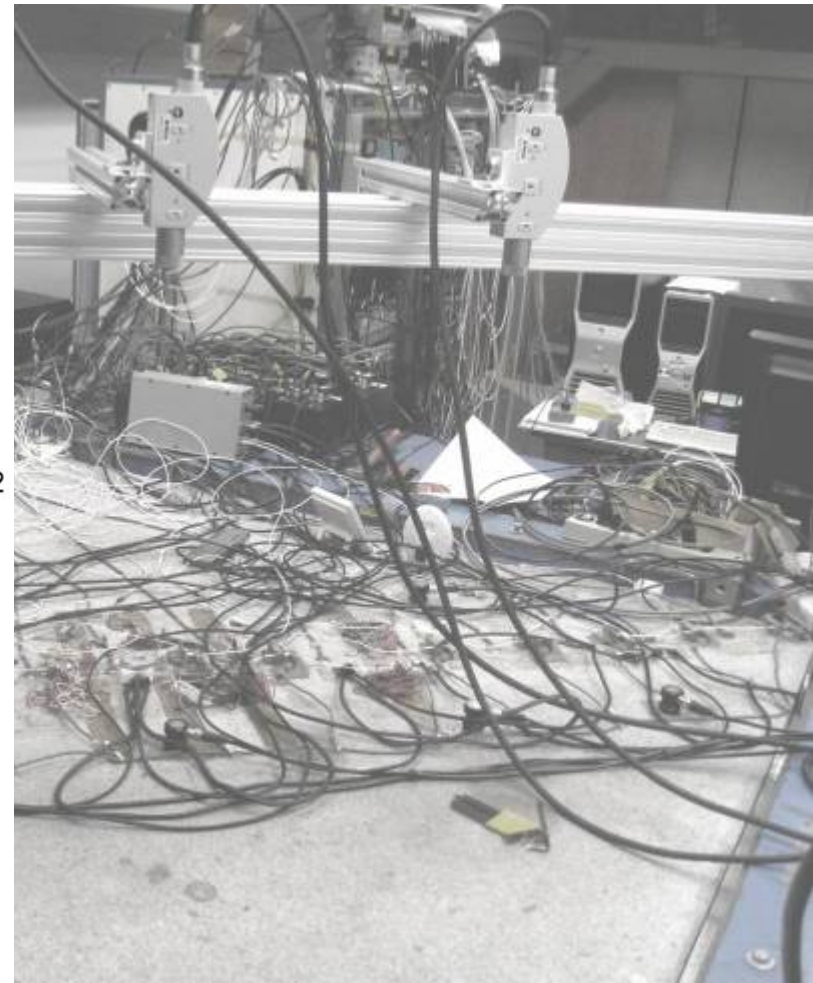
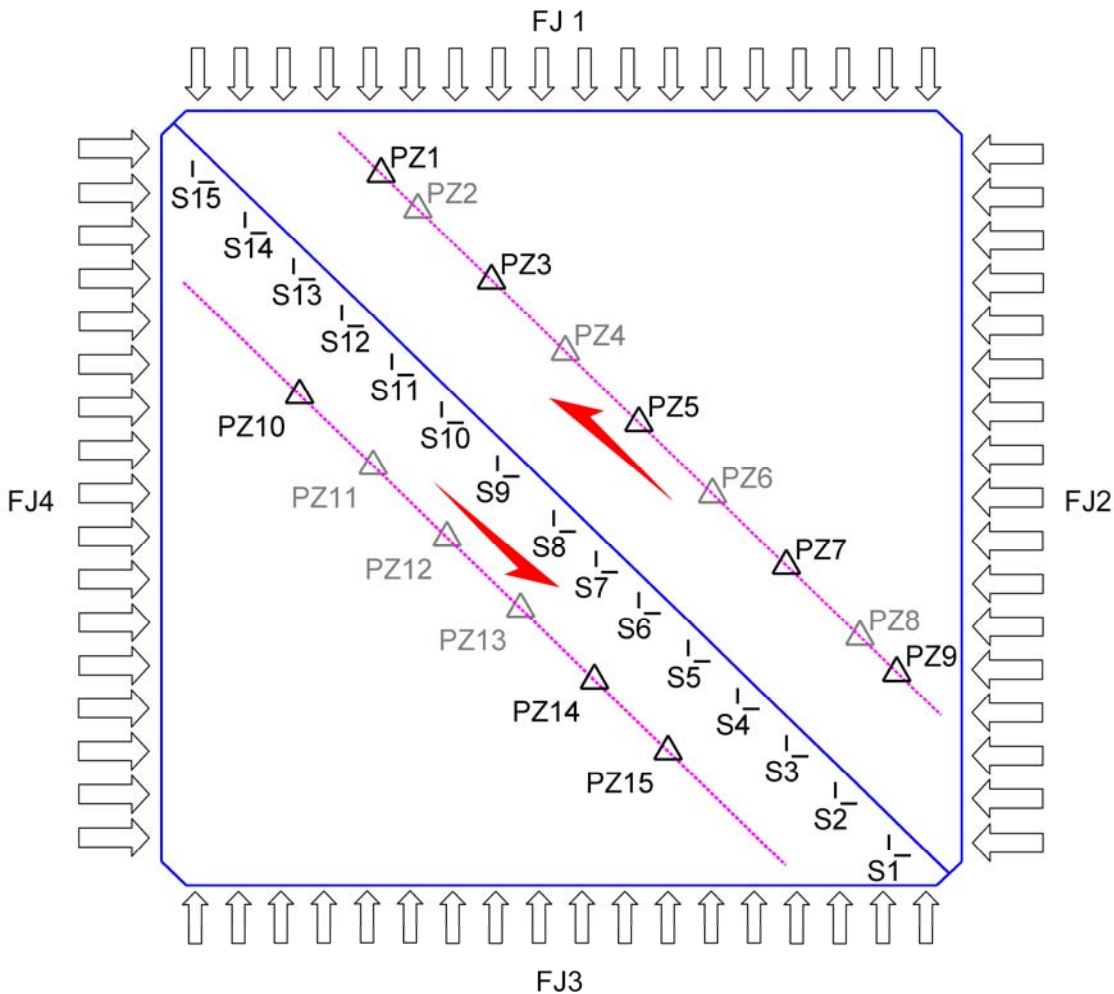
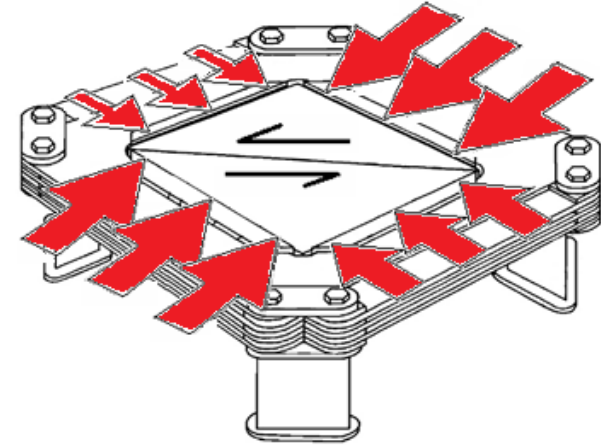


Laser vibrometer
(measures surface normal velocity)

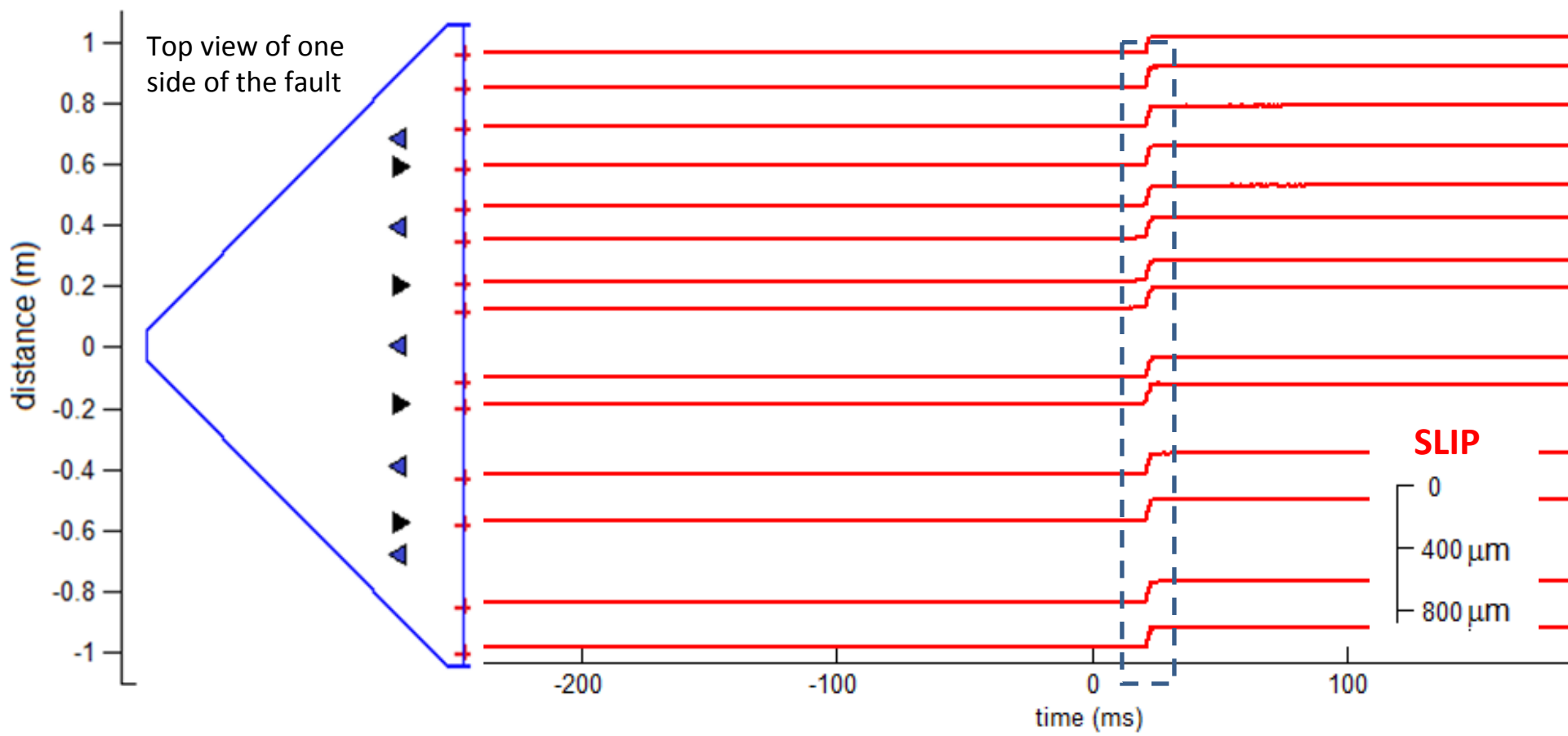
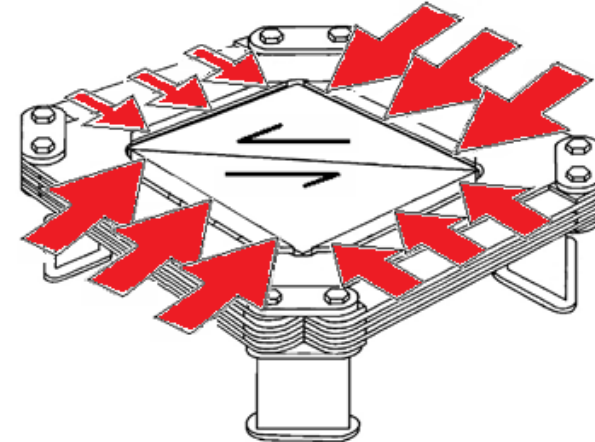


Strain gage pairs located along the fault measure
local shear stress at many locations

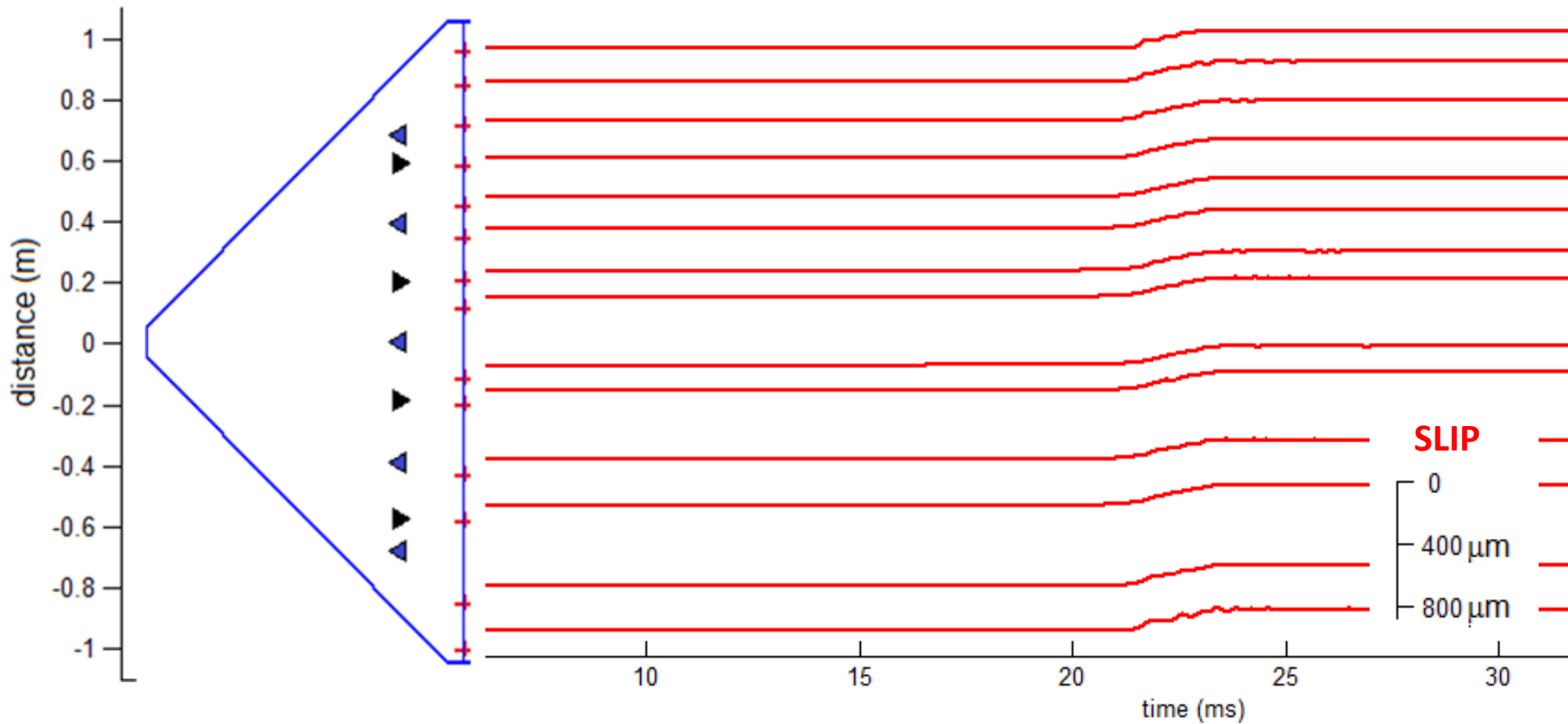
2 meter biaxial press: instrumentation



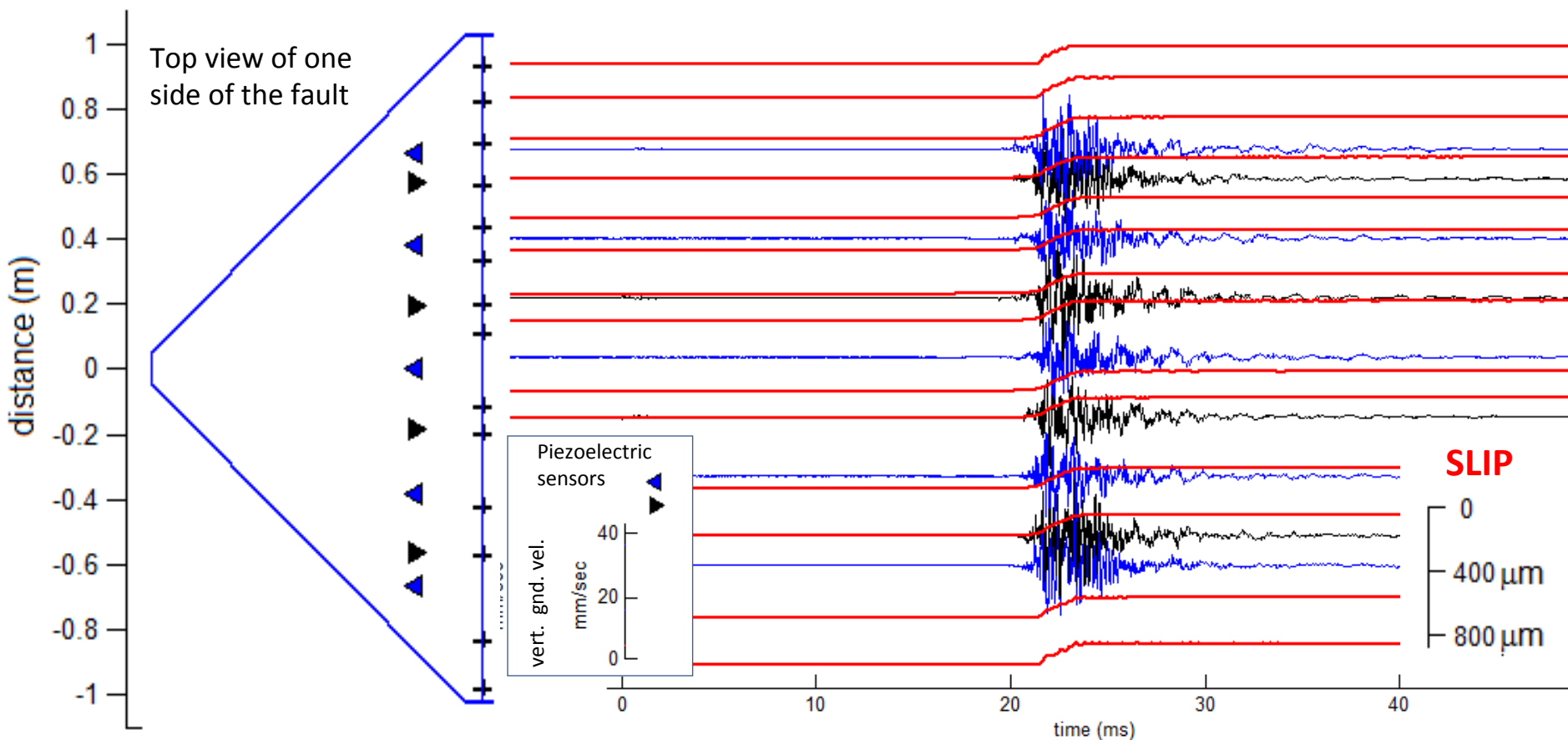
Stick slip



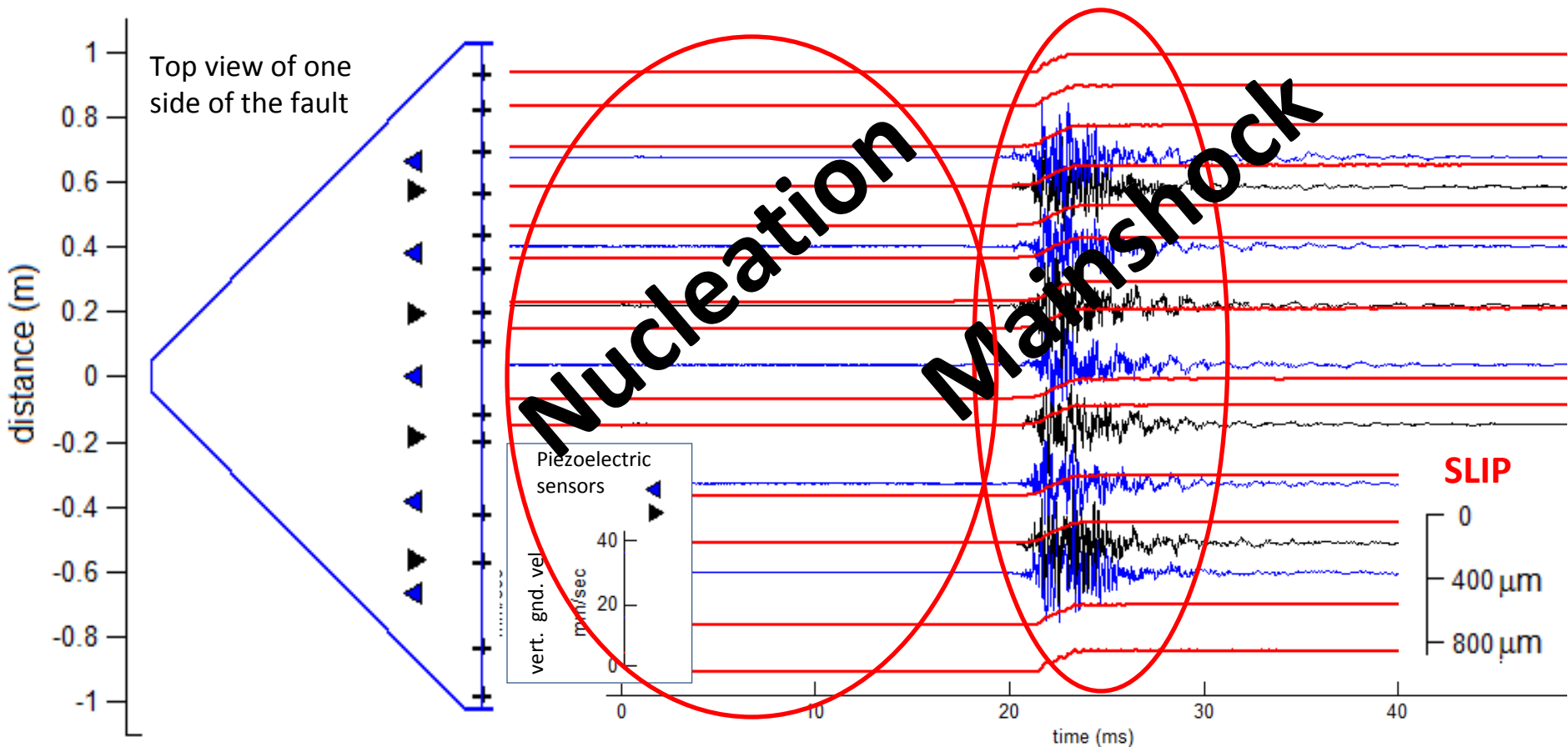
2 meter biaxial press



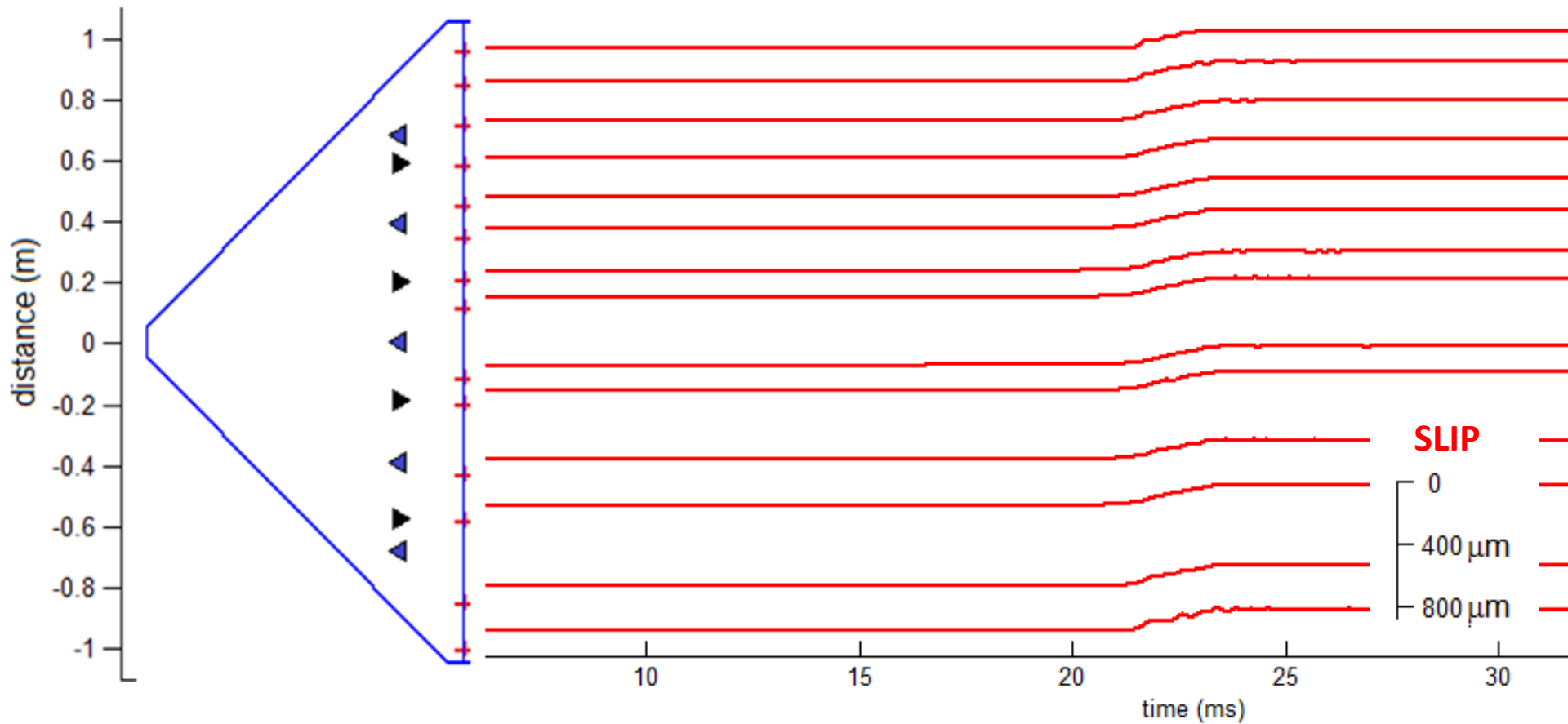
2 meter biaxial press: stick slip and laboratory earthquake



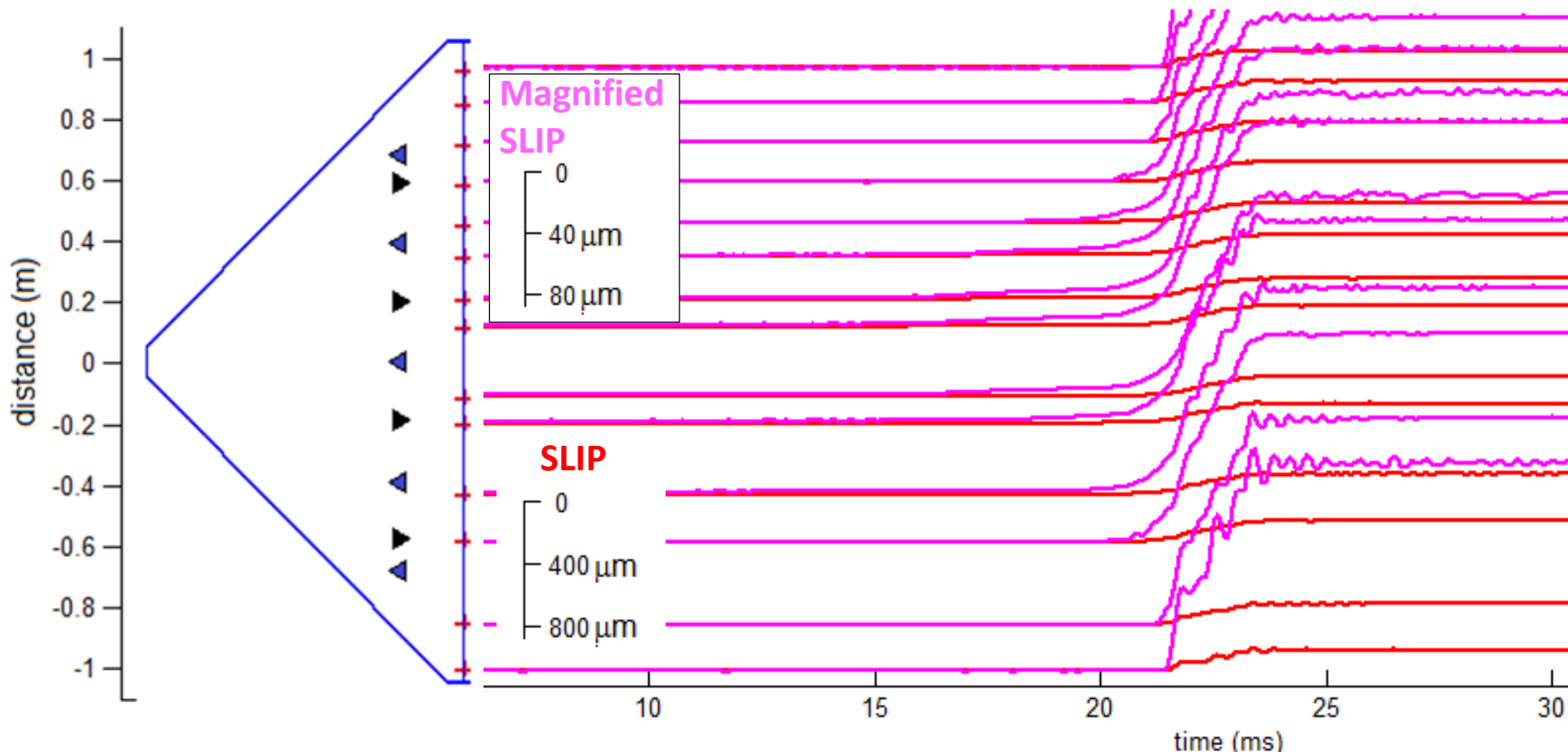
2 meter biaxial press: stick slip and laboratory earthquake



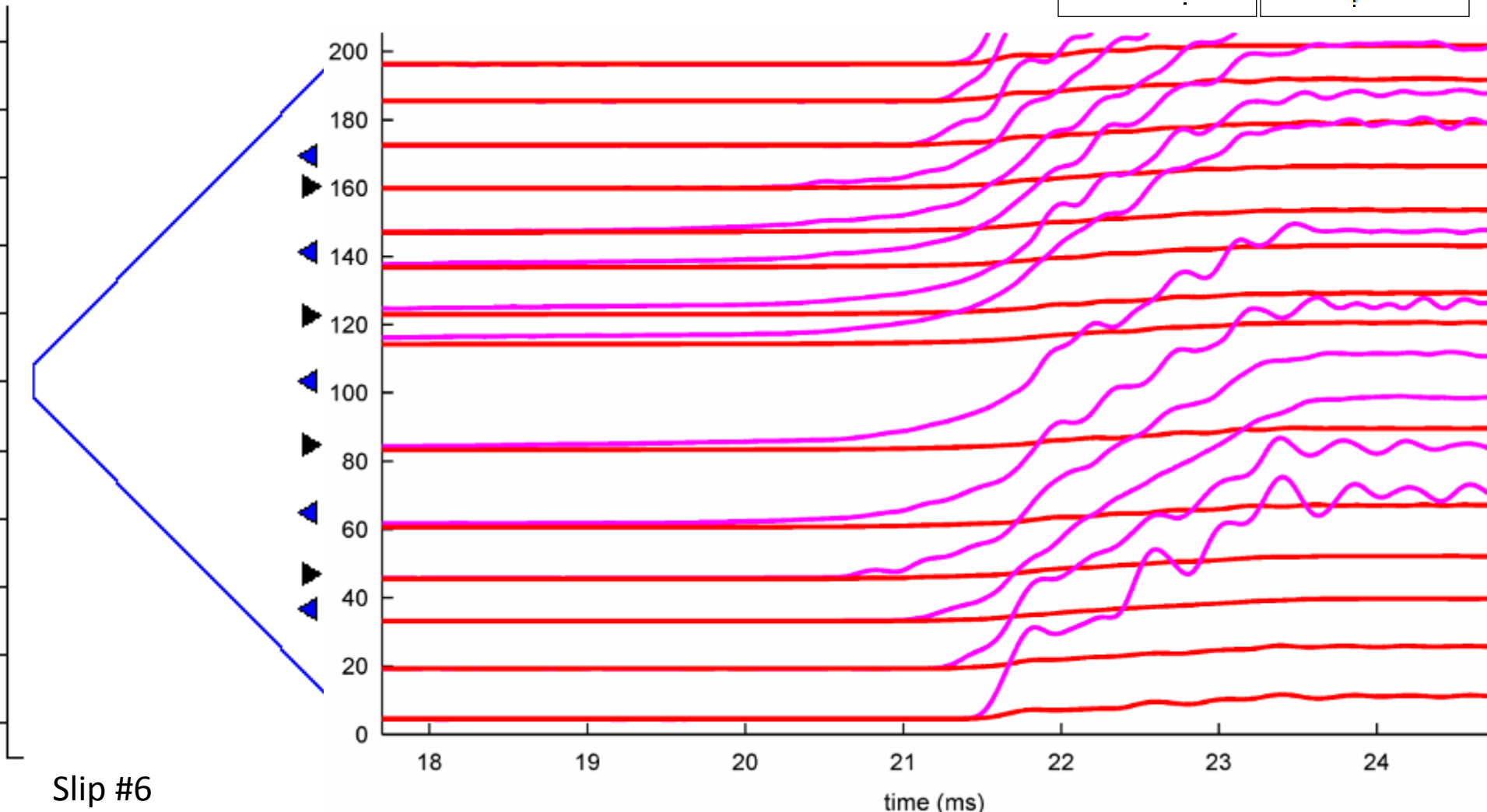
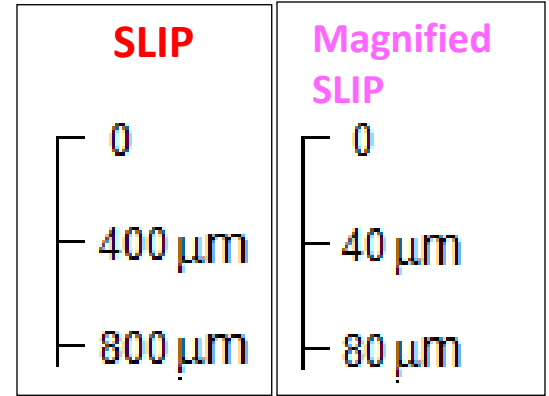
2 meter biaxial press



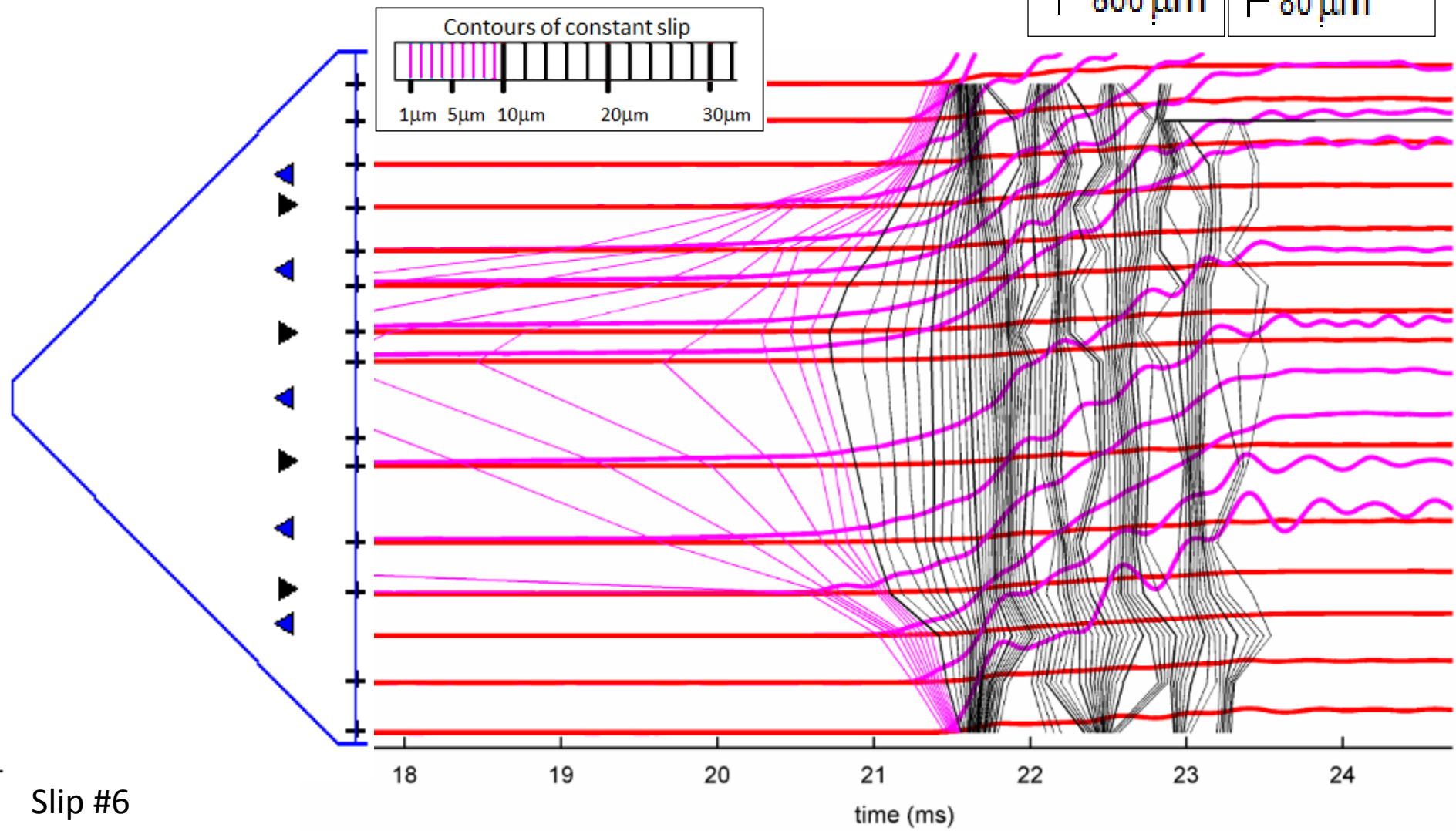
2 meter biaxial press



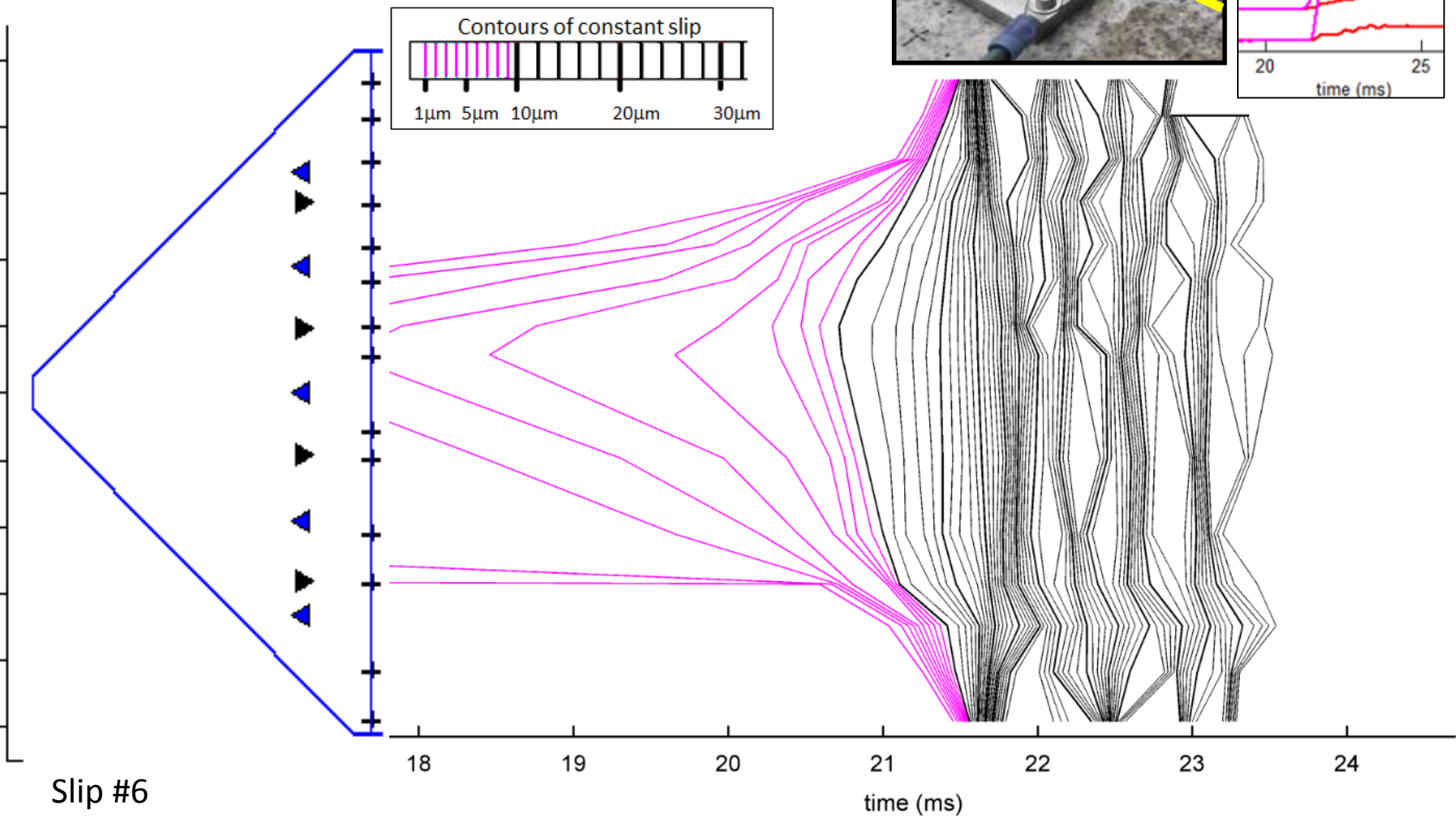
Mainshock



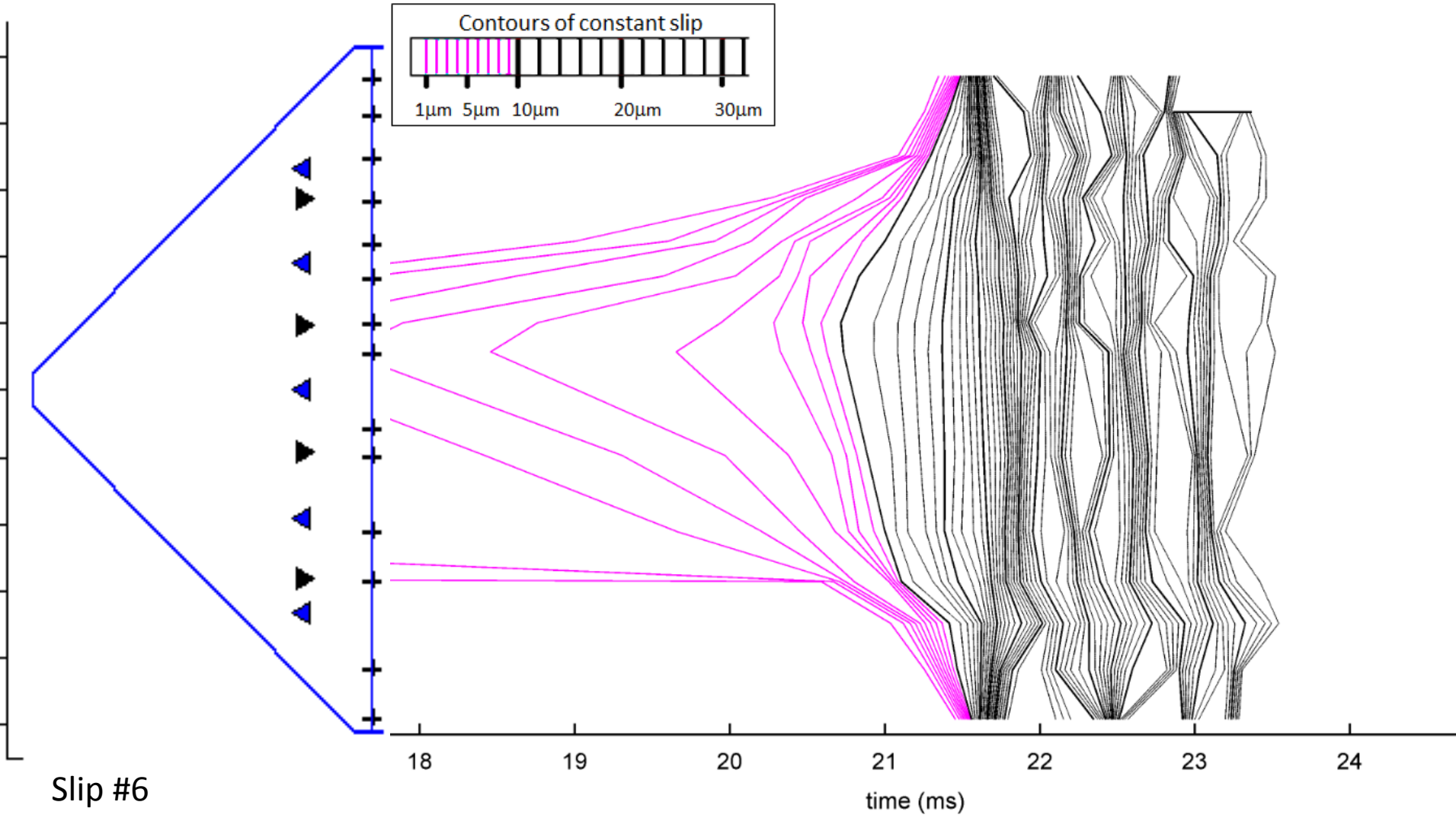
Mainshock



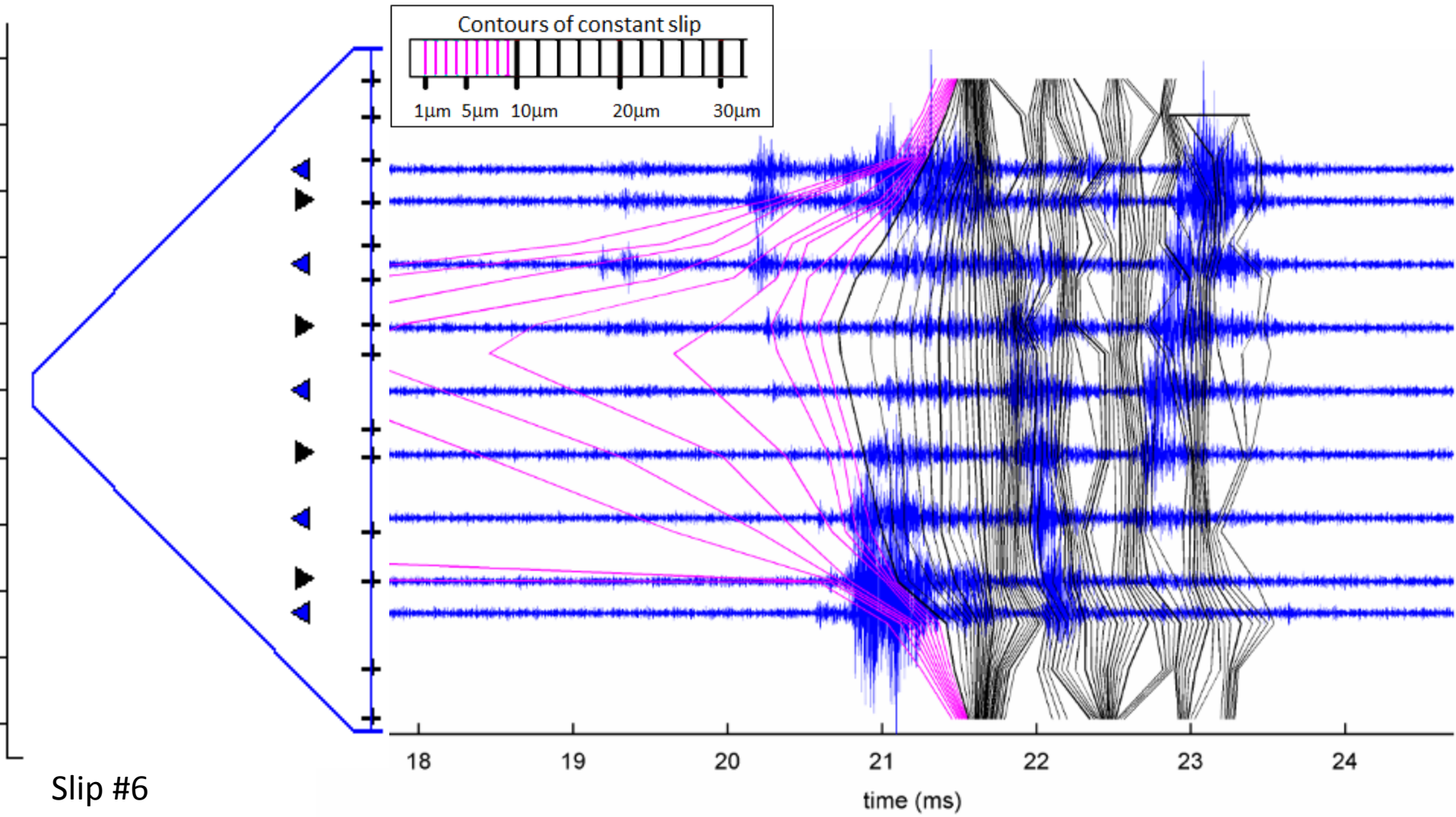
Mainshock



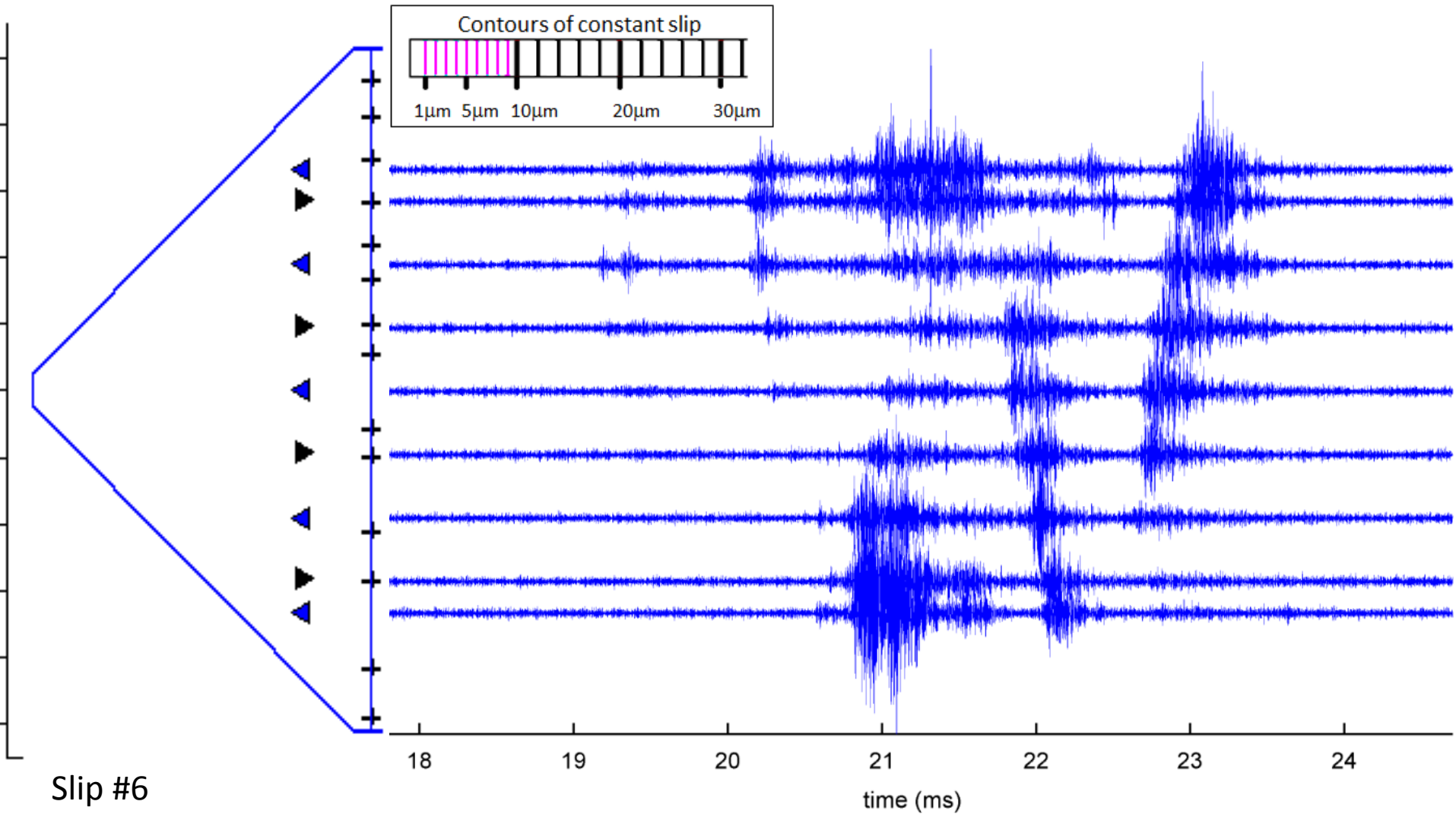
Mainshock



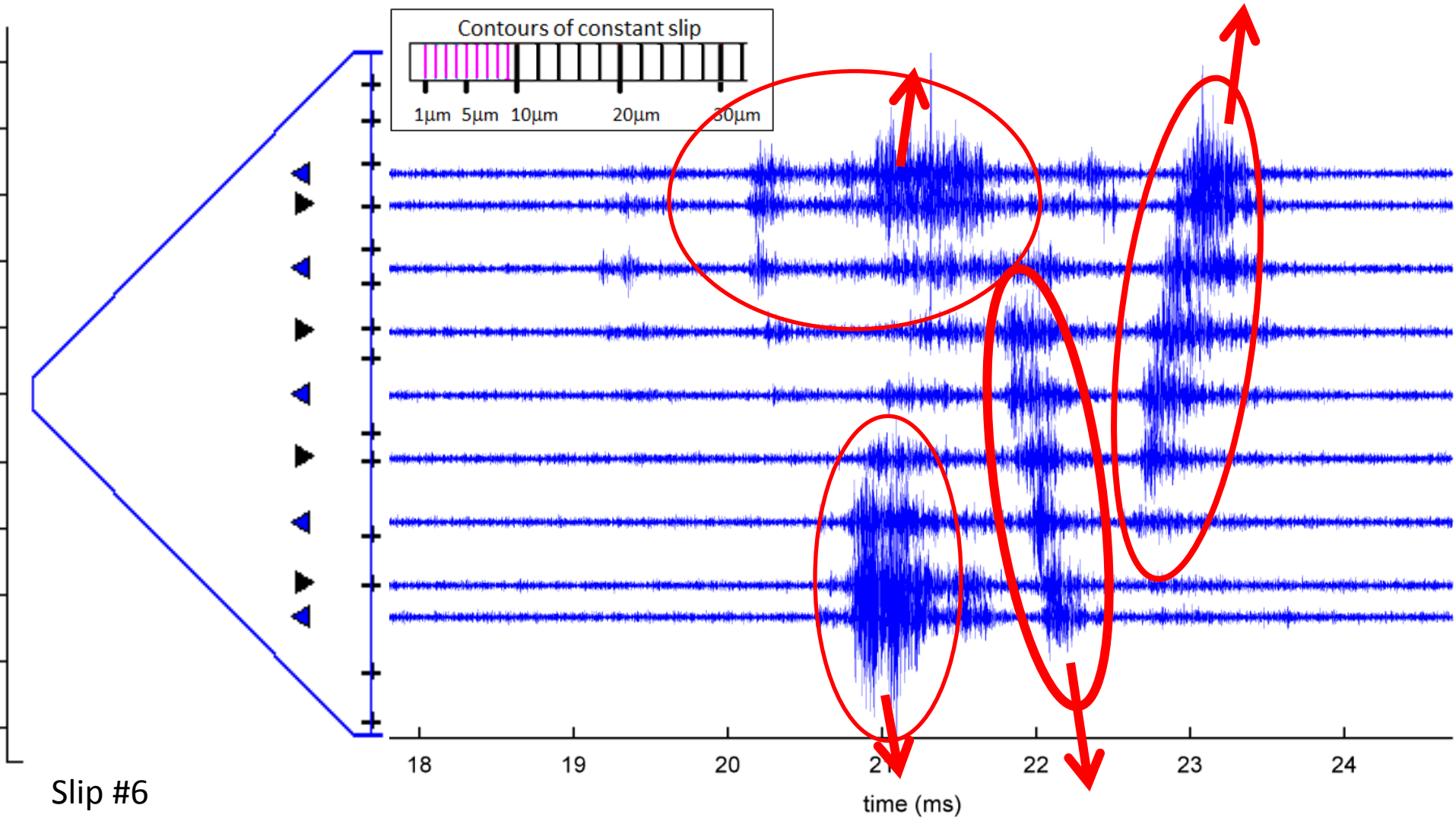
Mainshock



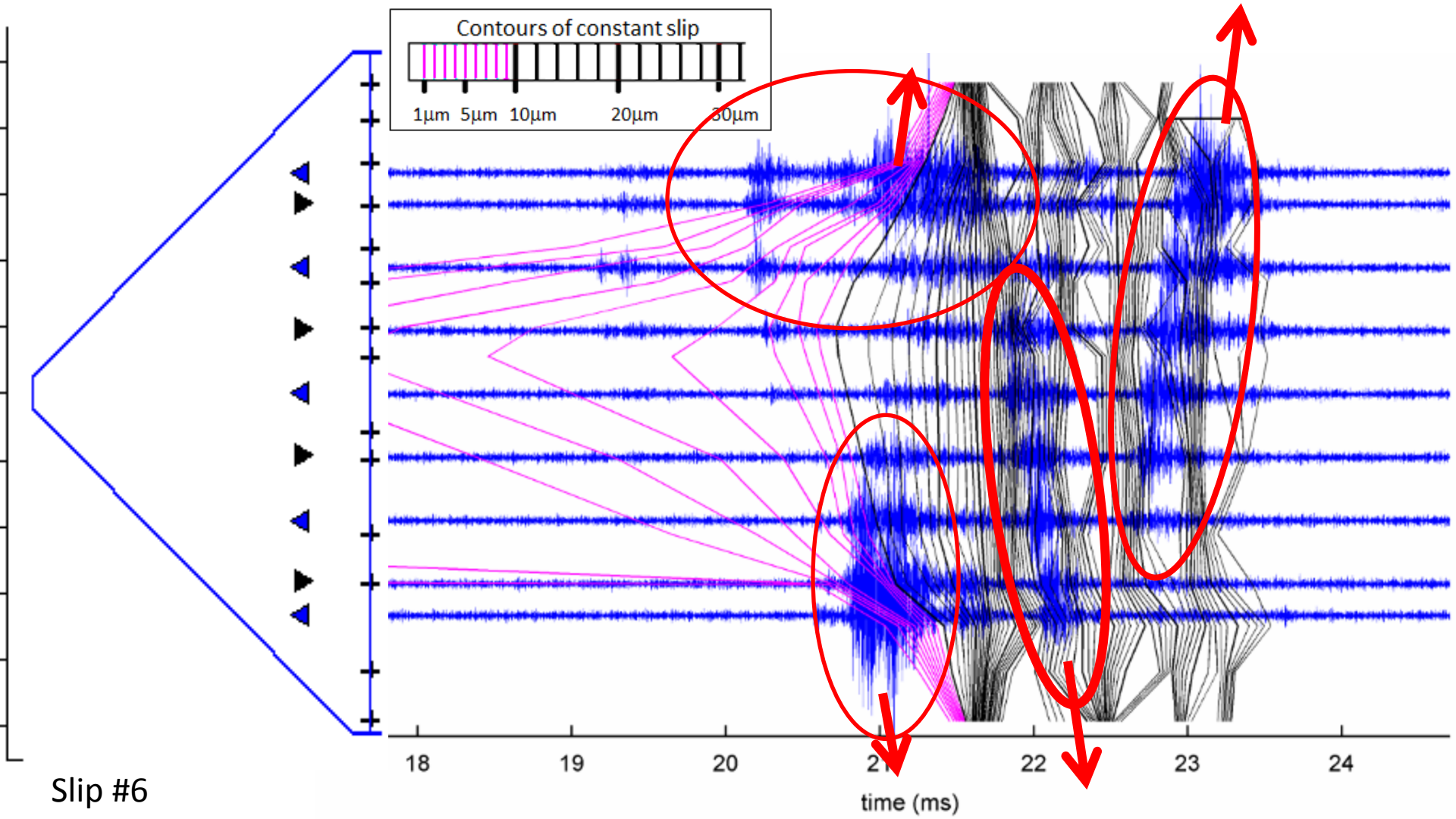
Mainshock



Mainshock

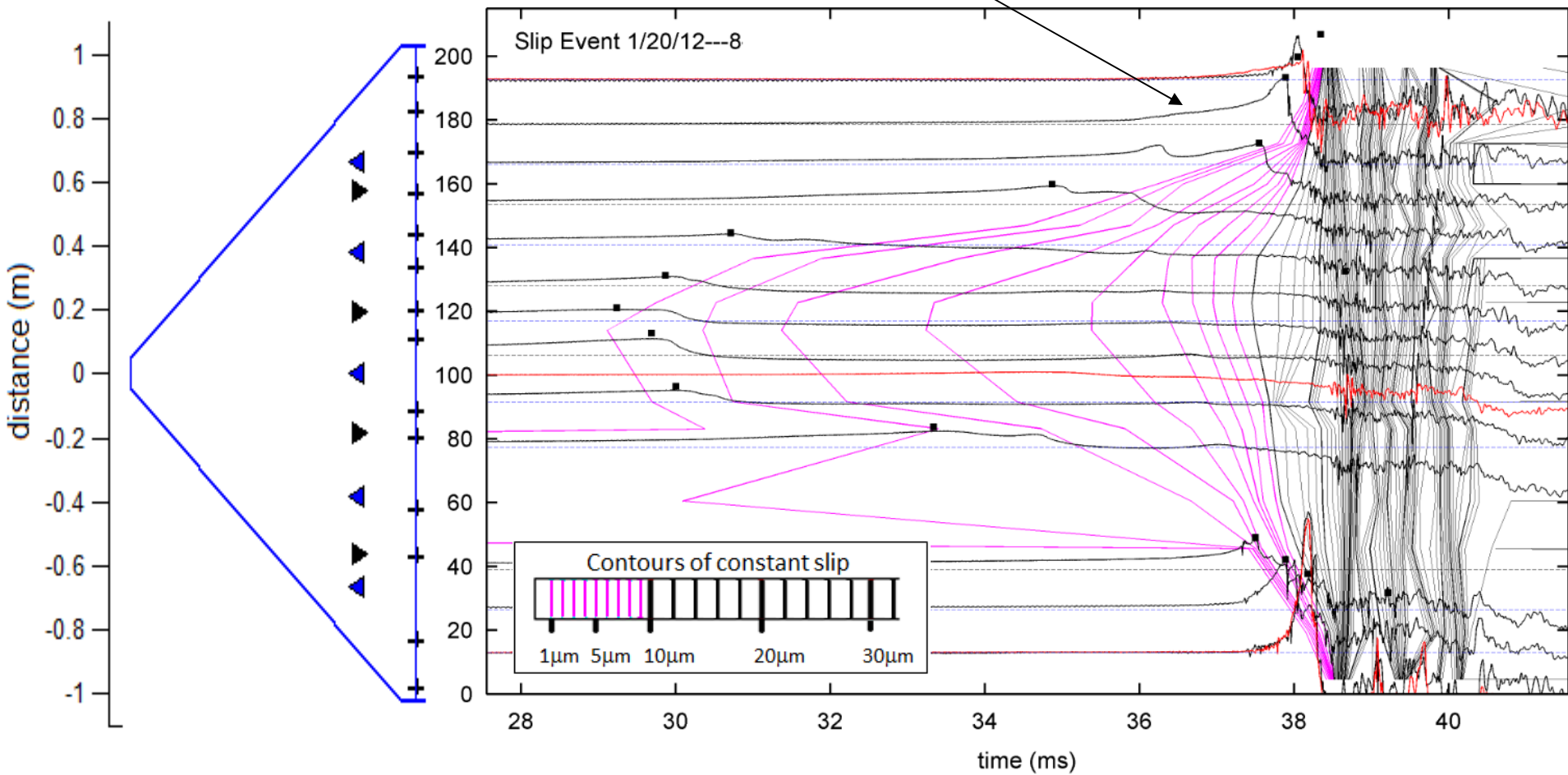


Mainshock



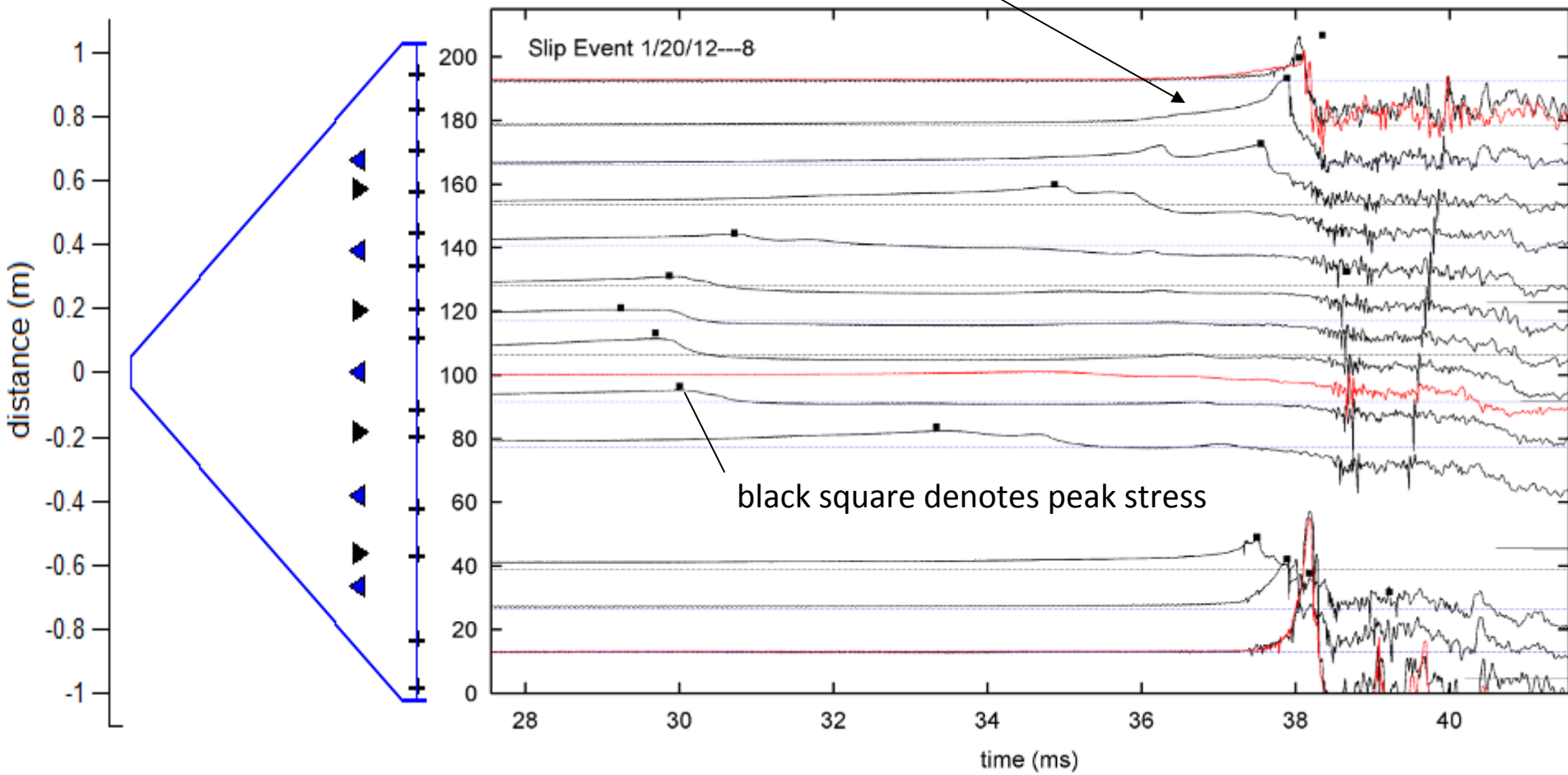
Slow slip during nucleation

Local shear stress measured from strain gage pairs along the fault

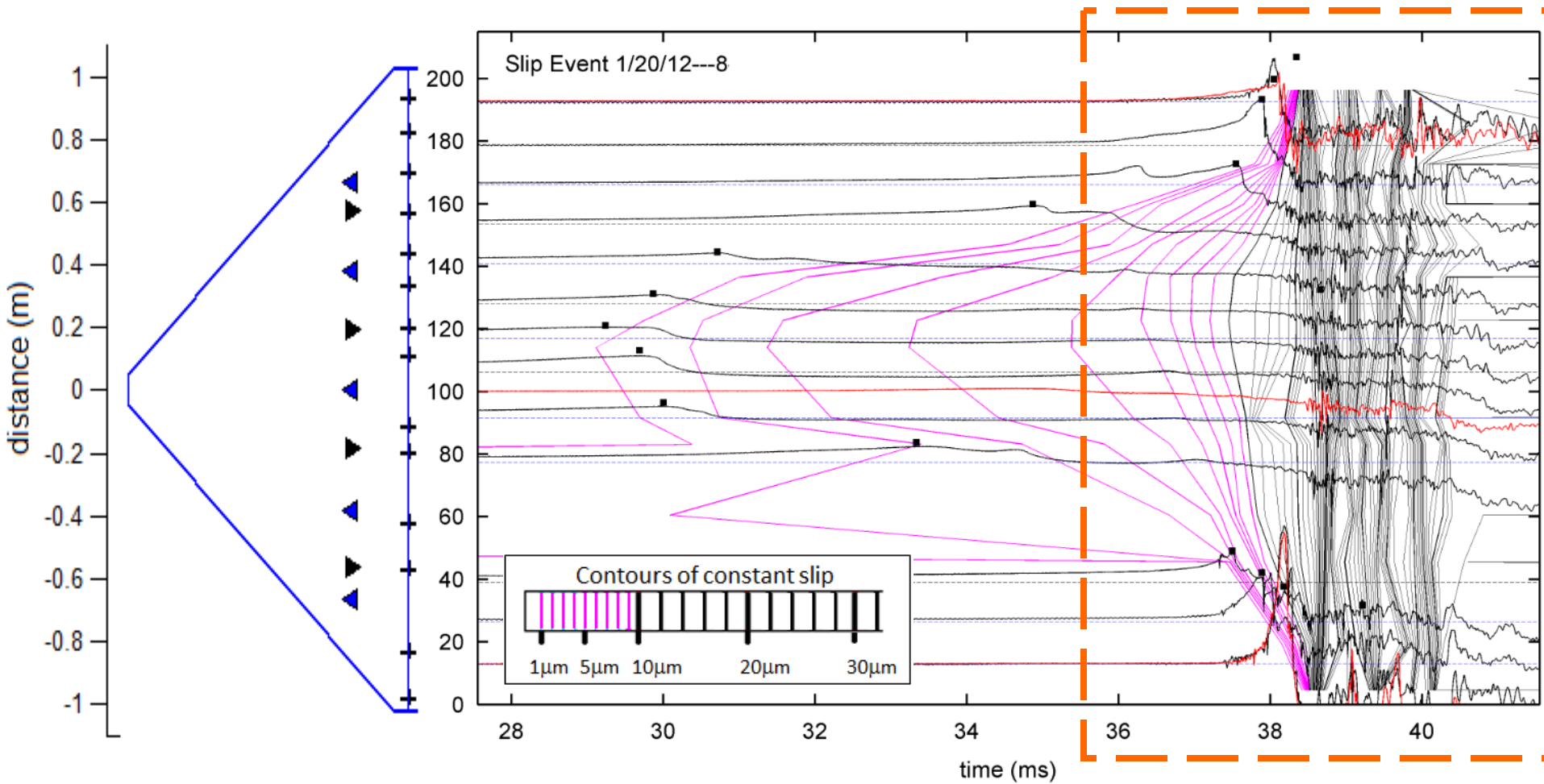


Slow slip during nucleation

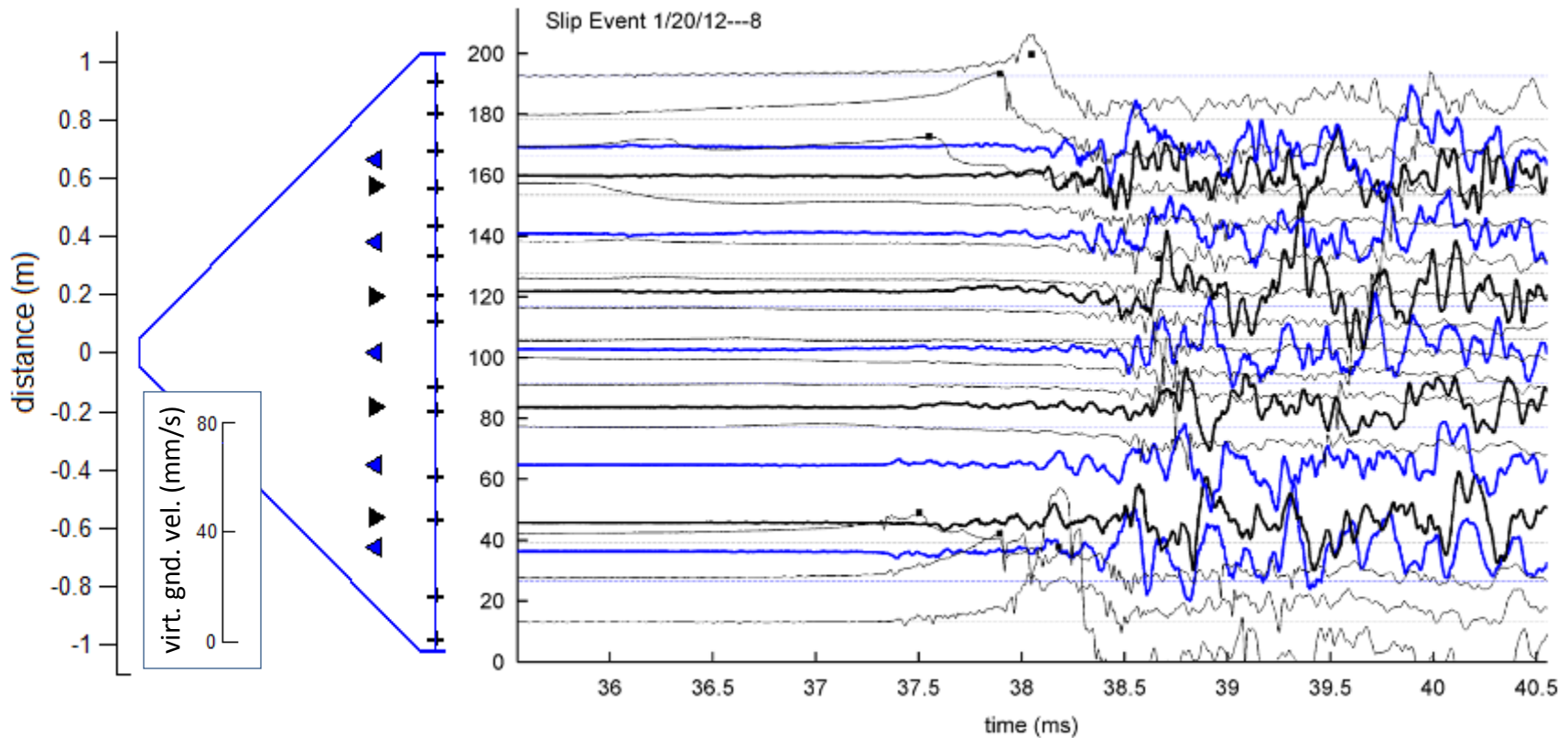
Local shear stress measured from strain gage pairs along the fault



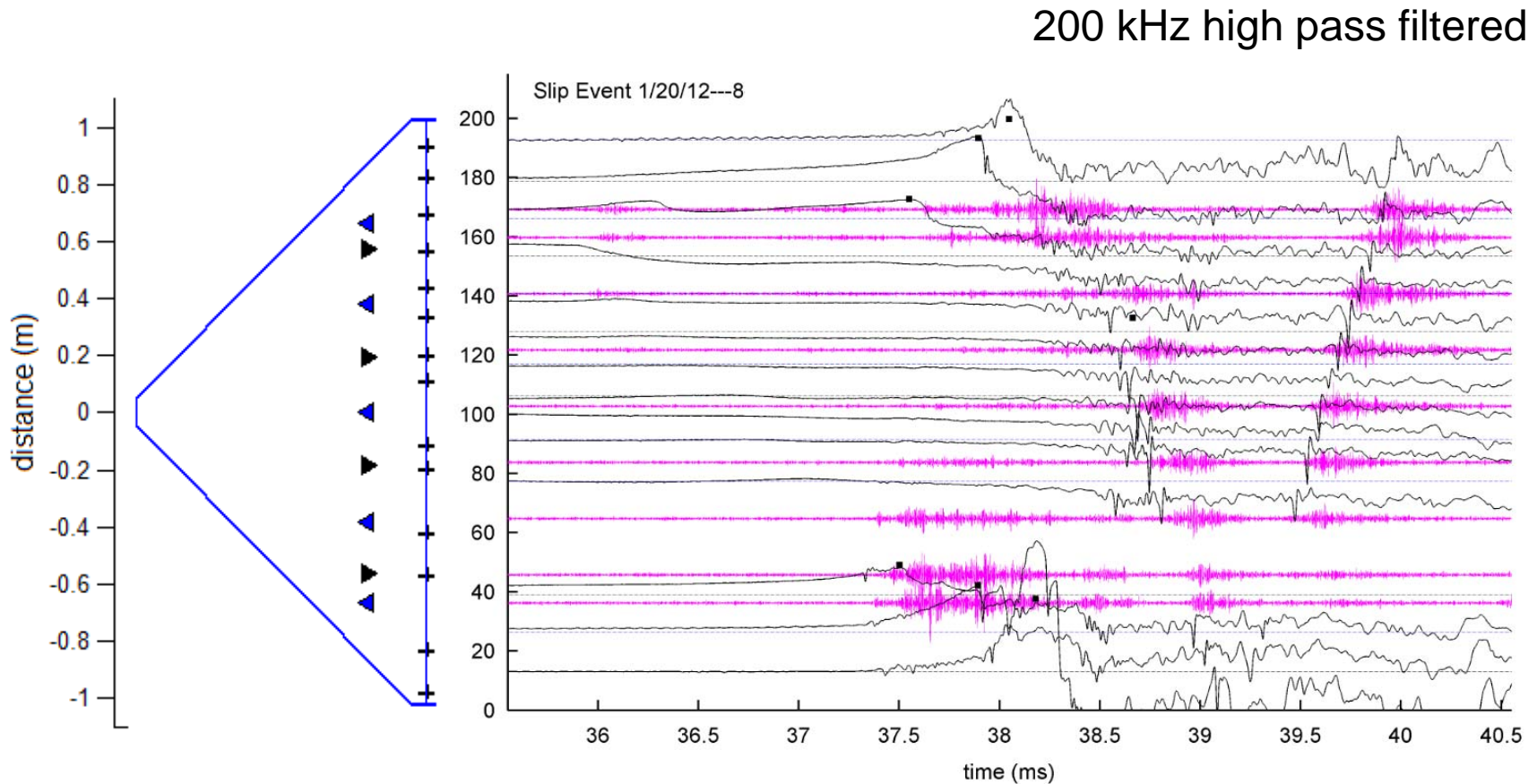
Zoom in on ground motions during “mainshock”



Complexity of HF seismicity during dynamic fault rupture



Complexity of HF seismicity during dynamic fault rupture

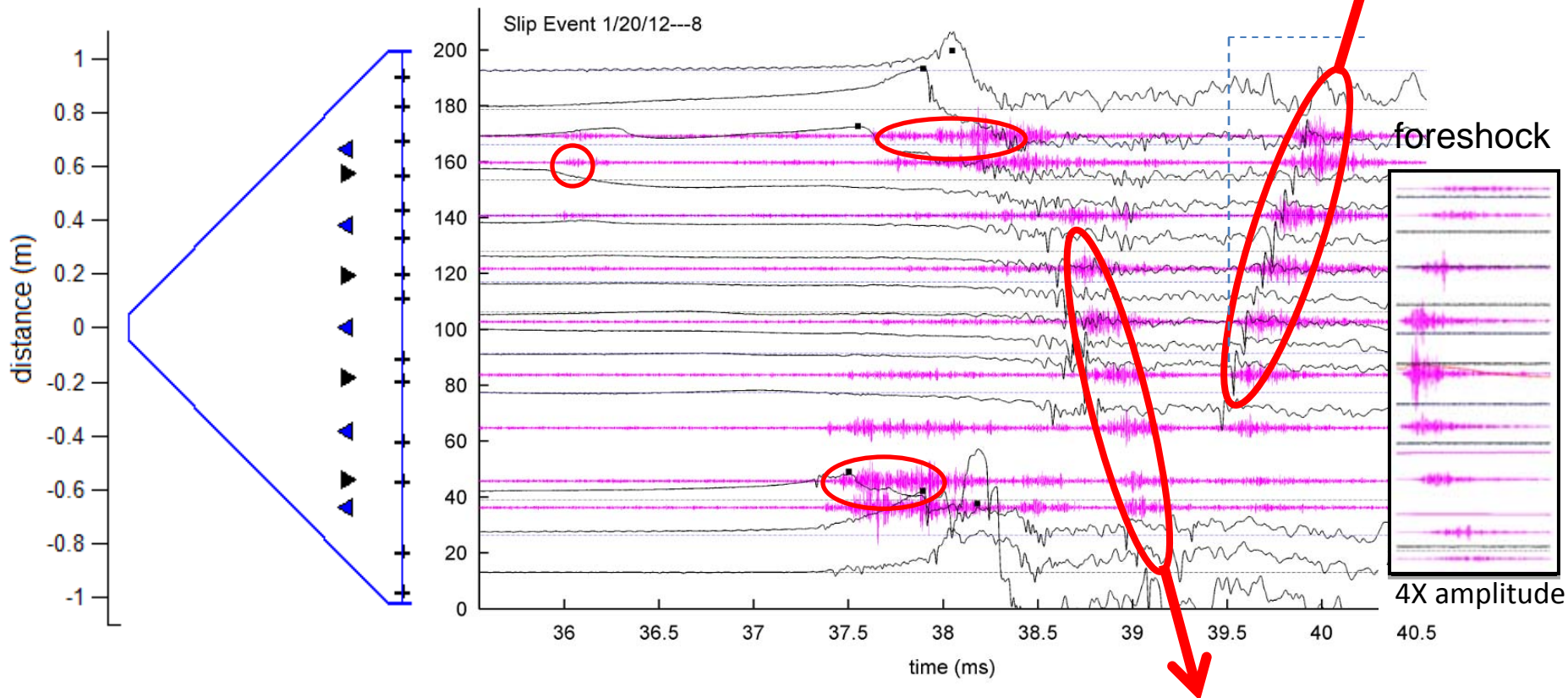


Moving sources of high frequency during dynamic fault rupture

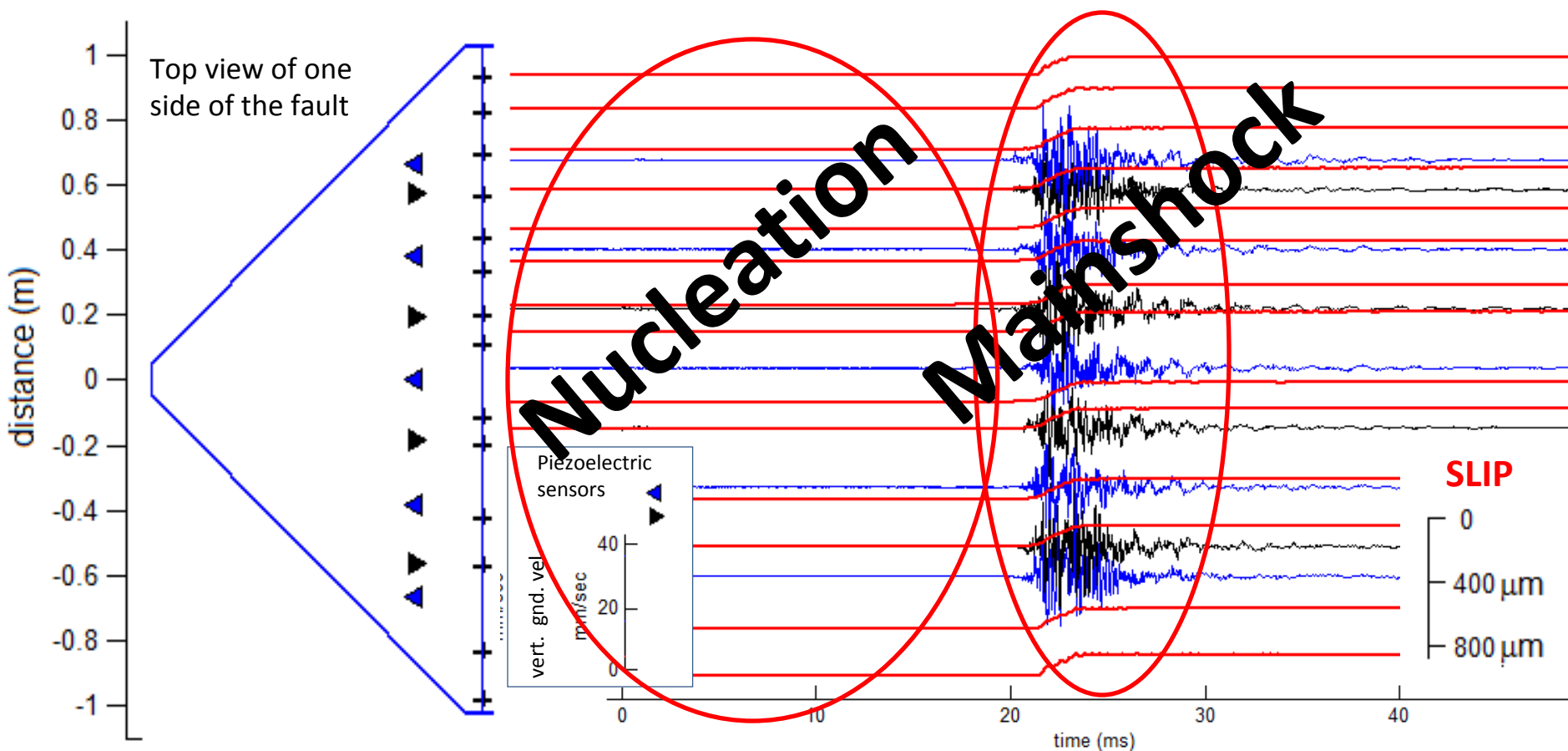
Magenta = 200 kHz high pass filtered ground motions recorded from piezoelectric sensor array.

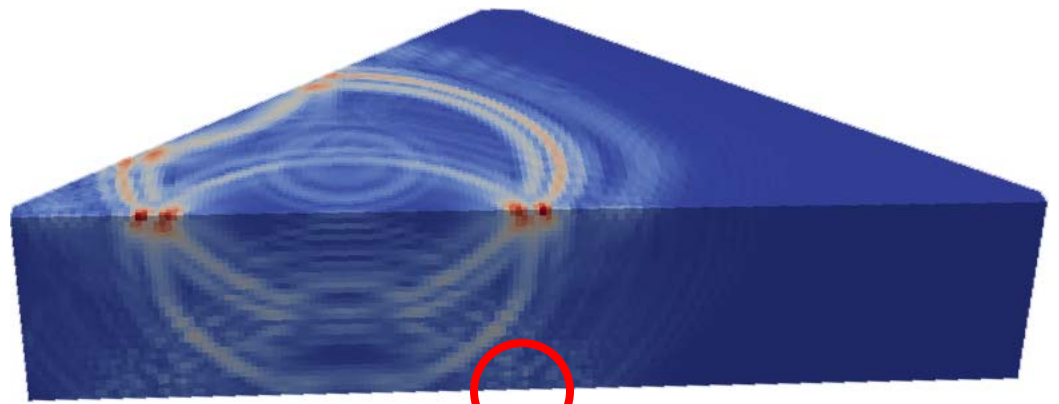
Propagation velocity
~2000 m/s

At these high frequencies attenuation is significant



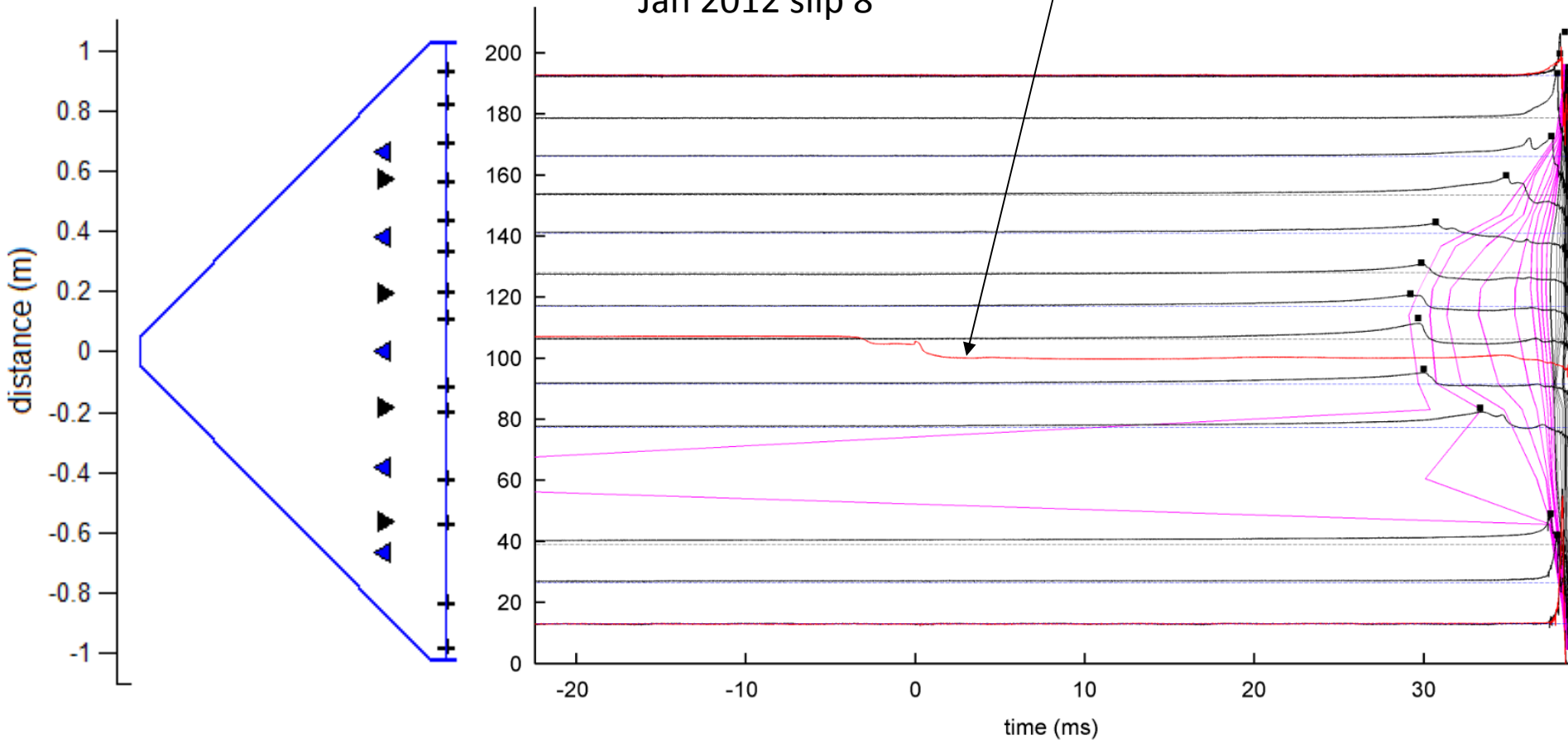
But this talk is about nucleation!





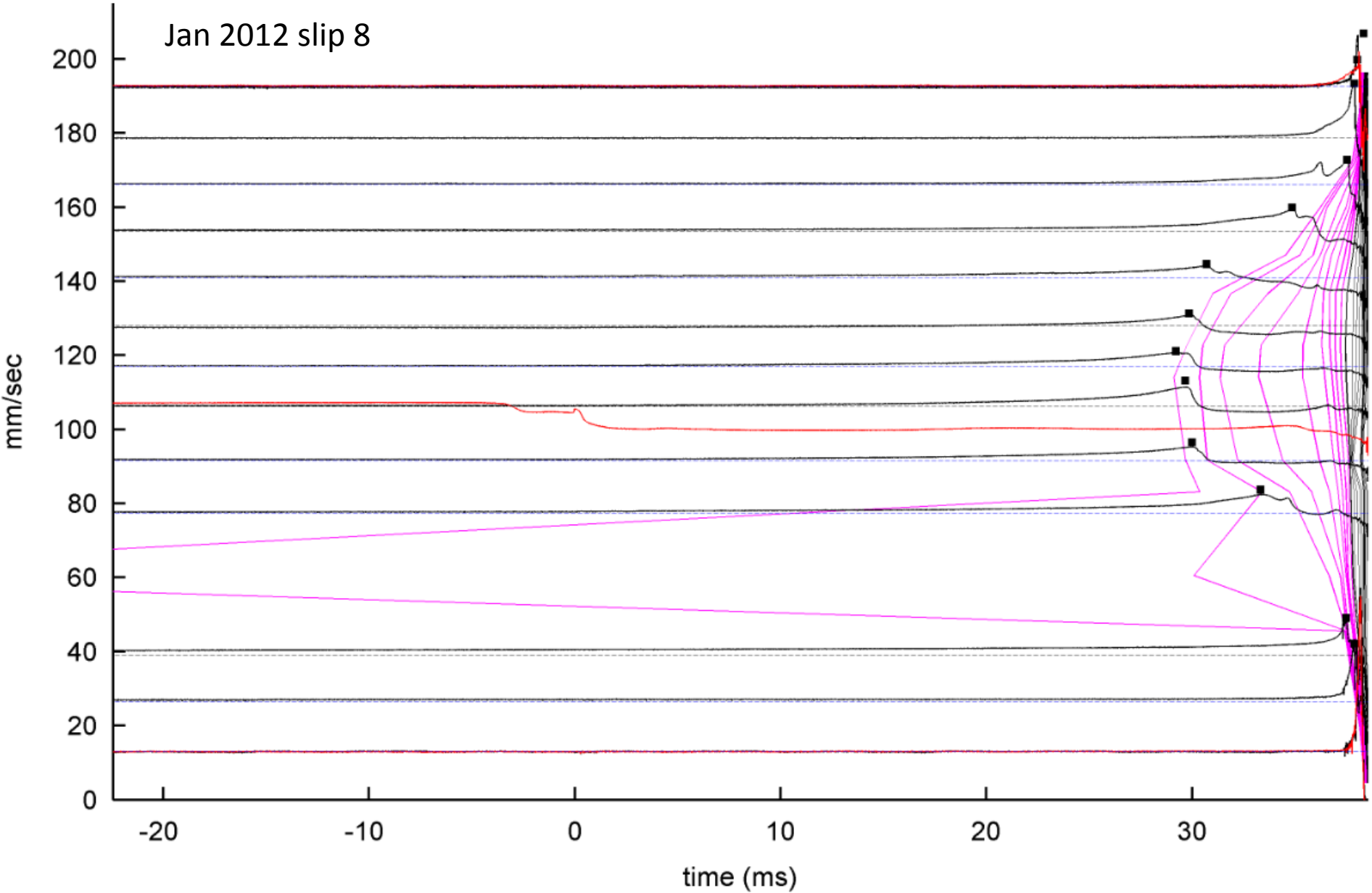
Local shear stress measured from strain gage pairs along the fault

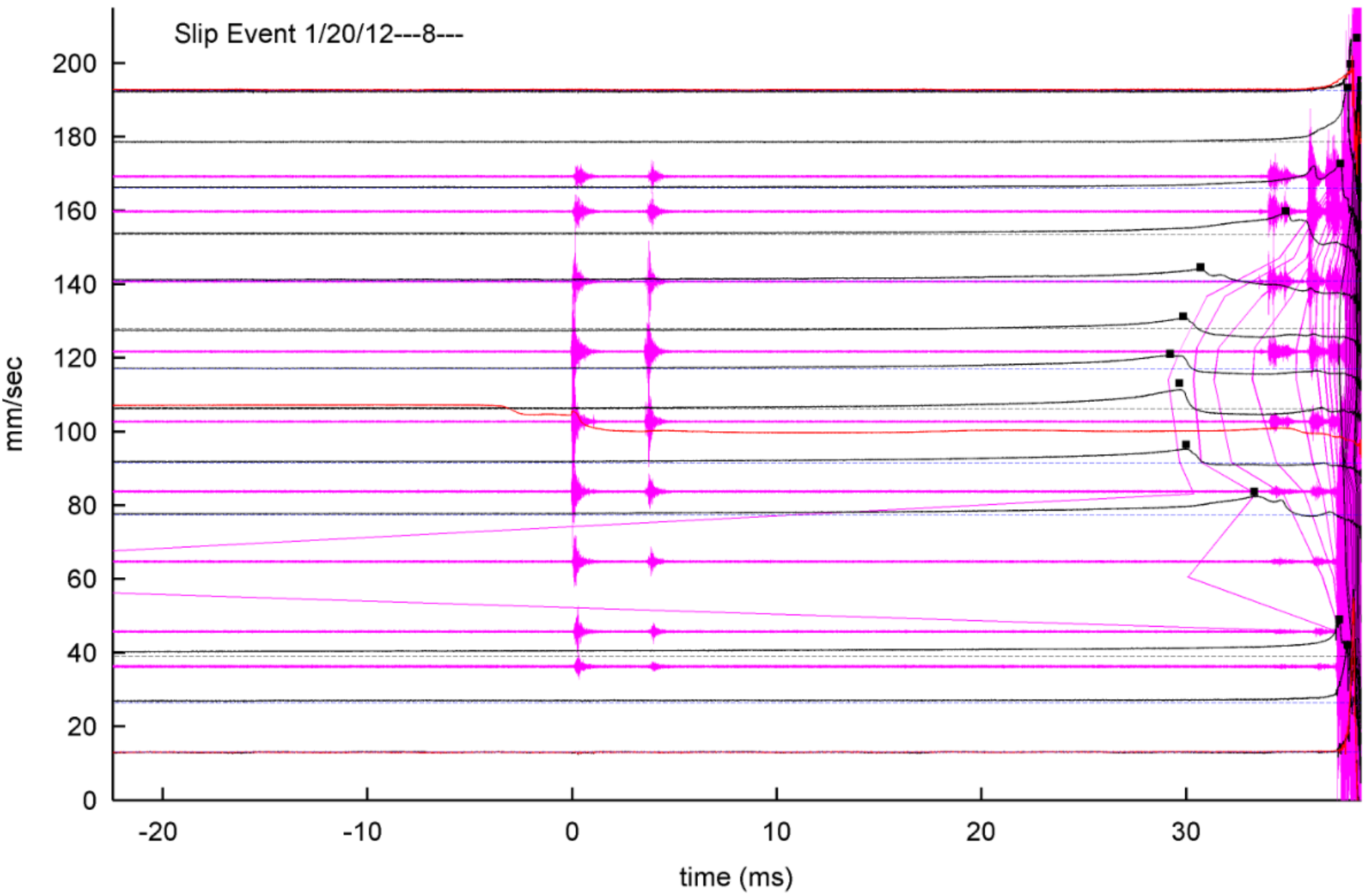
Jan 2012 slip 8



Local shear stress measured from strain gage pairs along the fault

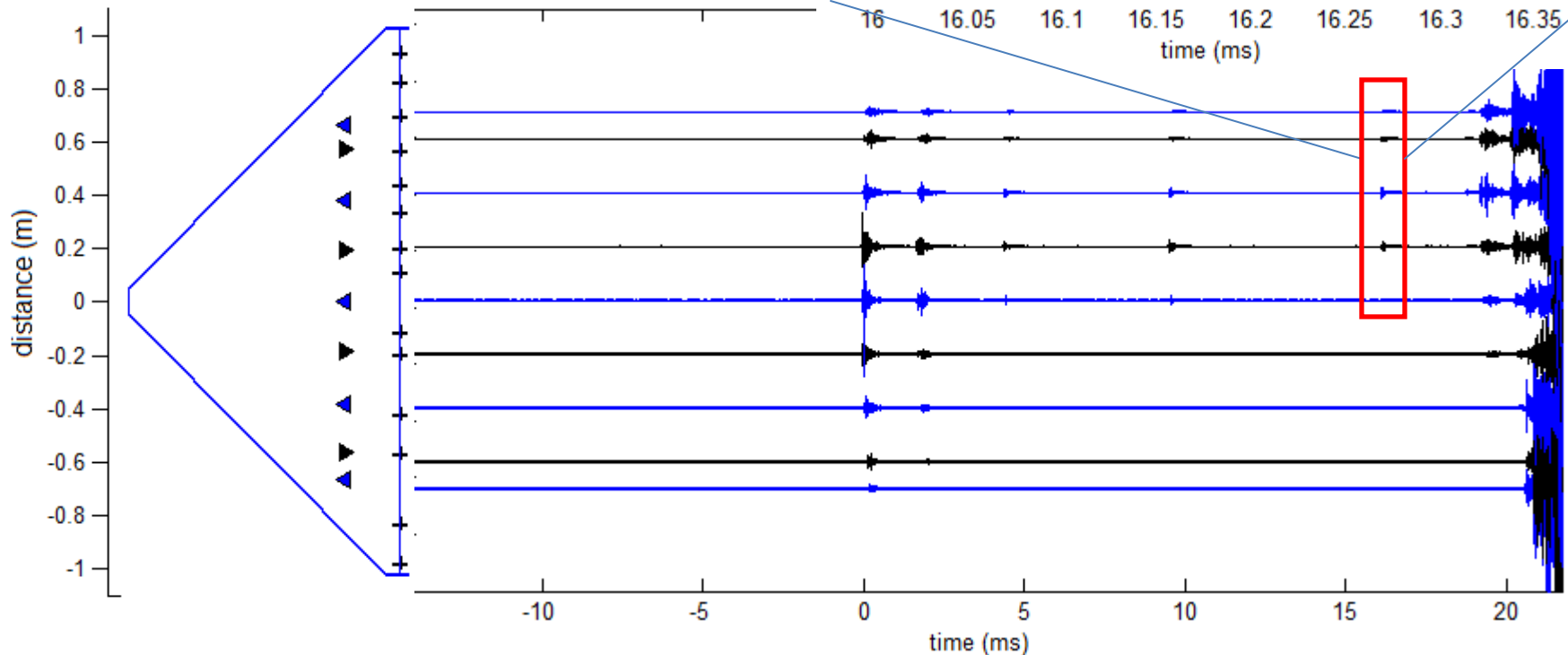
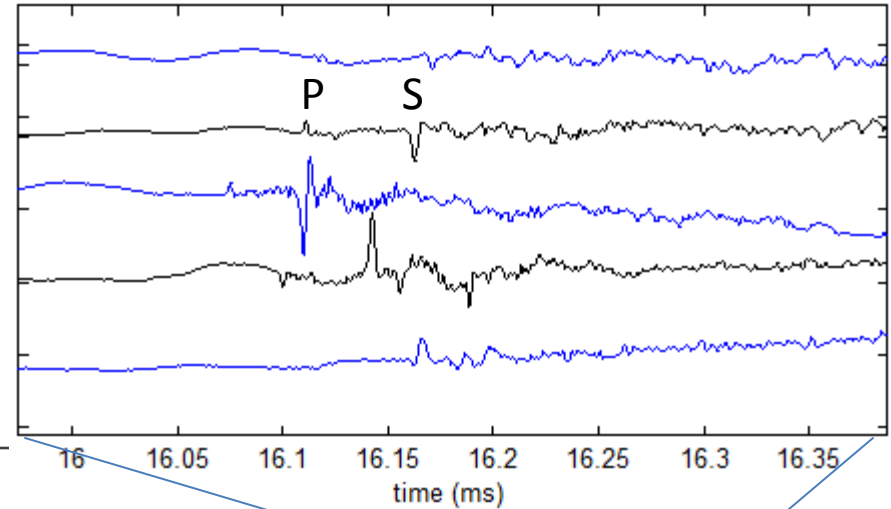
Jan 2012 slip 8





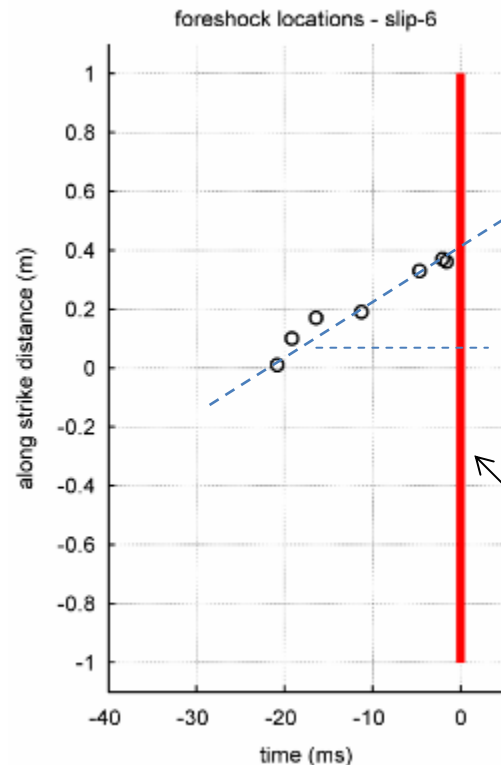
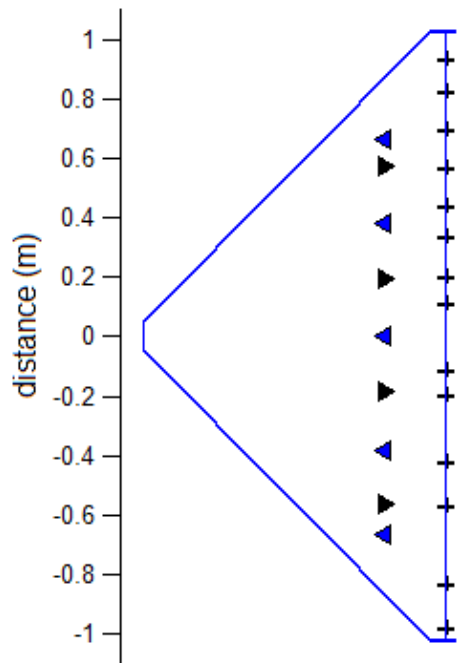
Nucleation: foreshocks

- Distinct, impulsive P and S waves
- Can locate to +/-10 mm accuracy
~3-10 μm source duration
- Max source radius 7-25 mm



Foreshock migration

- Impulsive foreshocks can be located to $\sim\pm 10$ mm accuracy
- Foreshock migration velocity along strike ~ 20 m/s
 - 2 orders of magnitude slower than migration of high frequency sources during dynamic rupture.

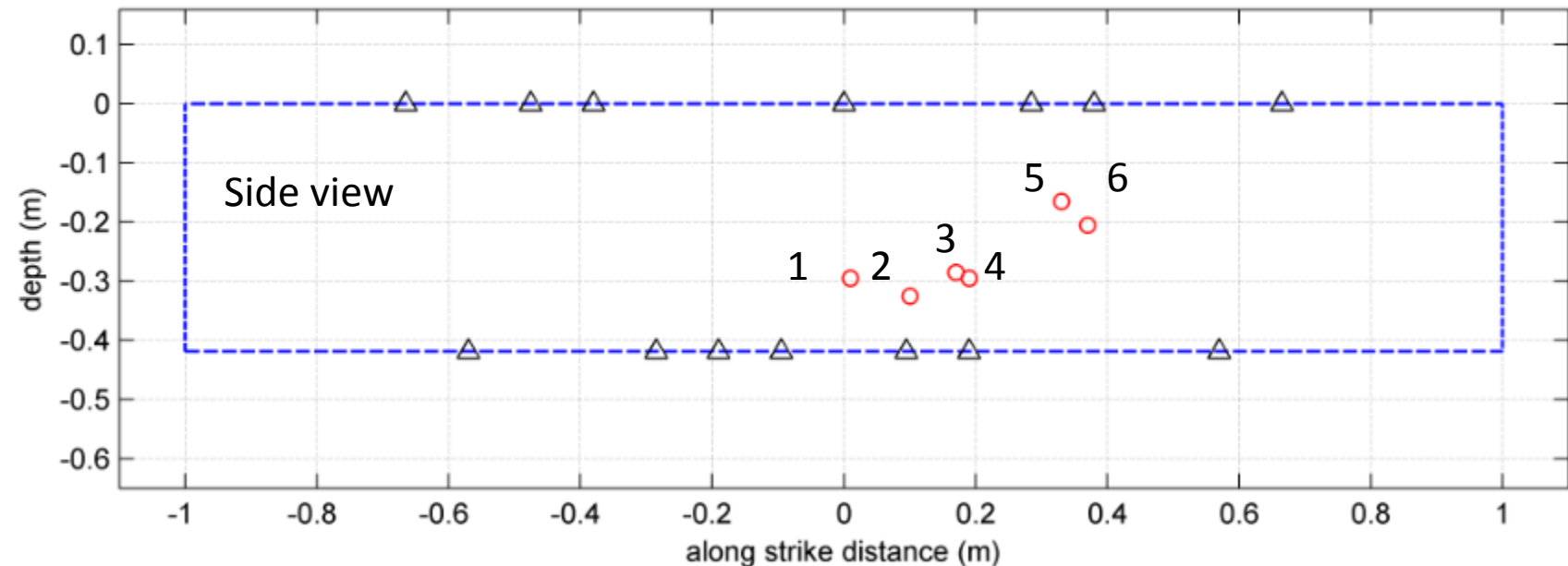
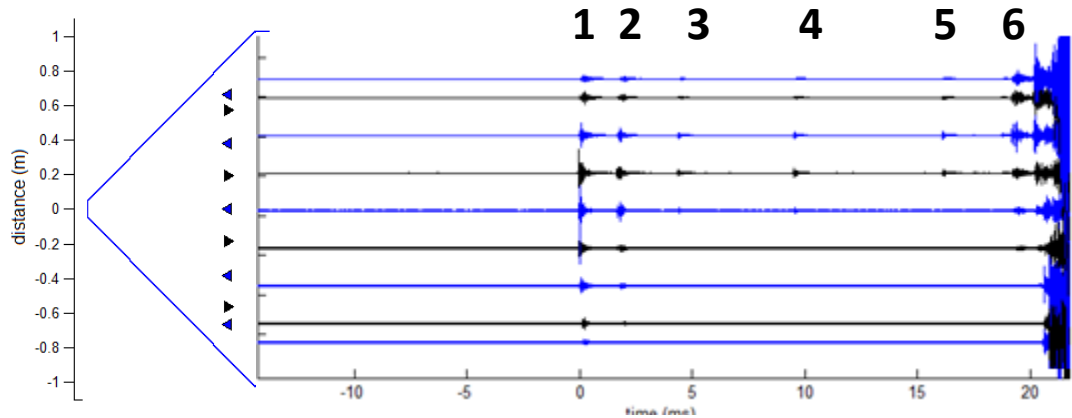


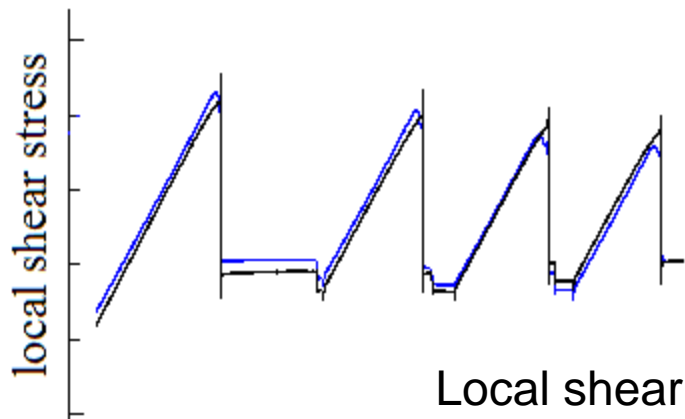
~ 20 m/s

time of onset of rapid sliding
(the mainshock).

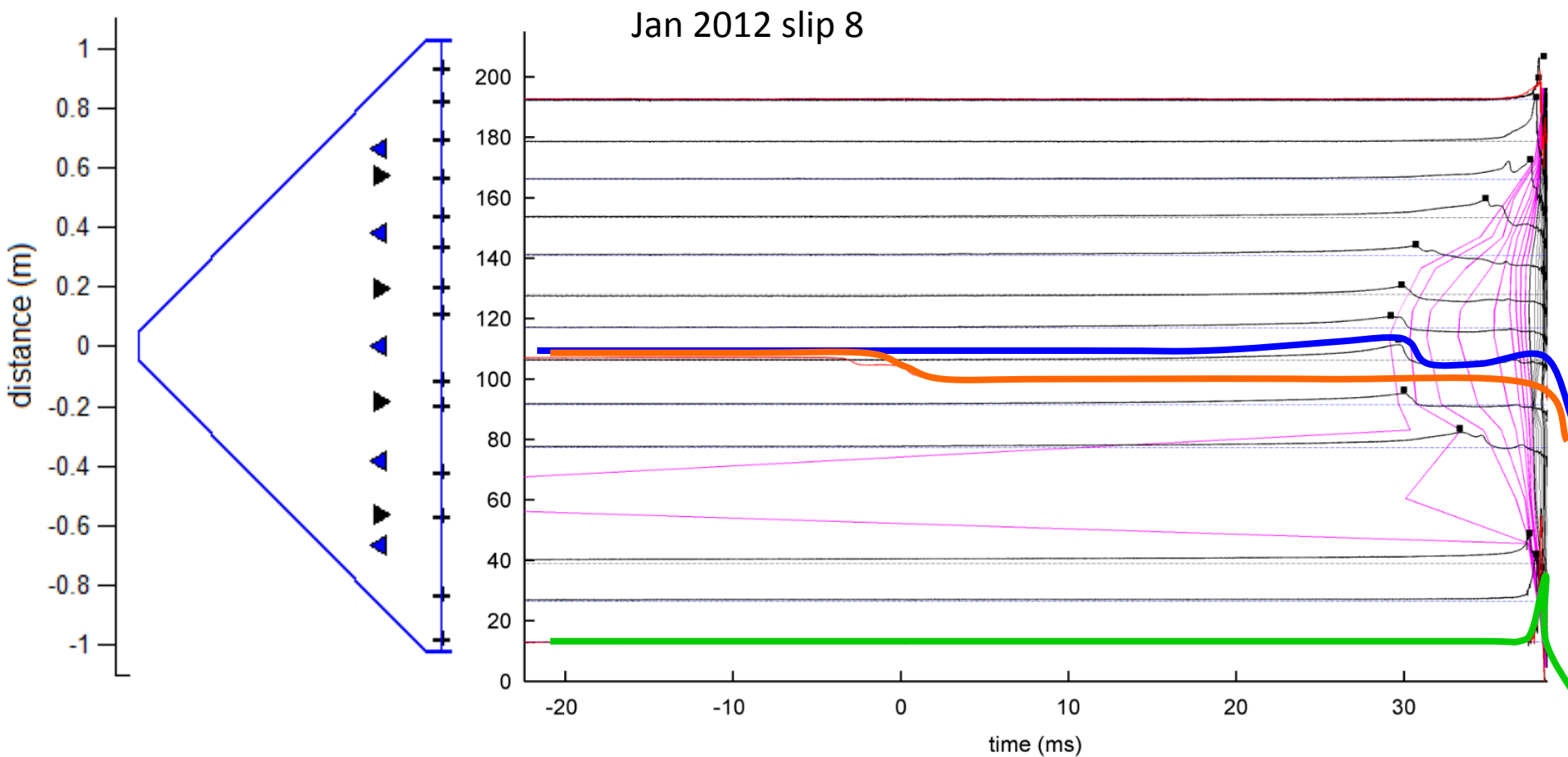
Foreshock locations: along strike and depth

- Foreshock locations from one stick slip event (Slip #6)

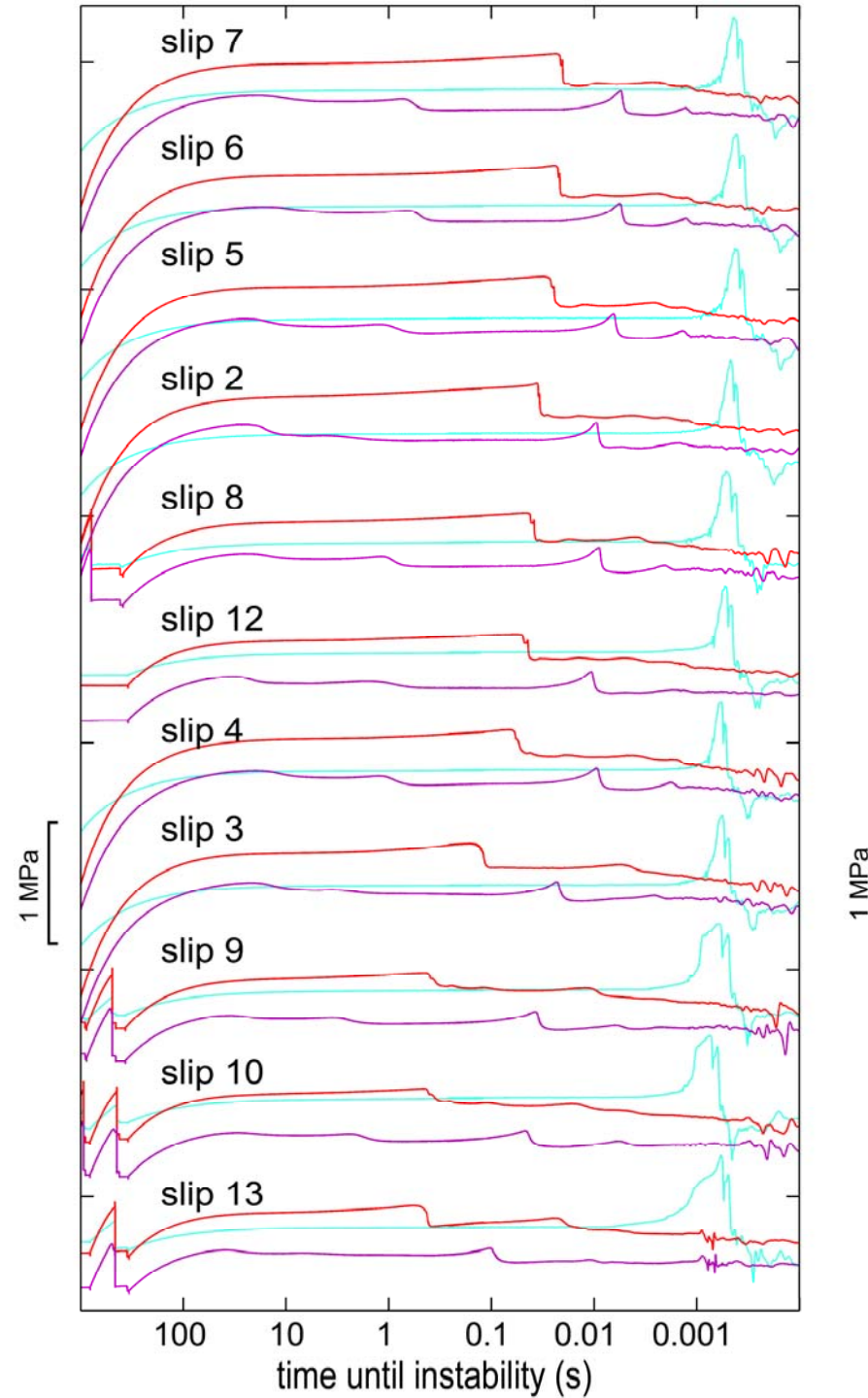
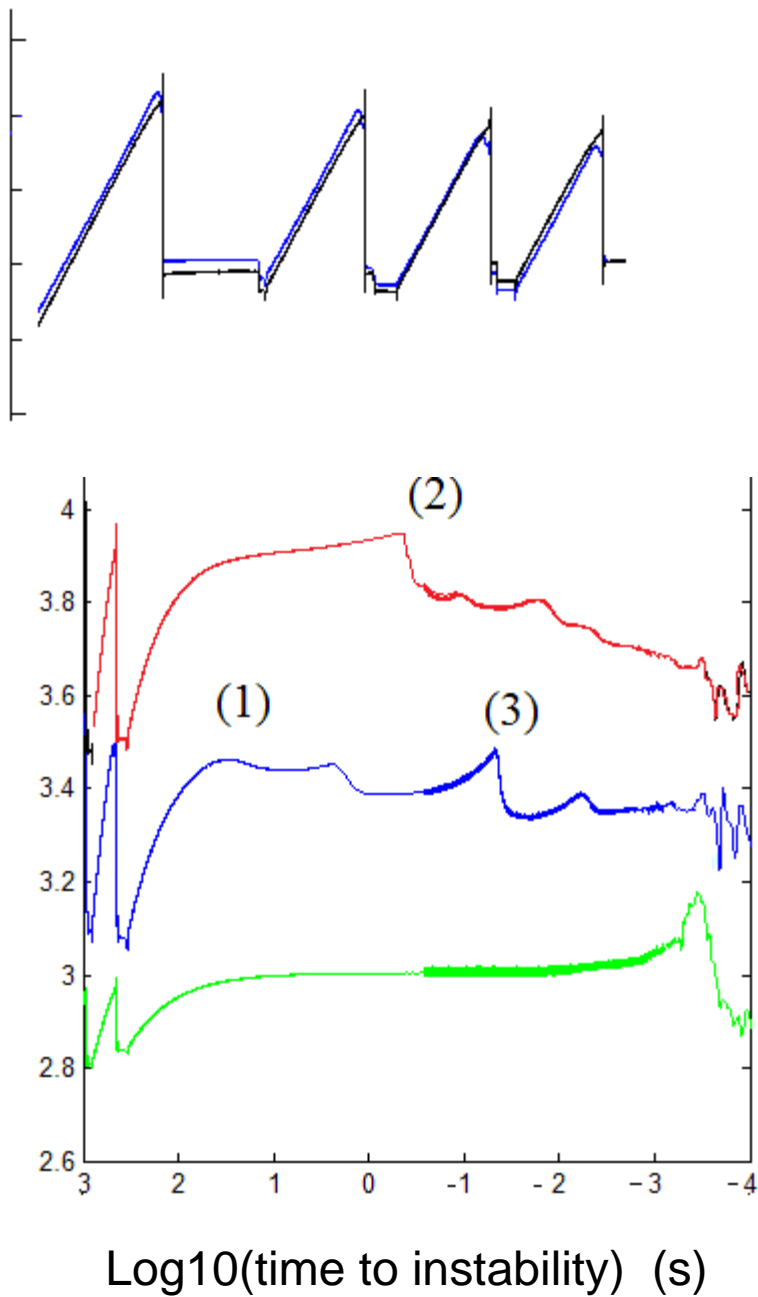


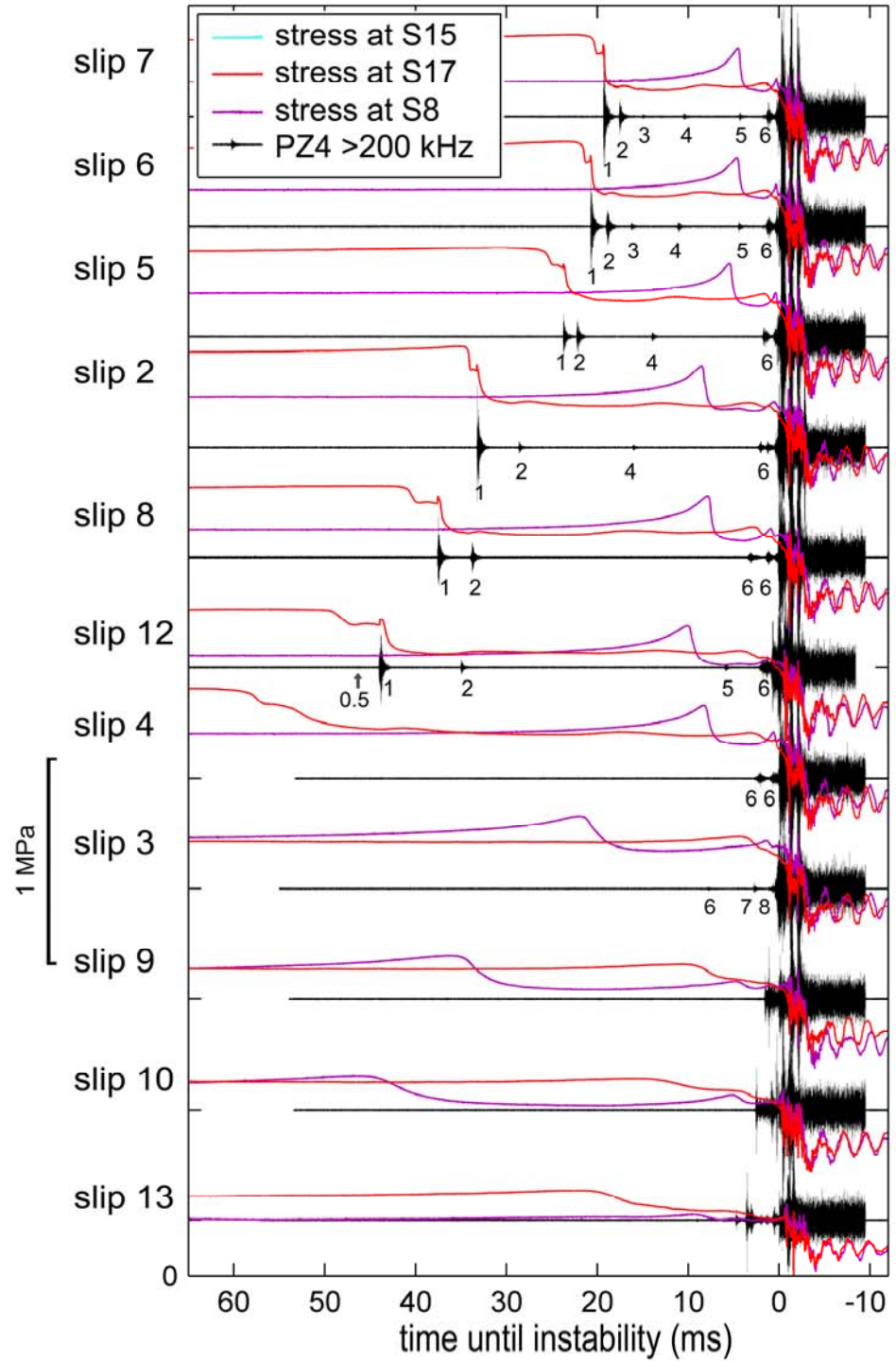
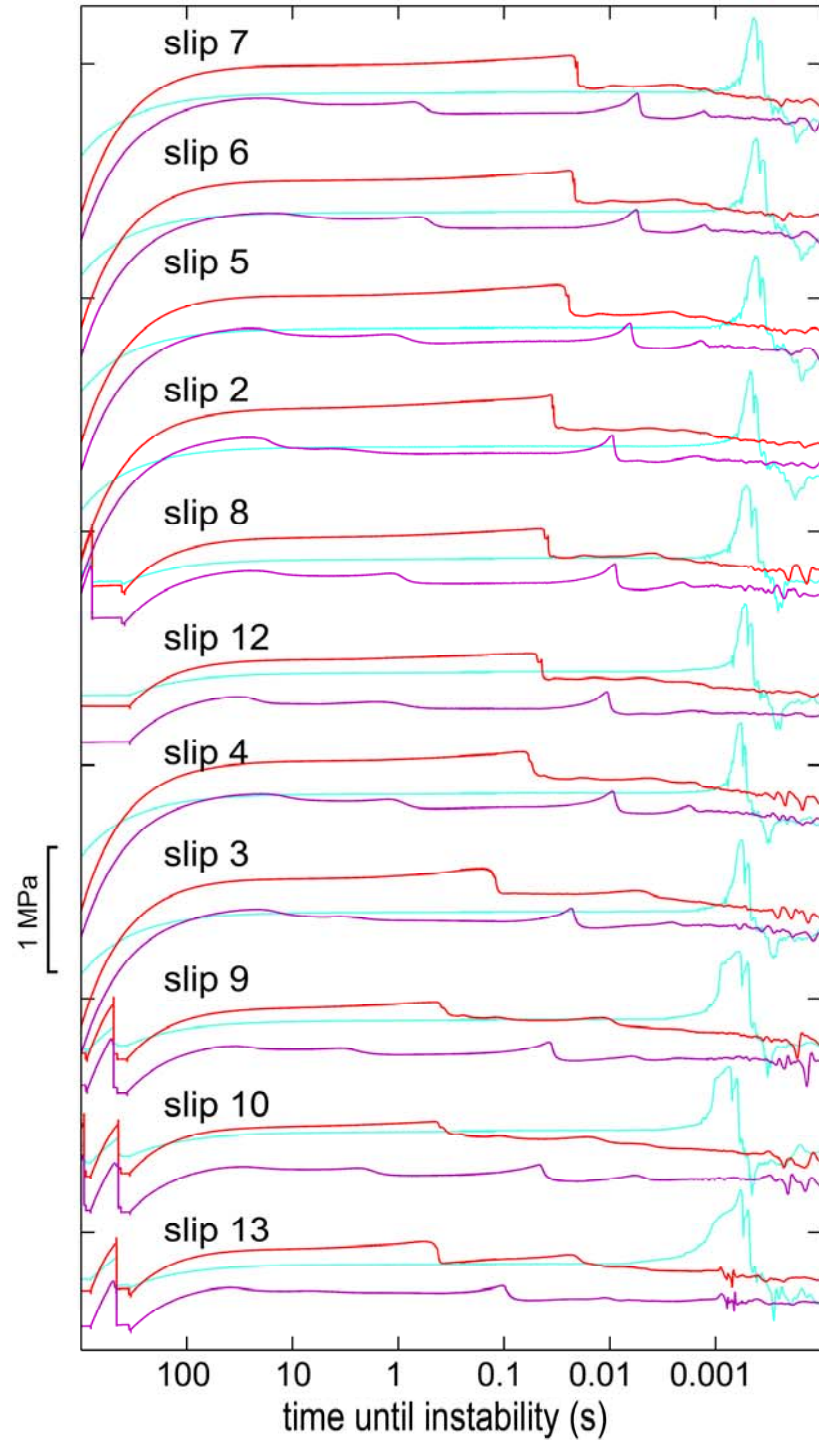


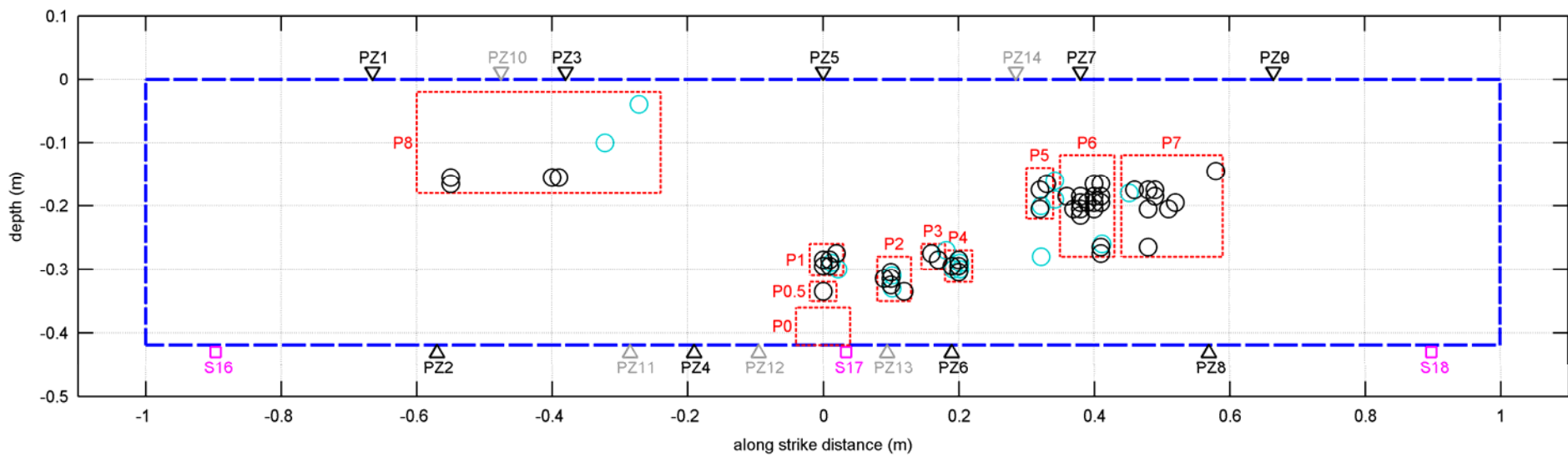
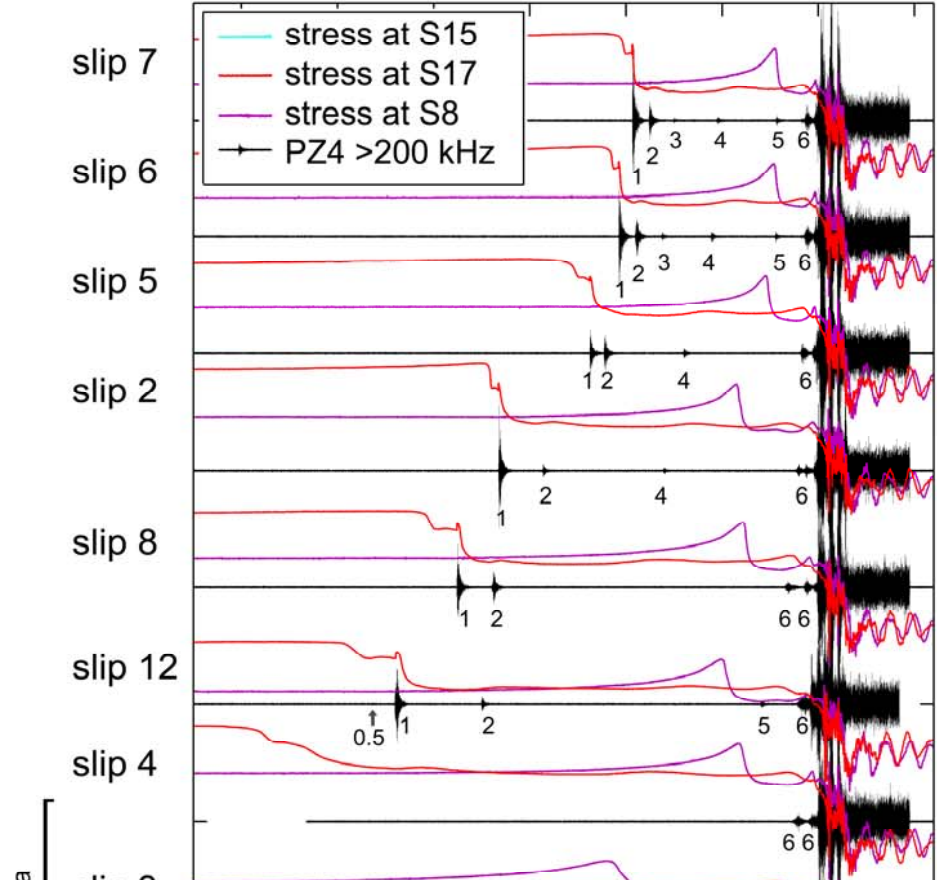
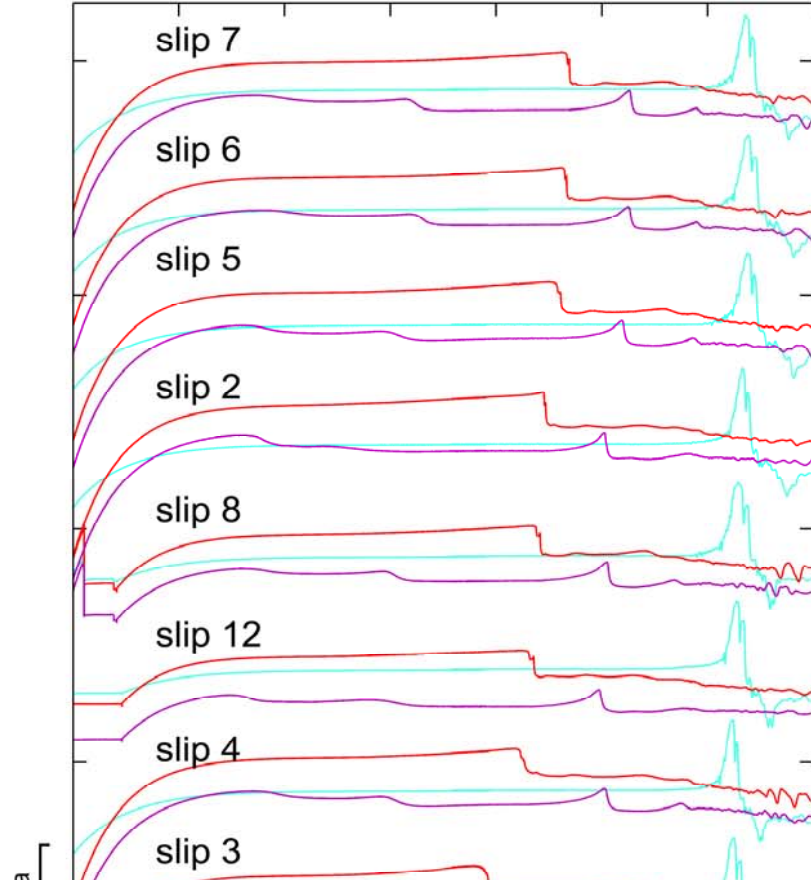
Local shear stress measured from strain gage pairs along the fault

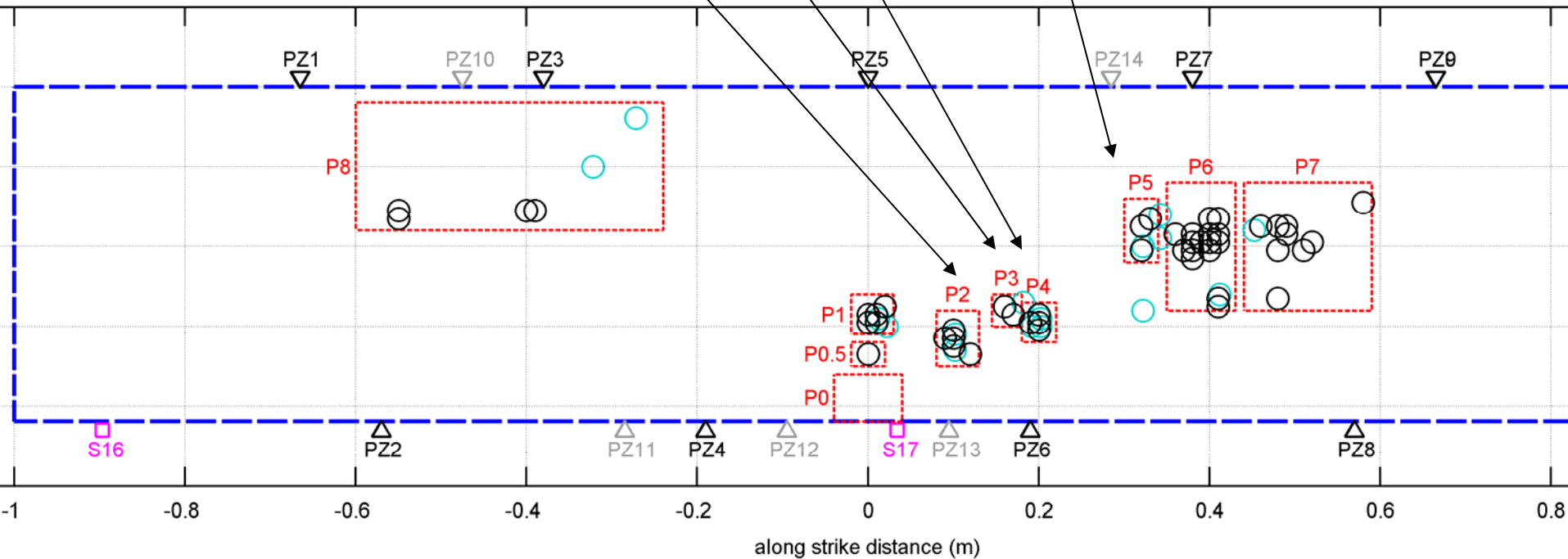
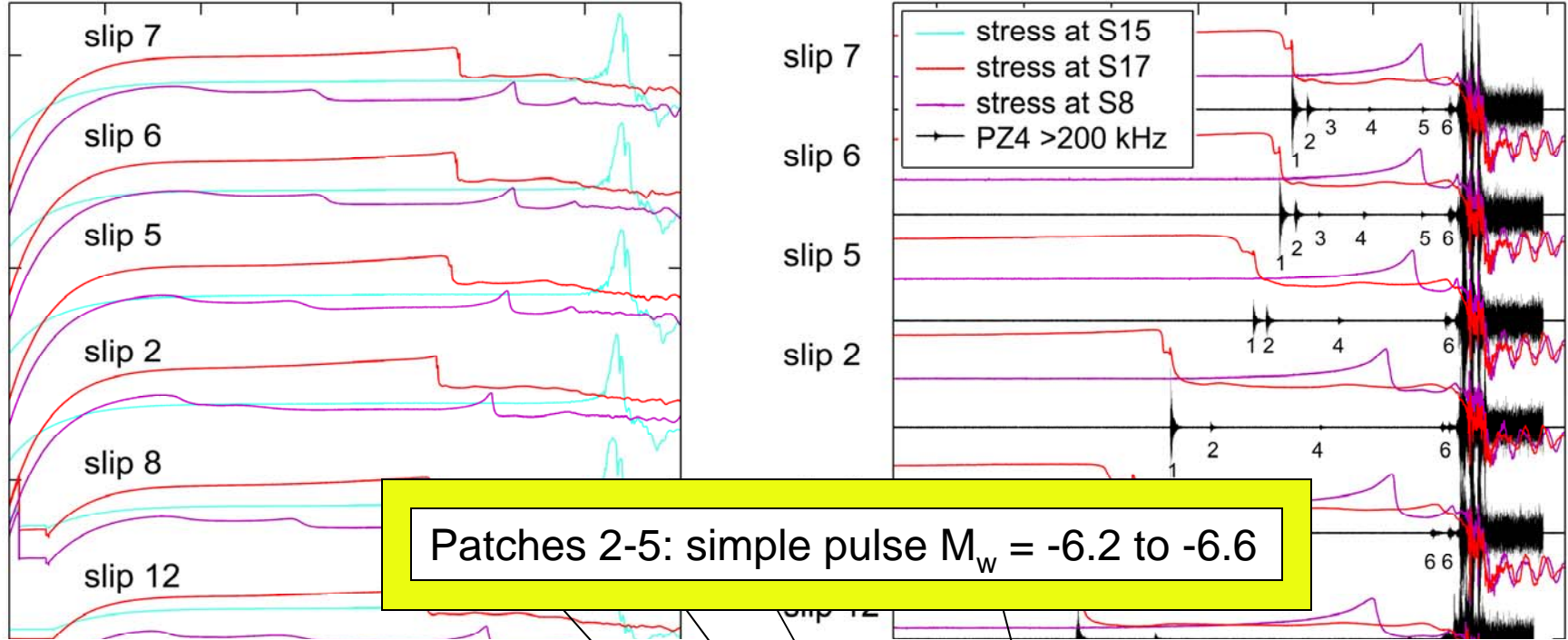


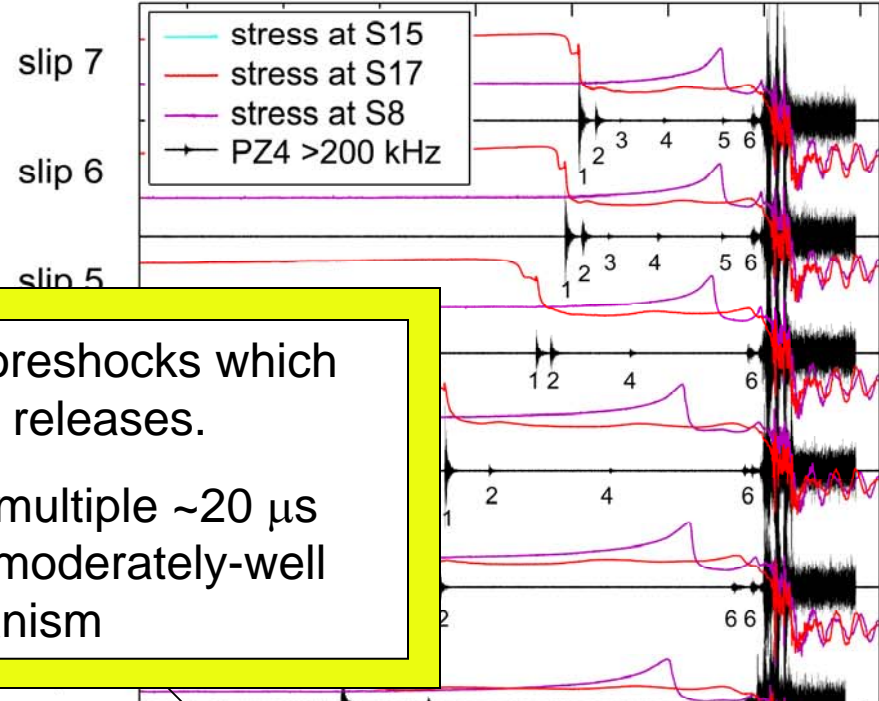
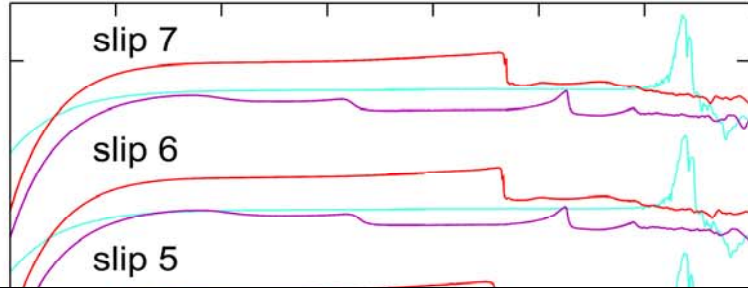
local shear stress



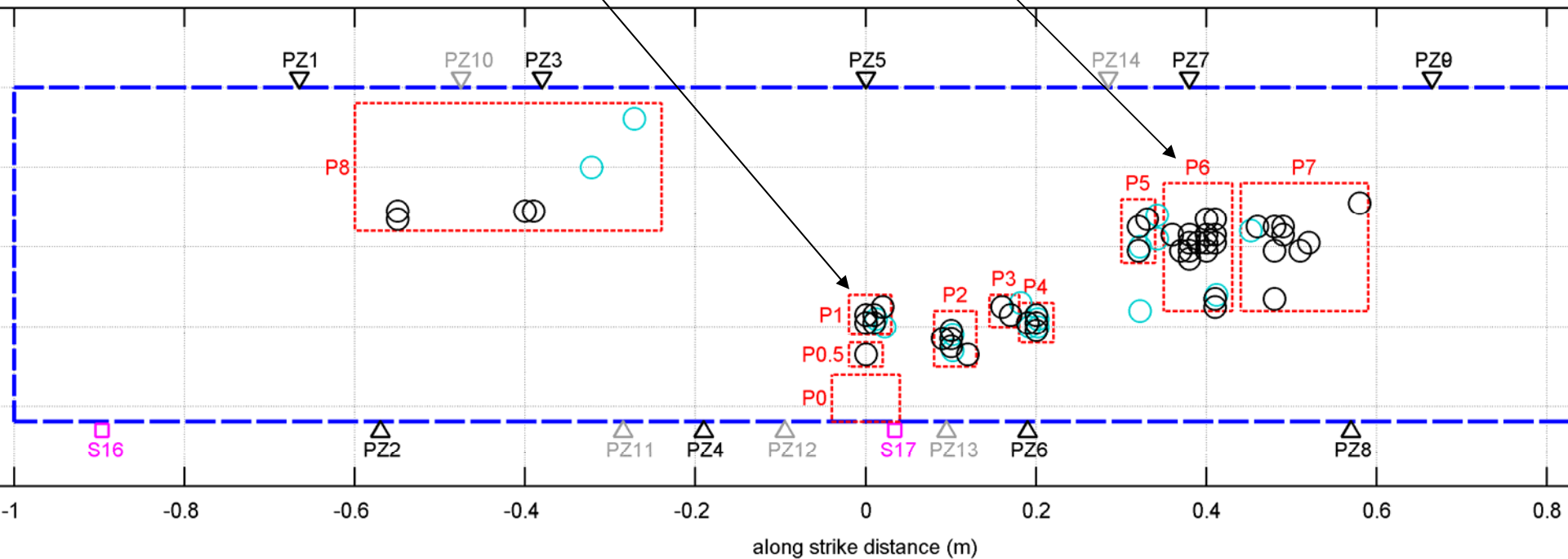


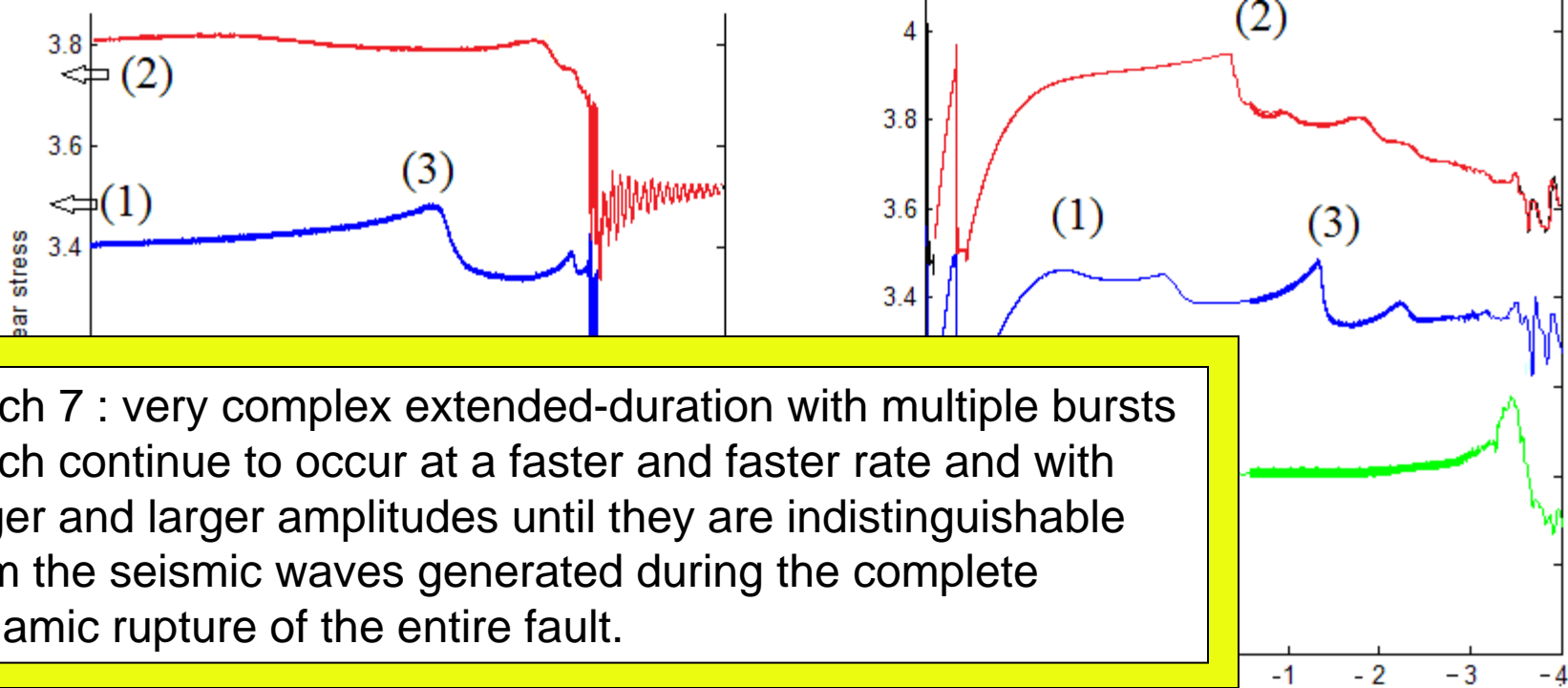




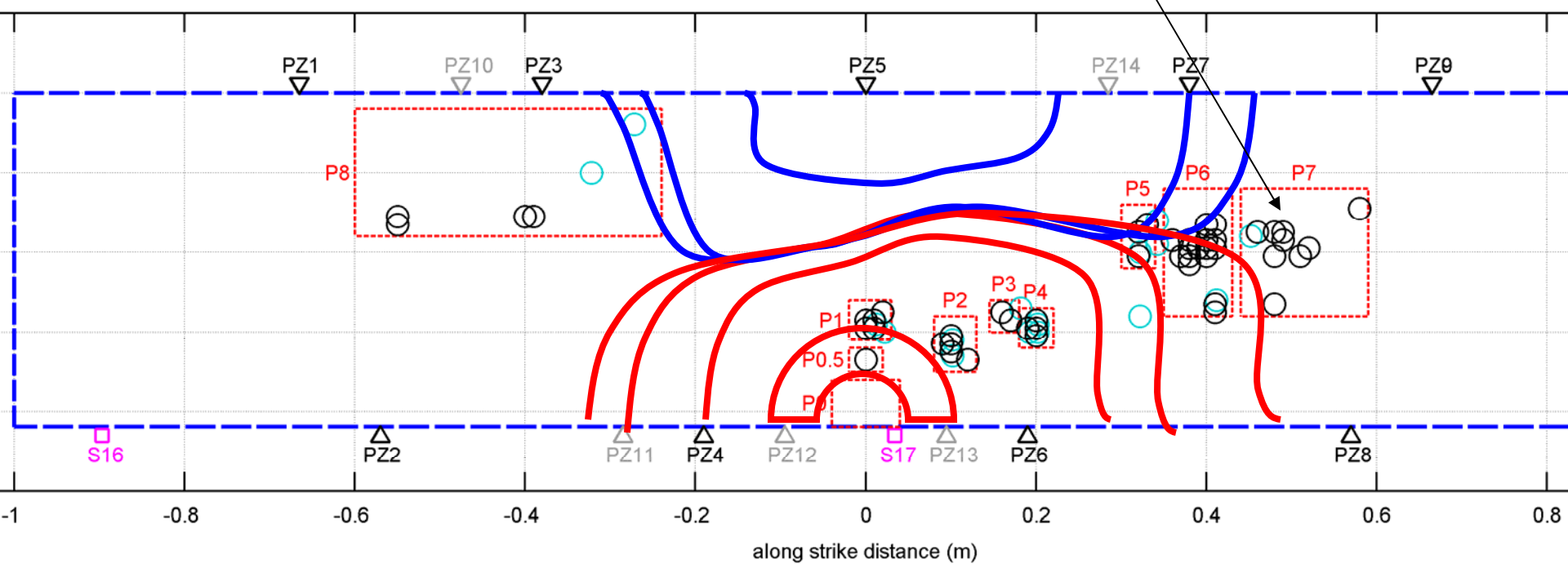


Patch1 and 6: complex extended-duration foreshocks which consist of multiple $M_w = -5.9$ to -4.8 moment releases.
 ~100 μ s in total and are often composed of multiple ~20 μ s bursts of moment release, each of which is moderately-well modeled by a shear dislocation focal mechanism



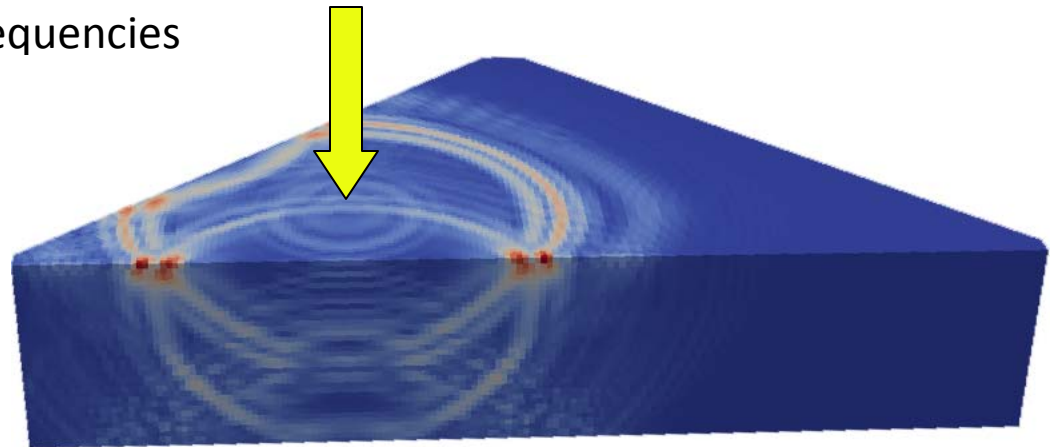


Patch 7 : very complex extended-duration with multiple bursts which continue to occur at a faster and faster rate and with larger and larger amplitudes until they are indistinguishable from the seismic waves generated during the complete dynamic rupture of the entire fault.



Modeling Laboratory Earthquakes

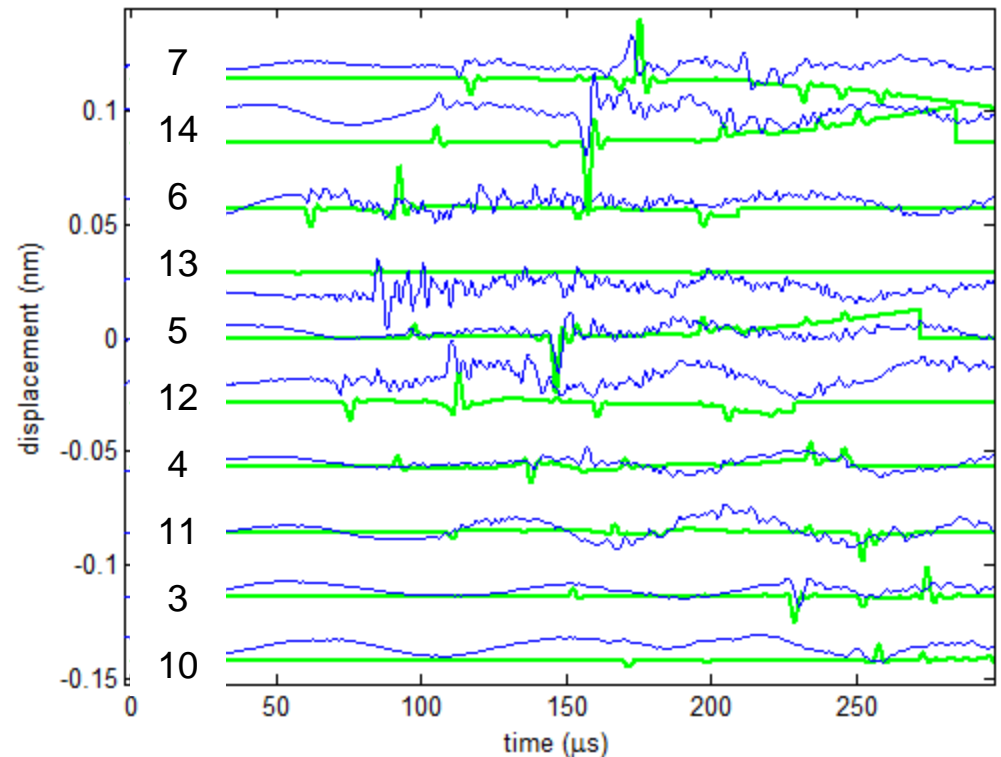
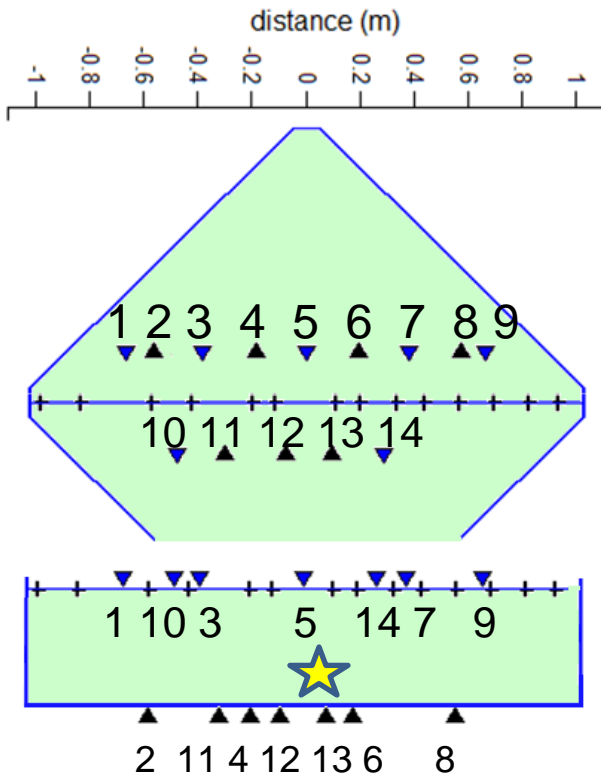
- Wave propagation modeling- Sierra White Granite
 - Goal is to develop a Green's function
 - Generalized ray theory, infinite plate
 - Good for simple geometry and few rays
 - Finite element models
 - Good for complicated geometry
 - Difficult to get high frequencies



Waveform modeling of impulsive, high-frequency foreshocks

- Foreshock modeled as a $\sim 3 \mu\text{s}$ pulse in moment rate
- Patch size ($\sim 7 \text{ mm}$ upper bound)

Patch 2 foreshock

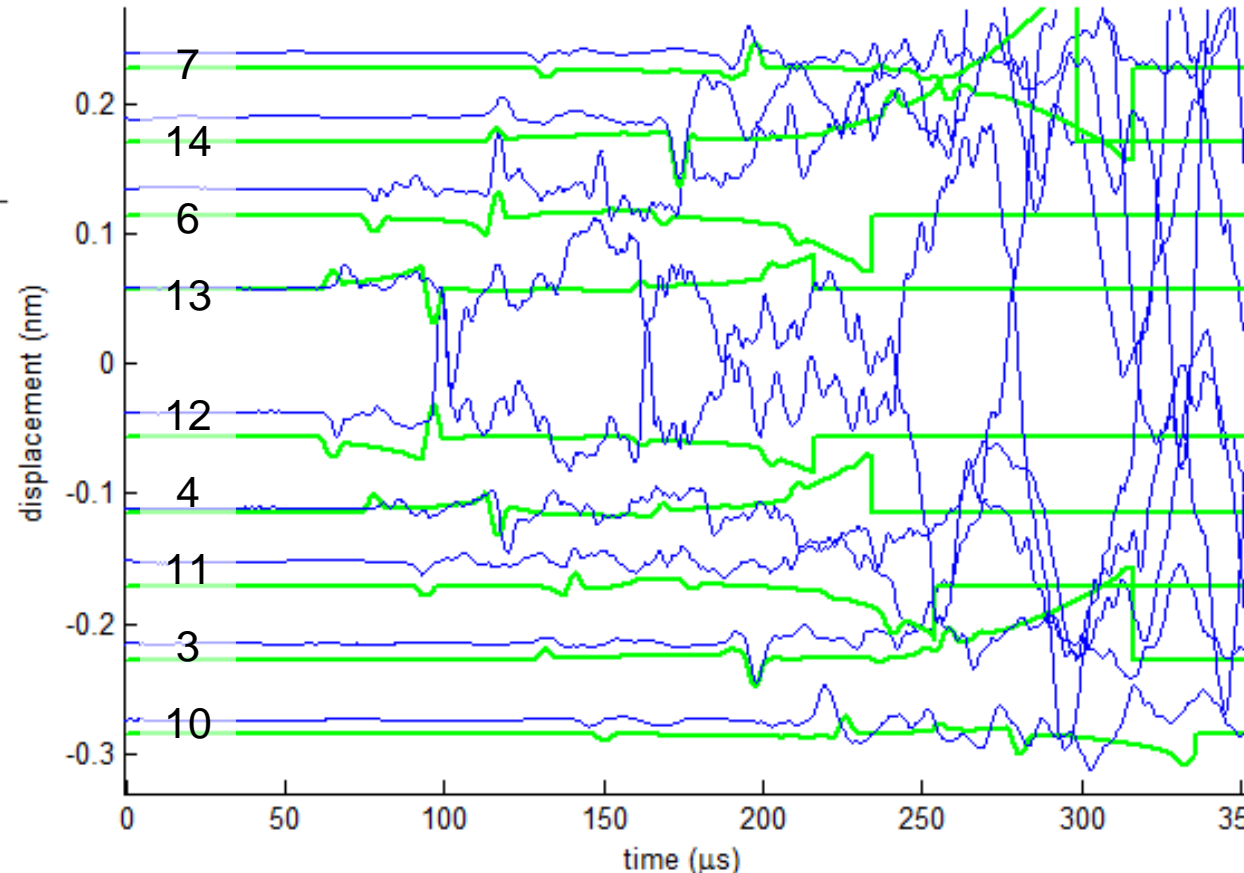
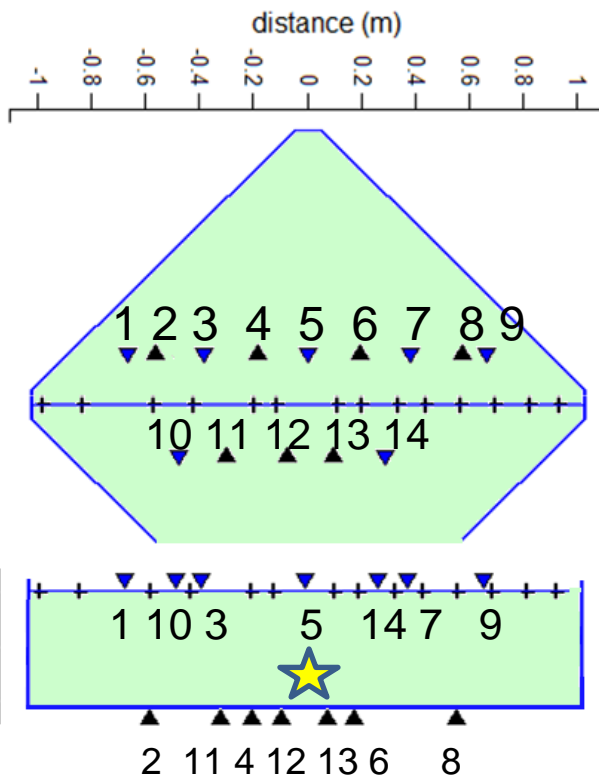


Waveform modeling of impulsive, high-frequency foreshocks

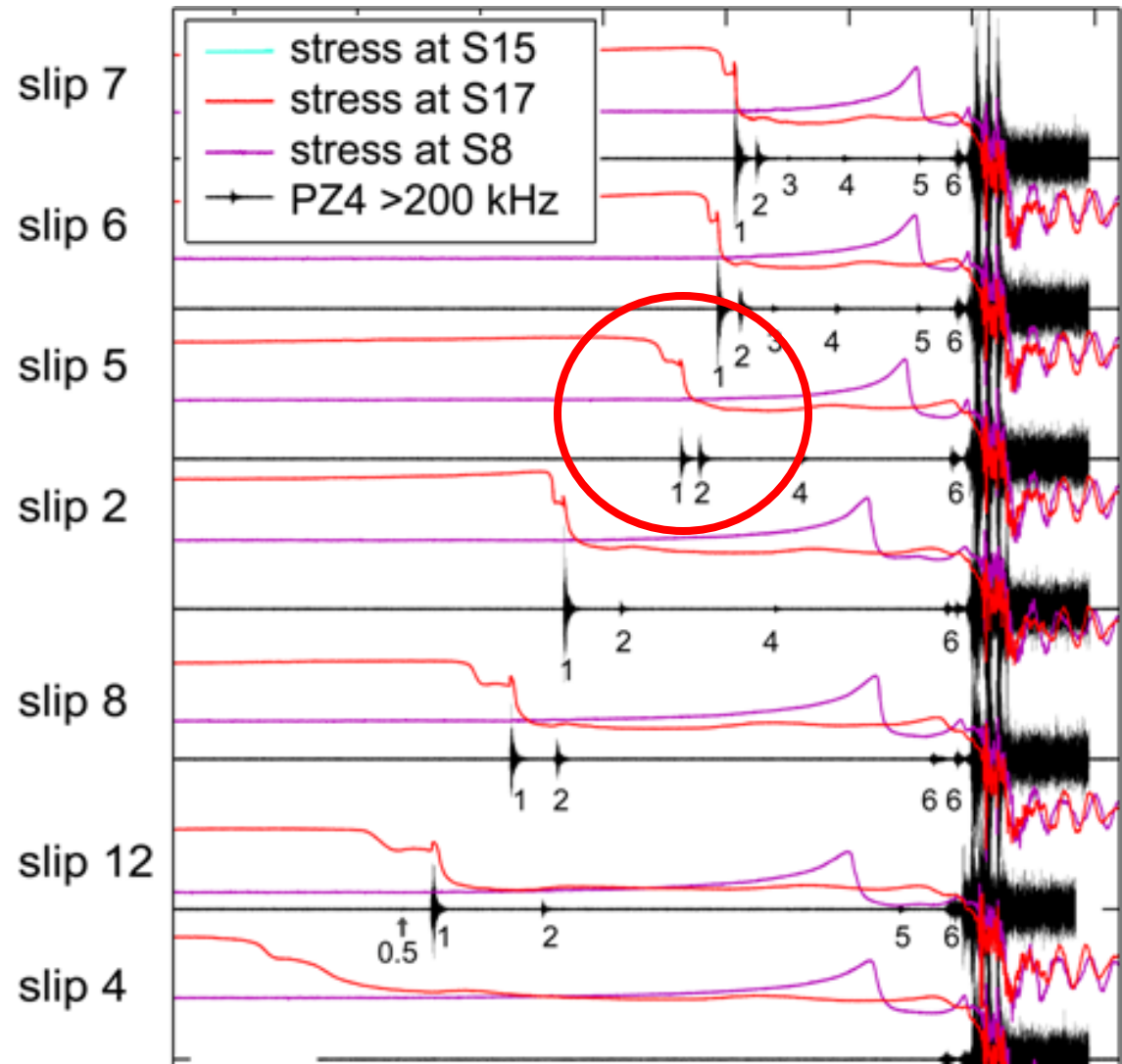
- Foreshock modeled as a $\sim 8 \mu\text{s}$ pulse in moment rate, more complexity later.

Patch 1 foreshock

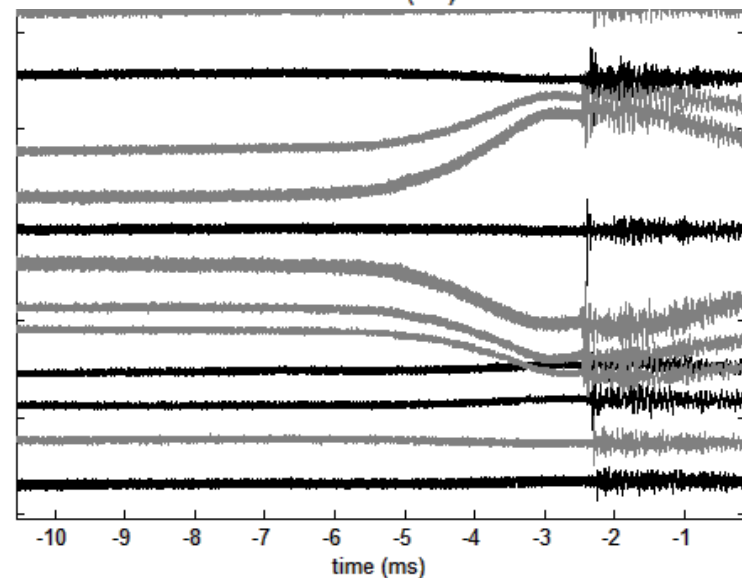
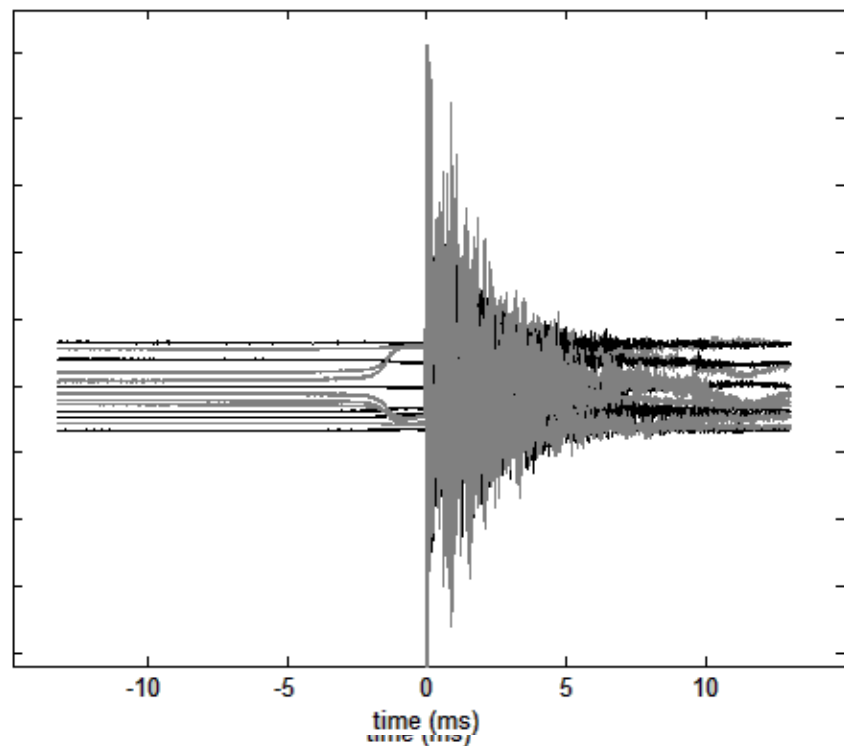
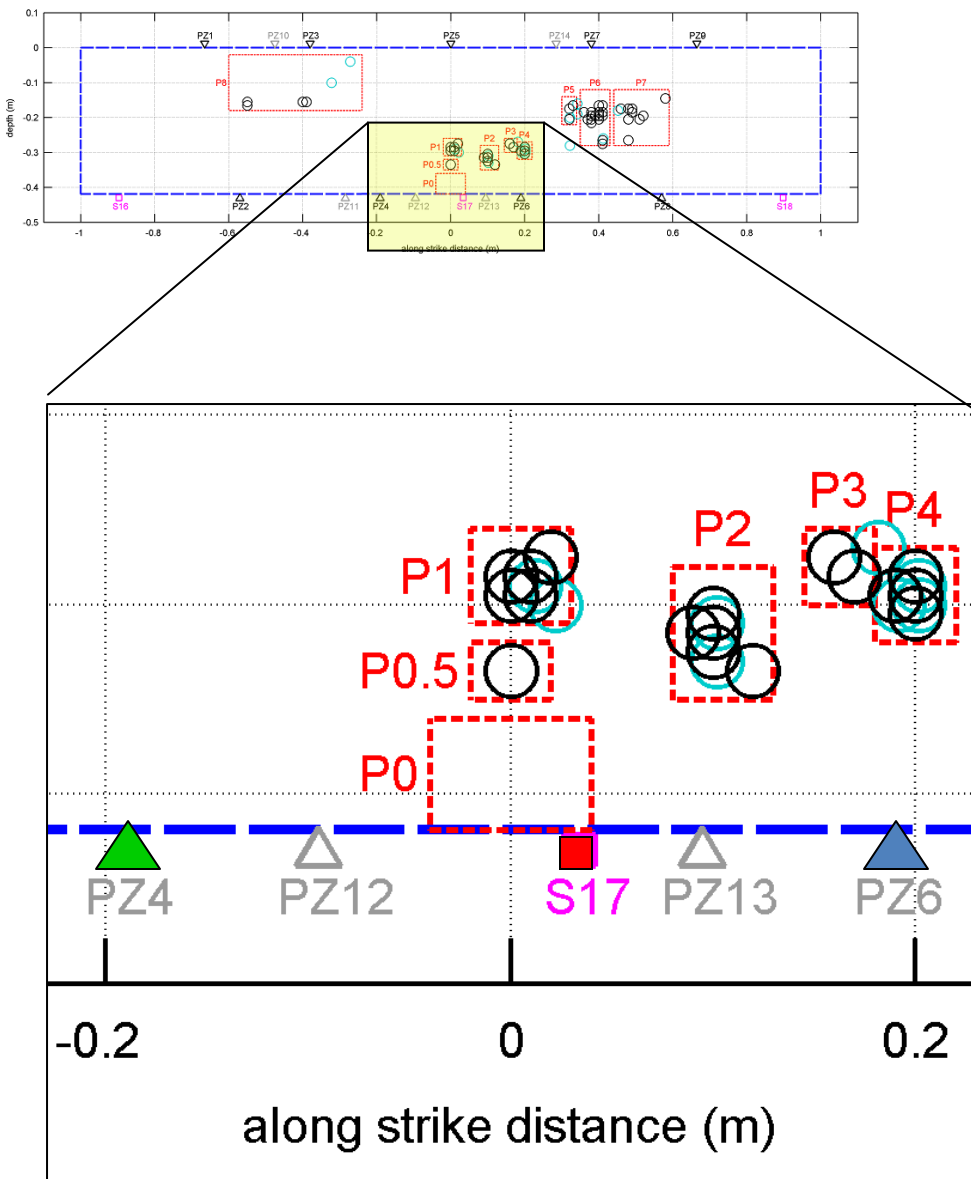
patch 1 foreshock from Slip 12 Jan 2012 exps



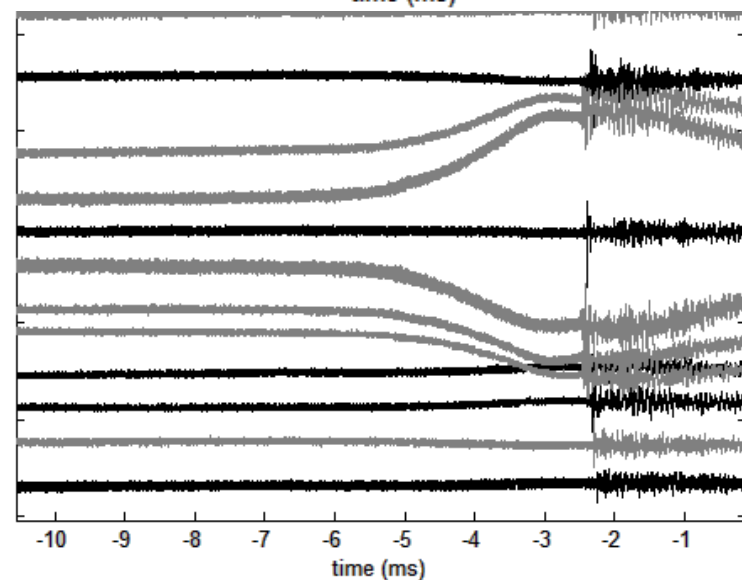
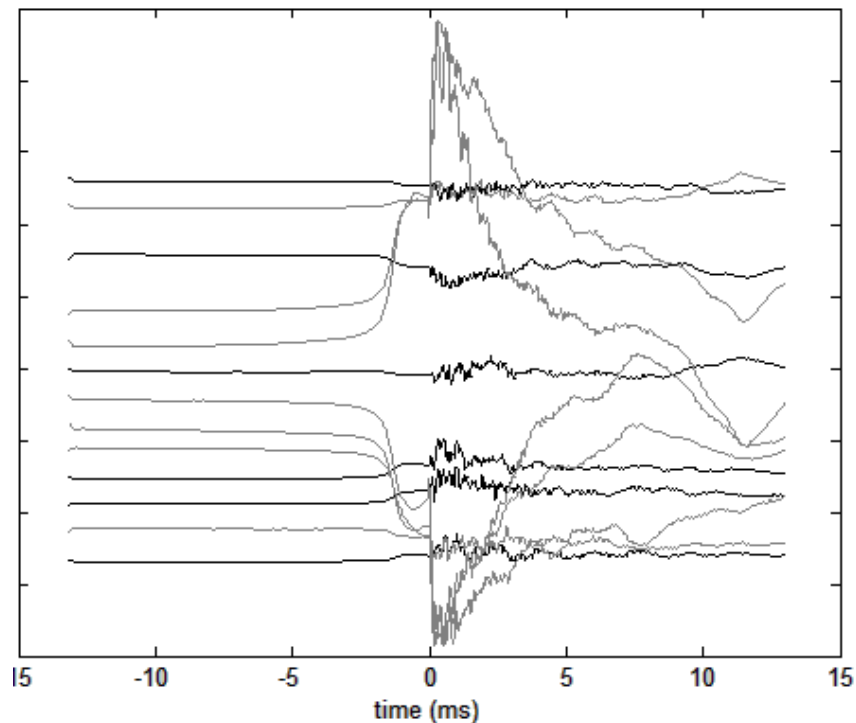
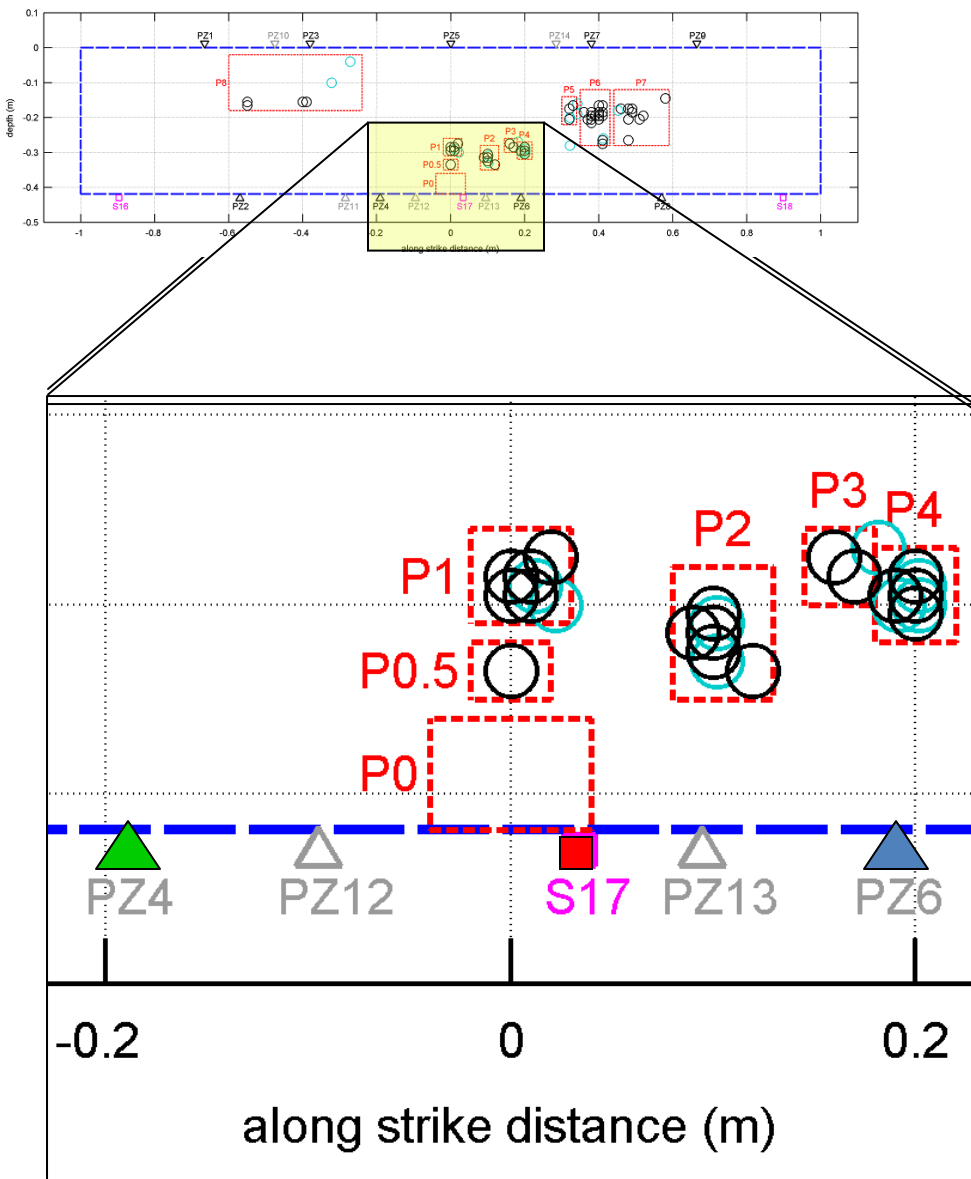
Patch 1 foreshock/stress interactions

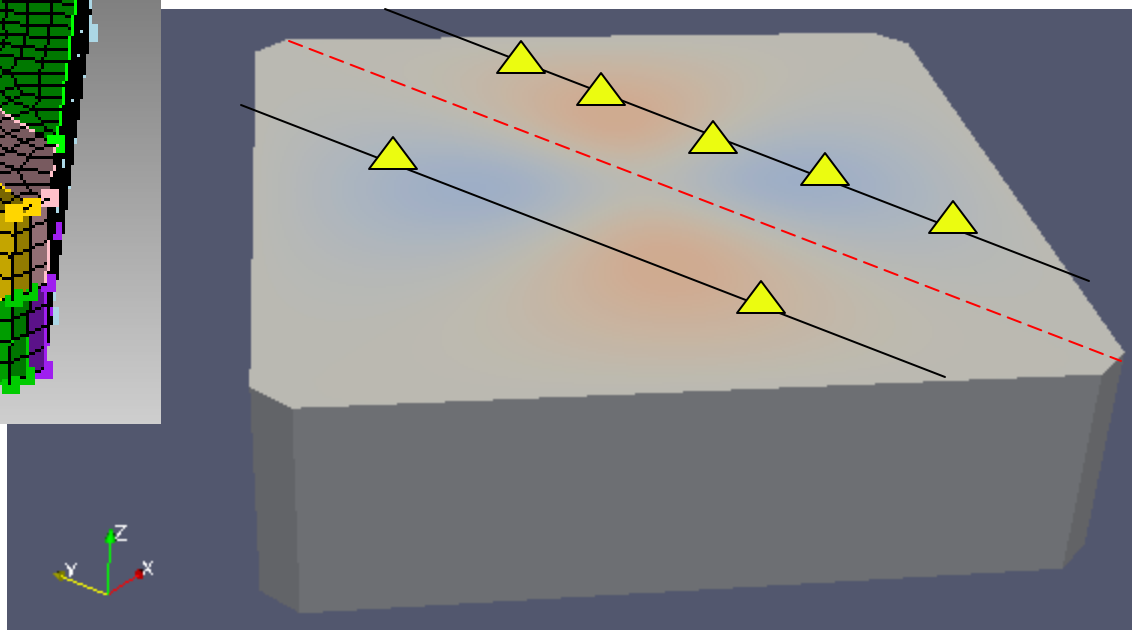
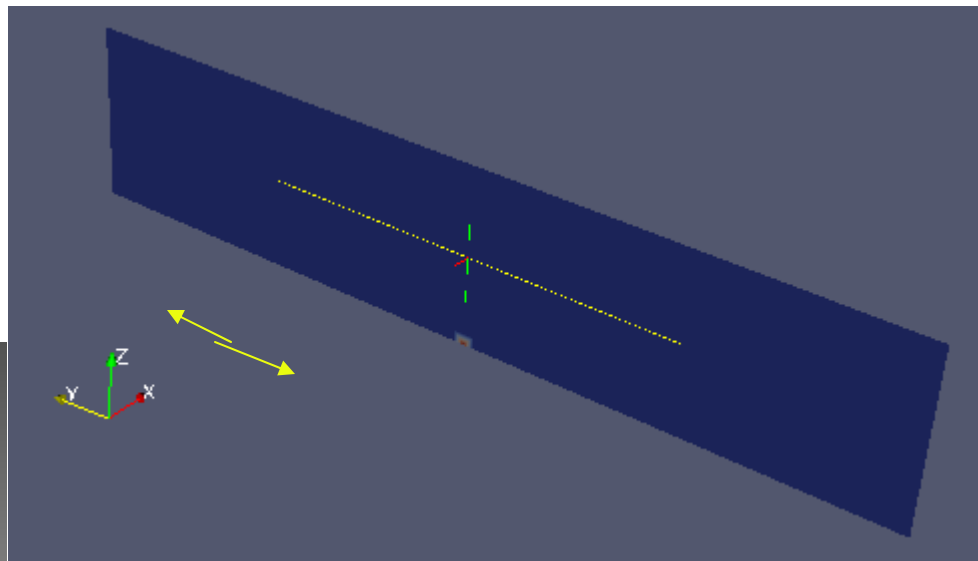
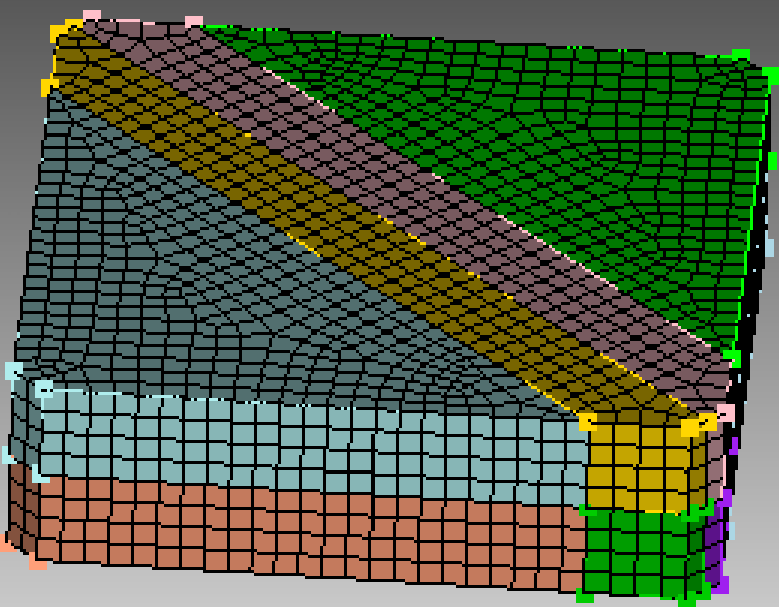


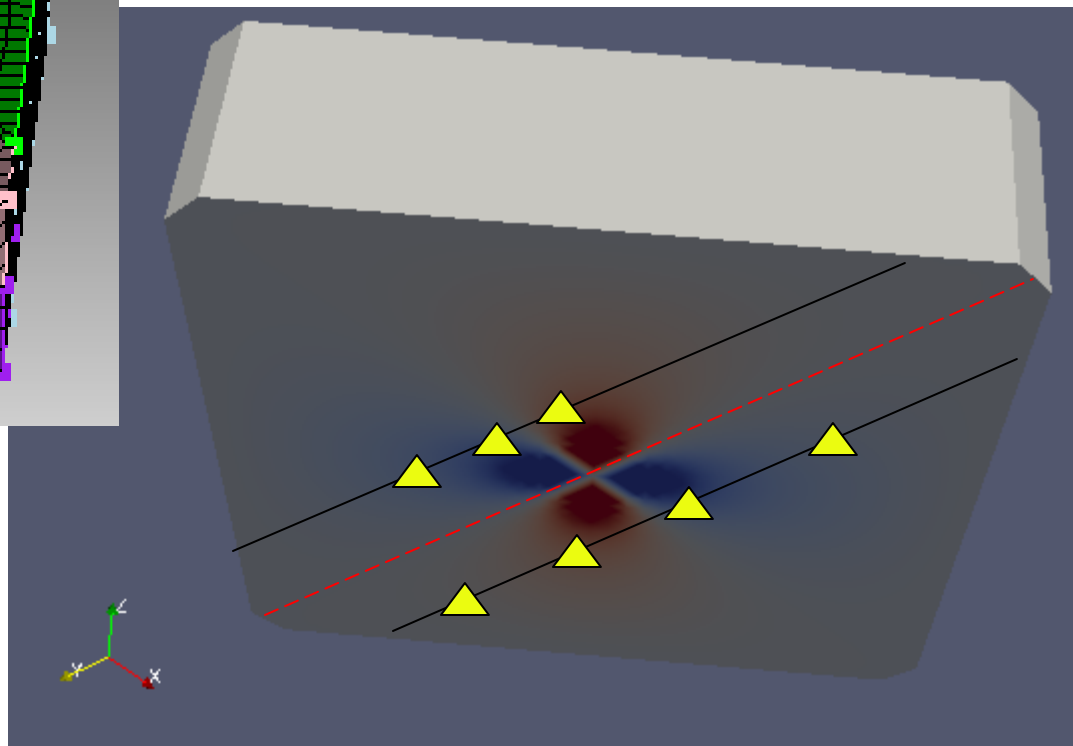
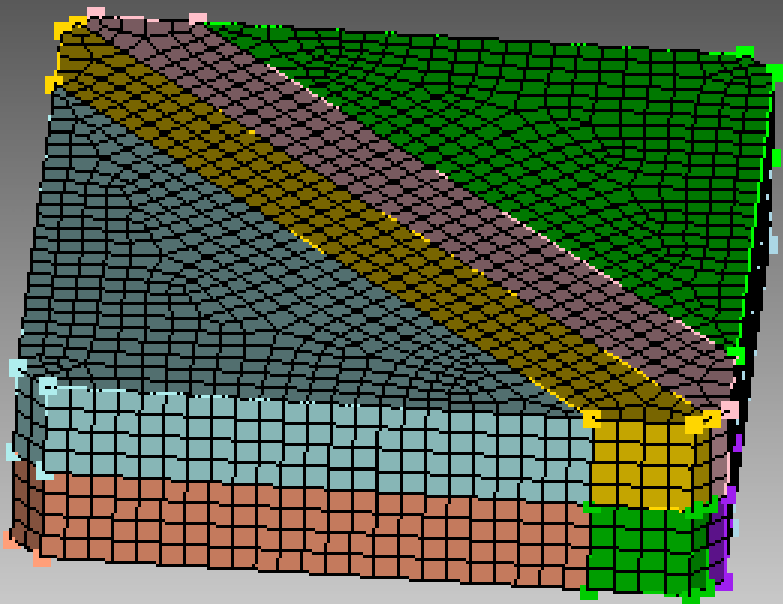
Patch 1 interactions

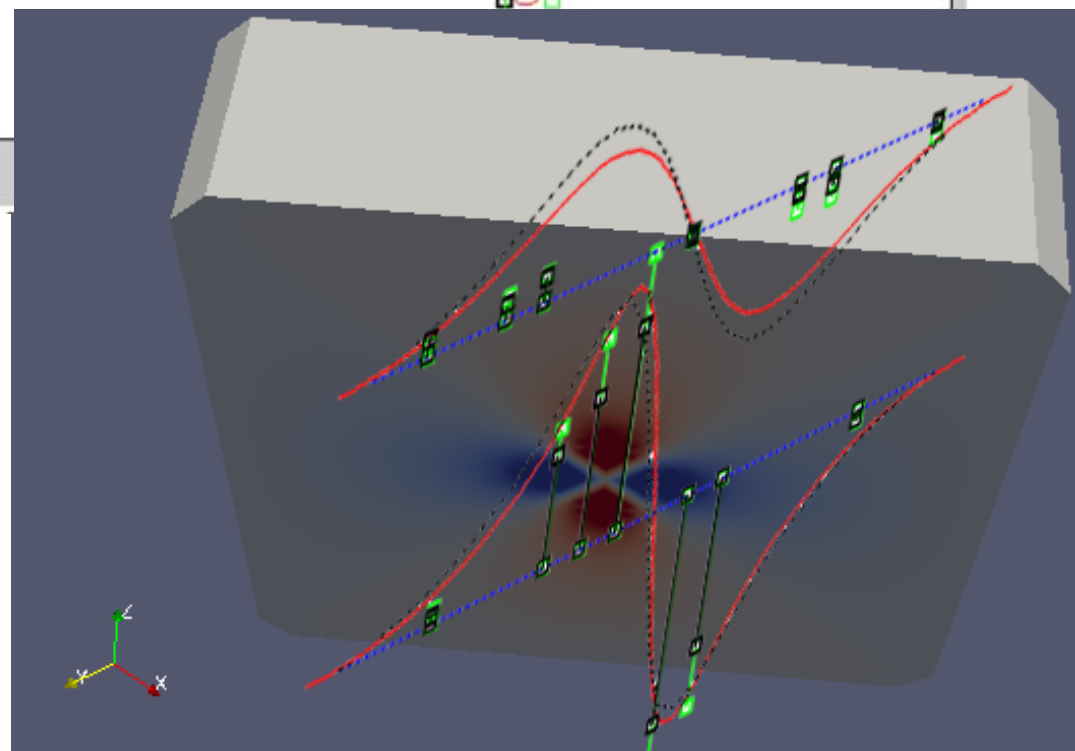
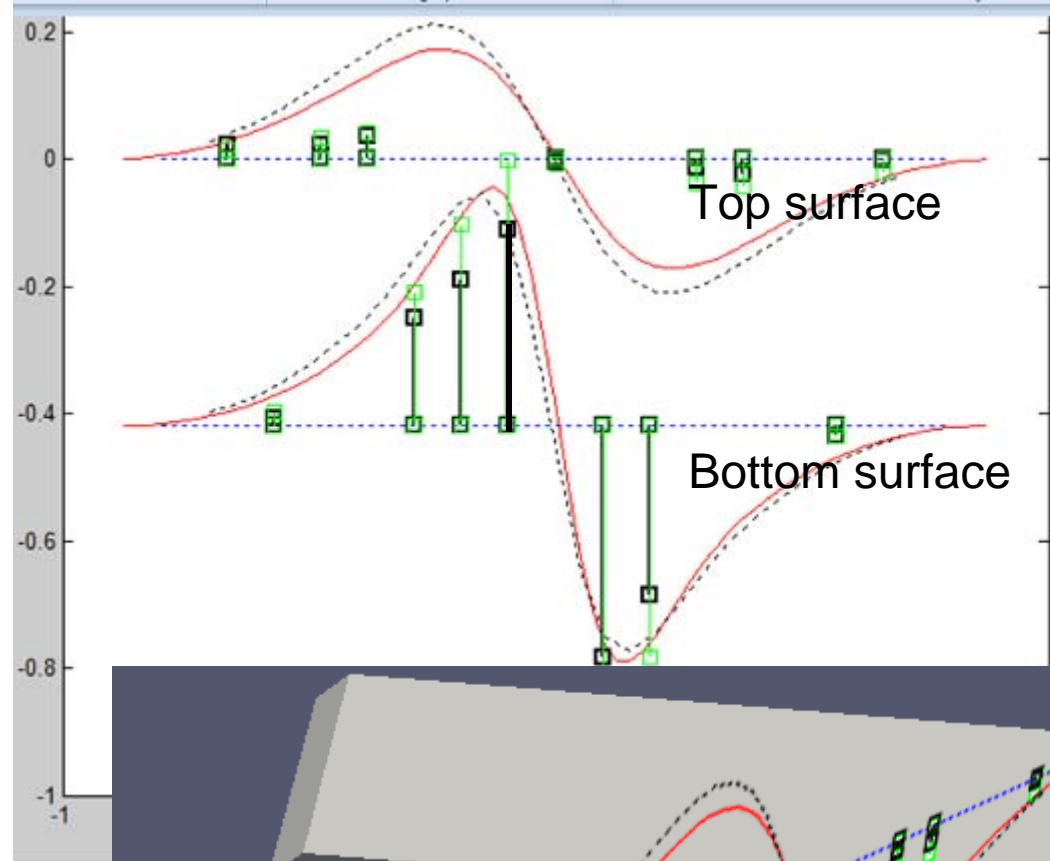
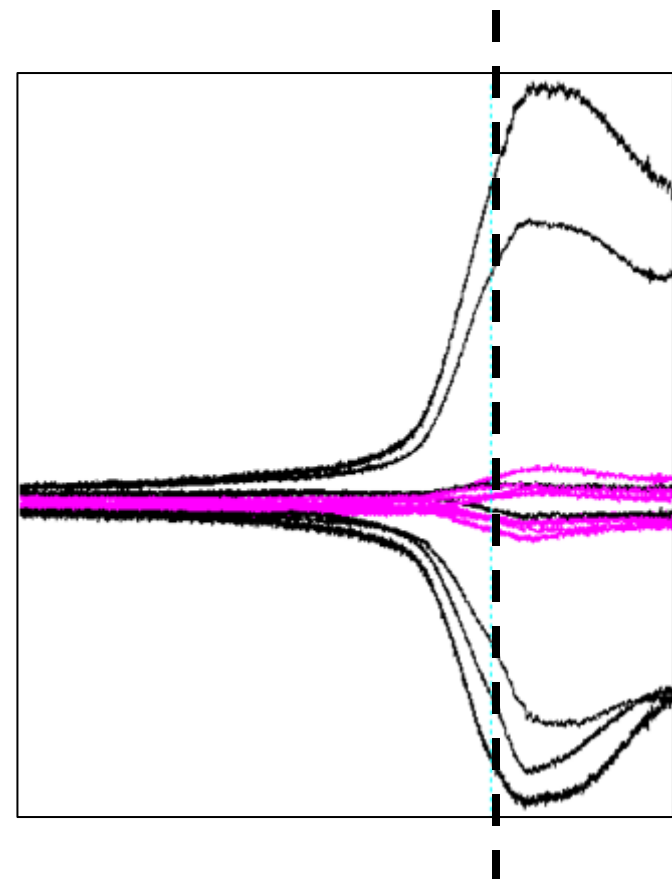


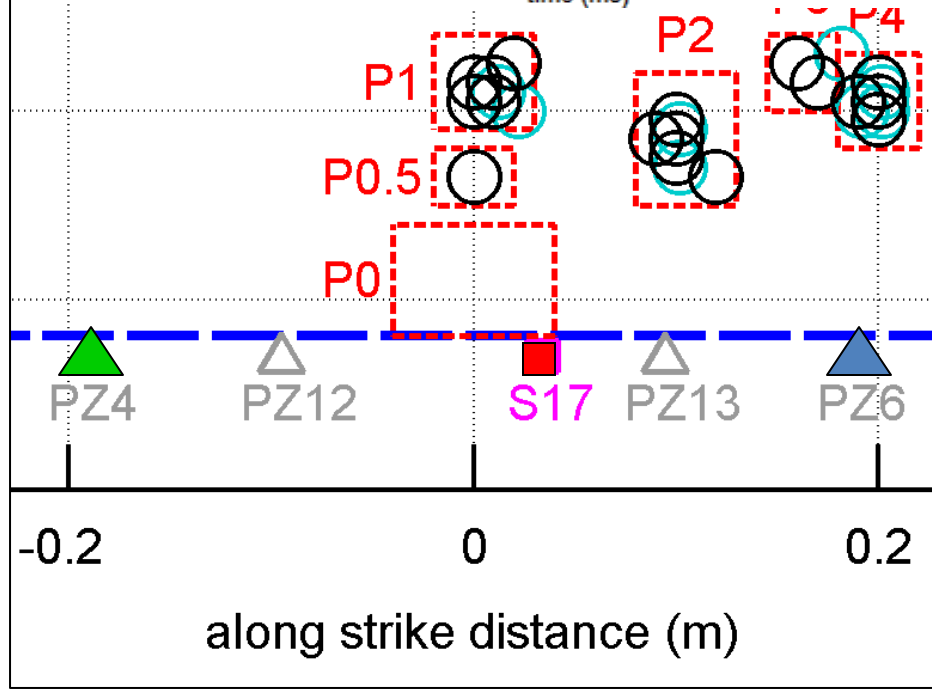
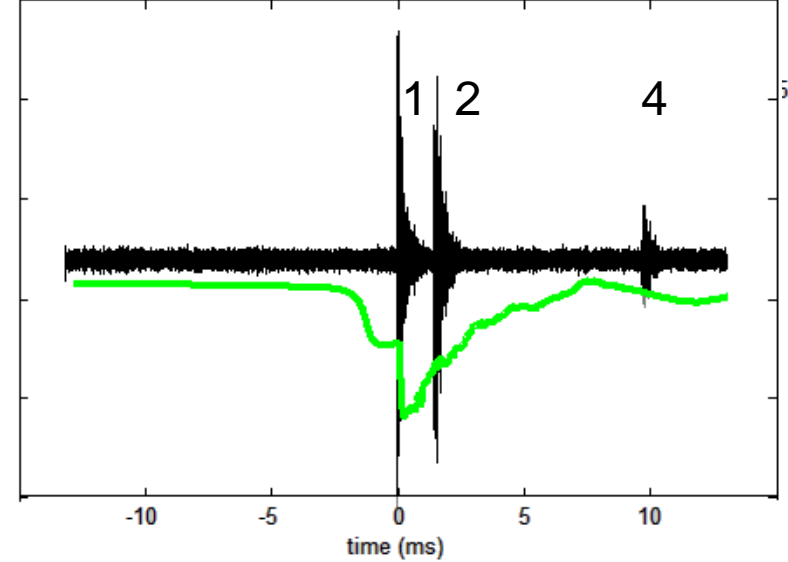
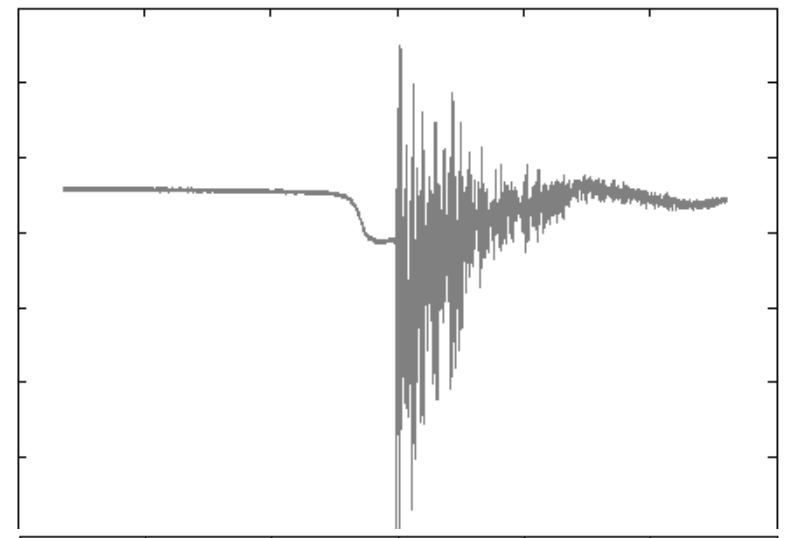
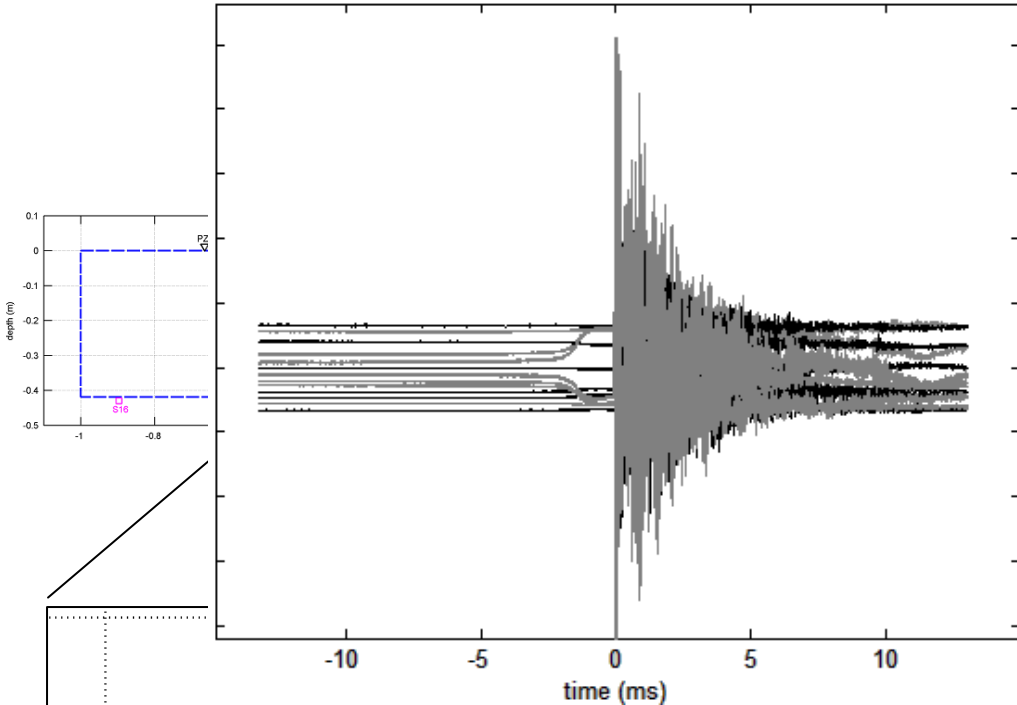
Patch 1 interactions

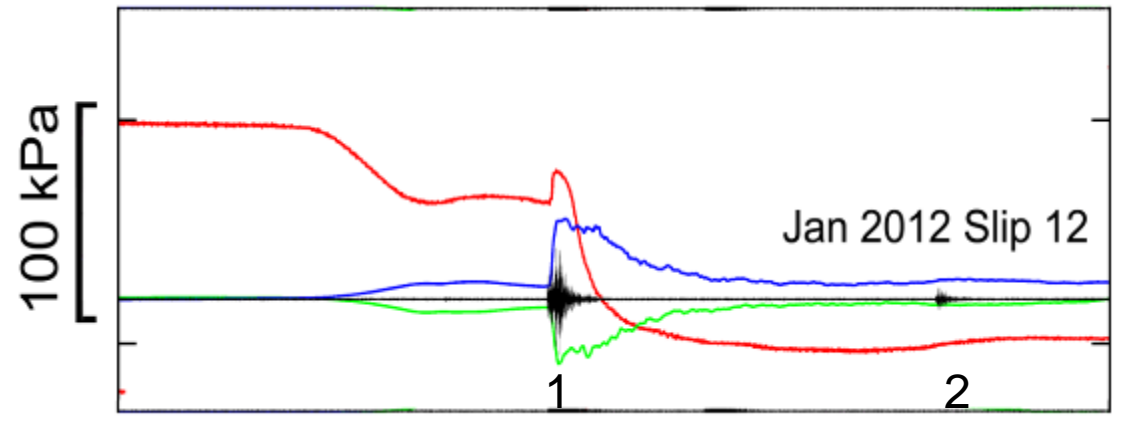
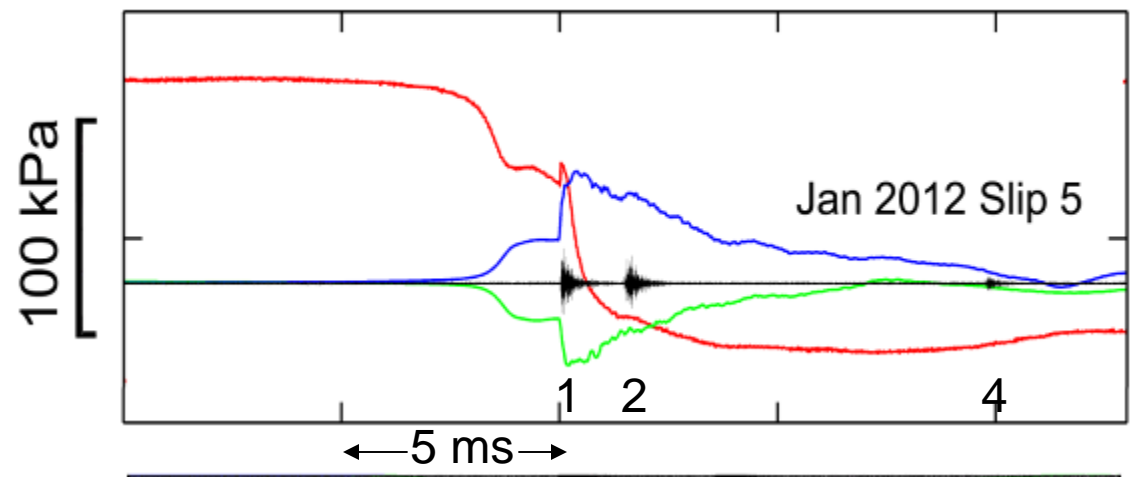
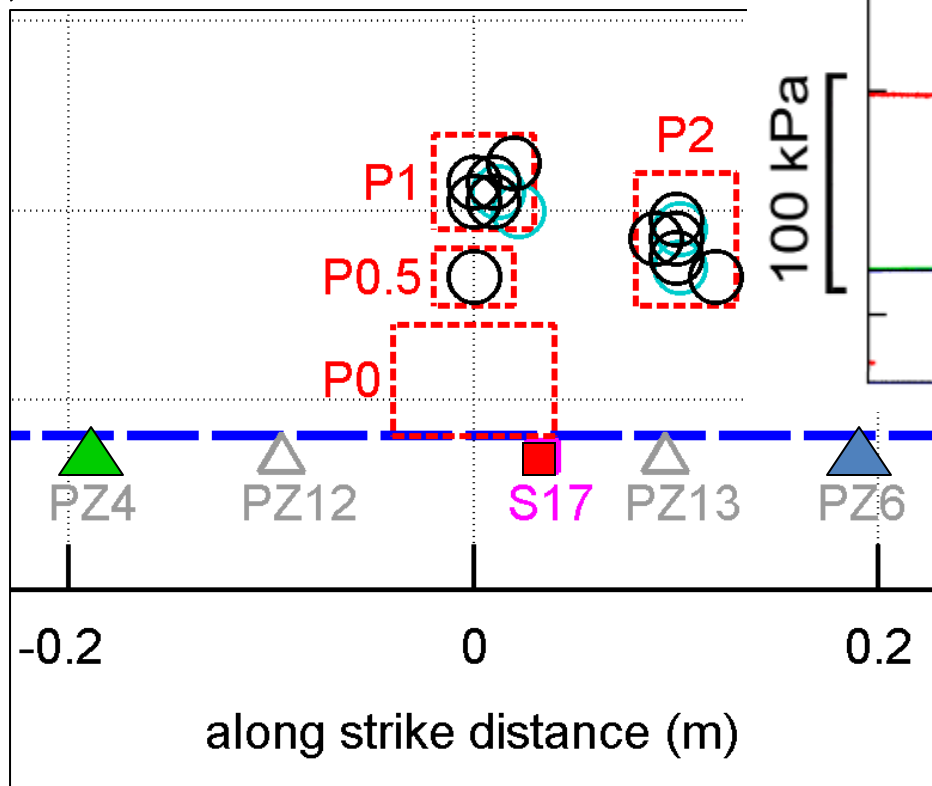
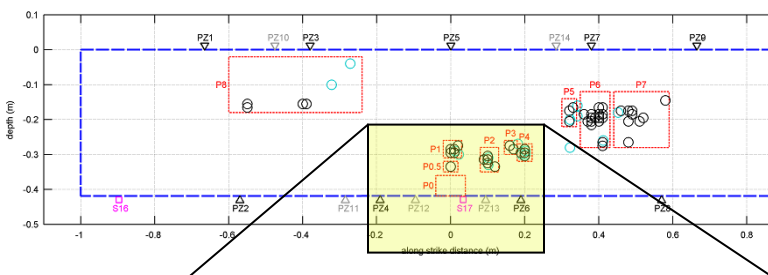


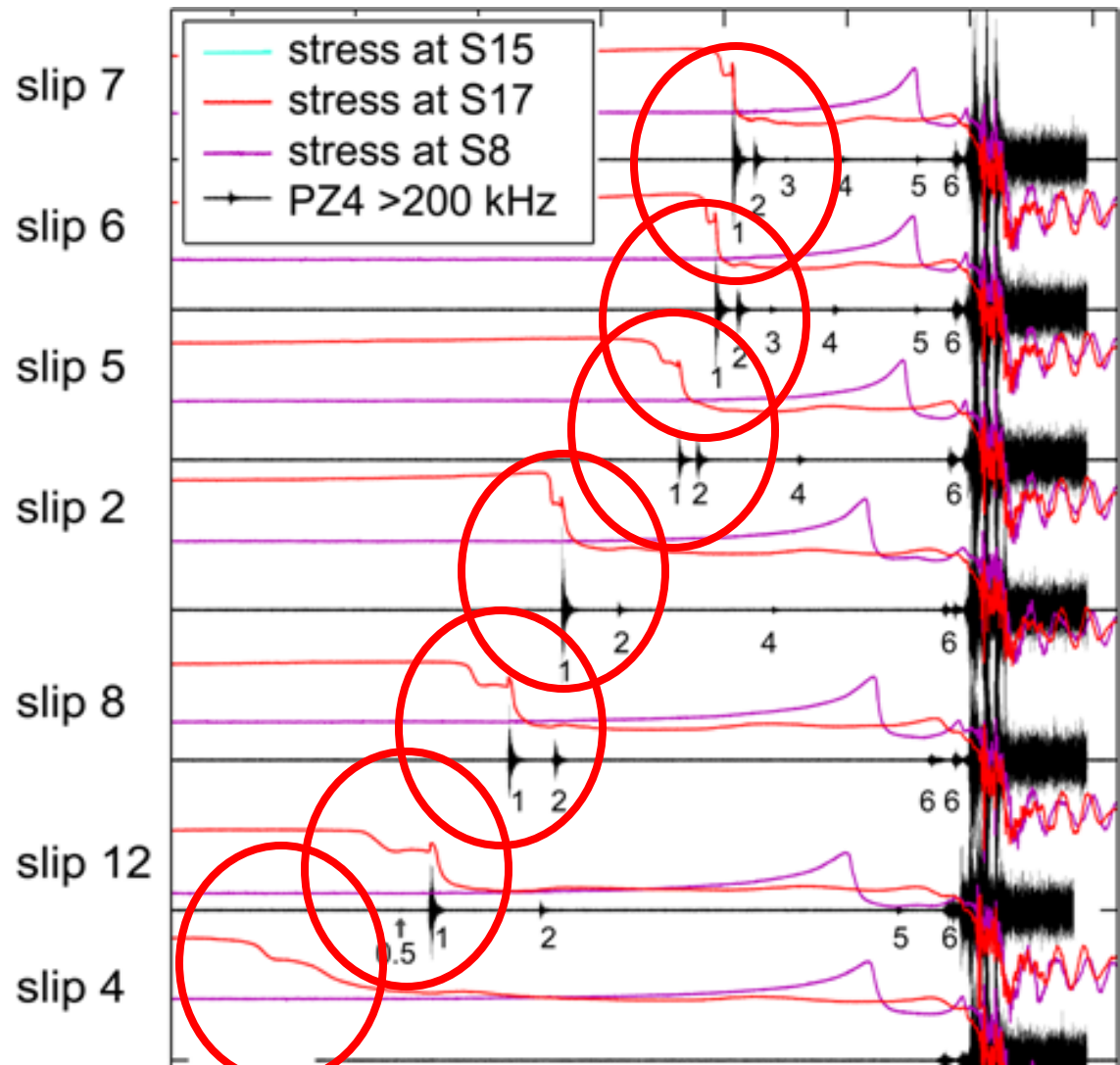


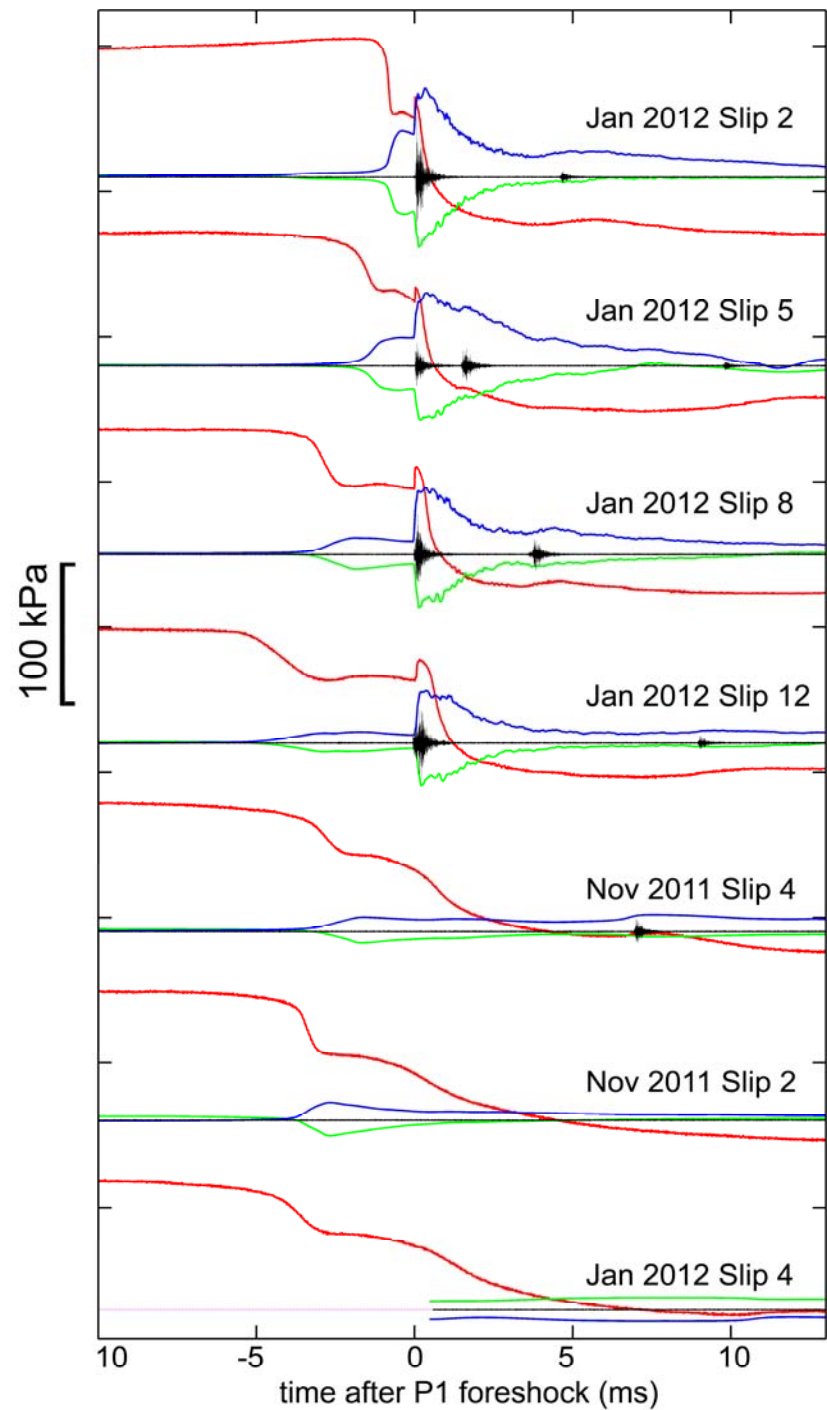
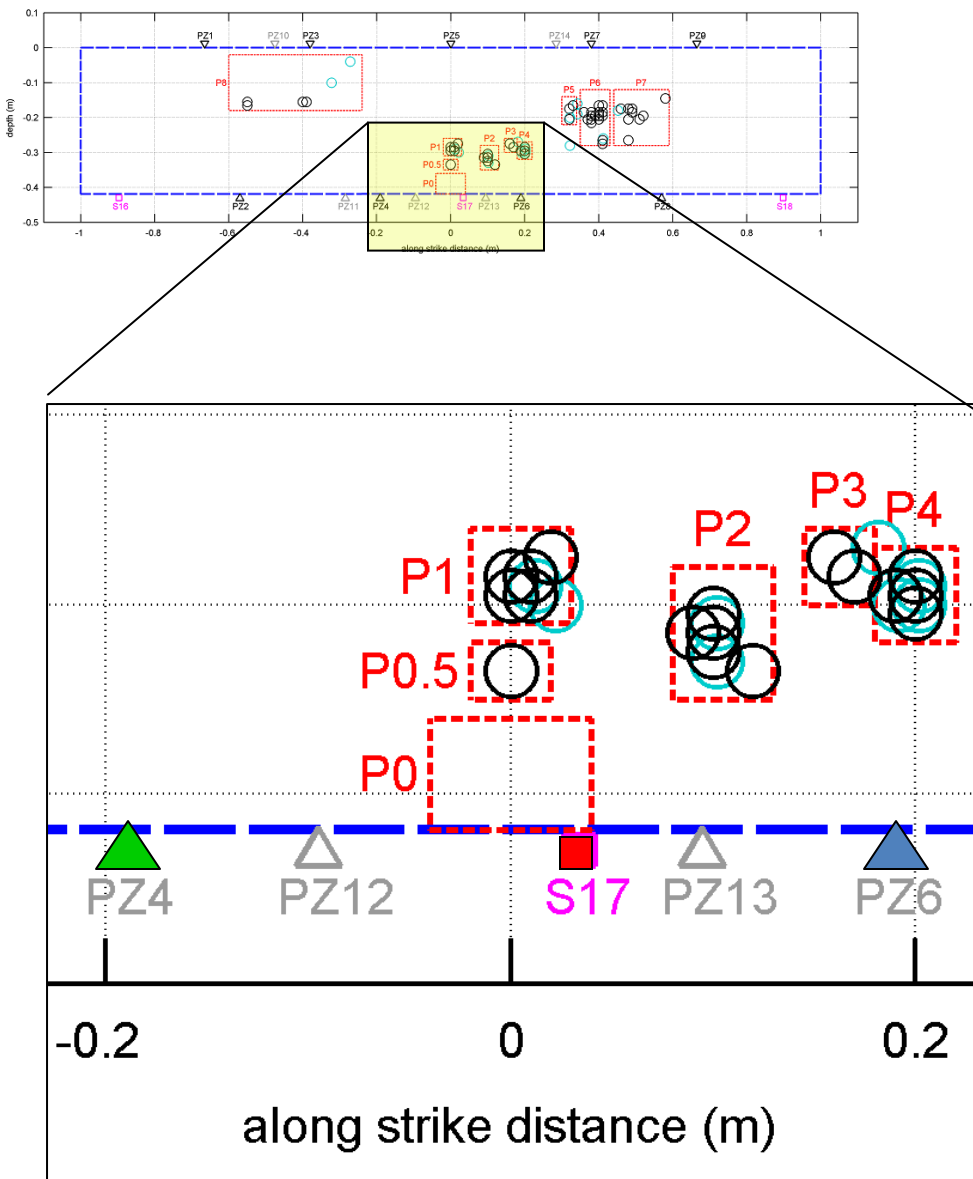










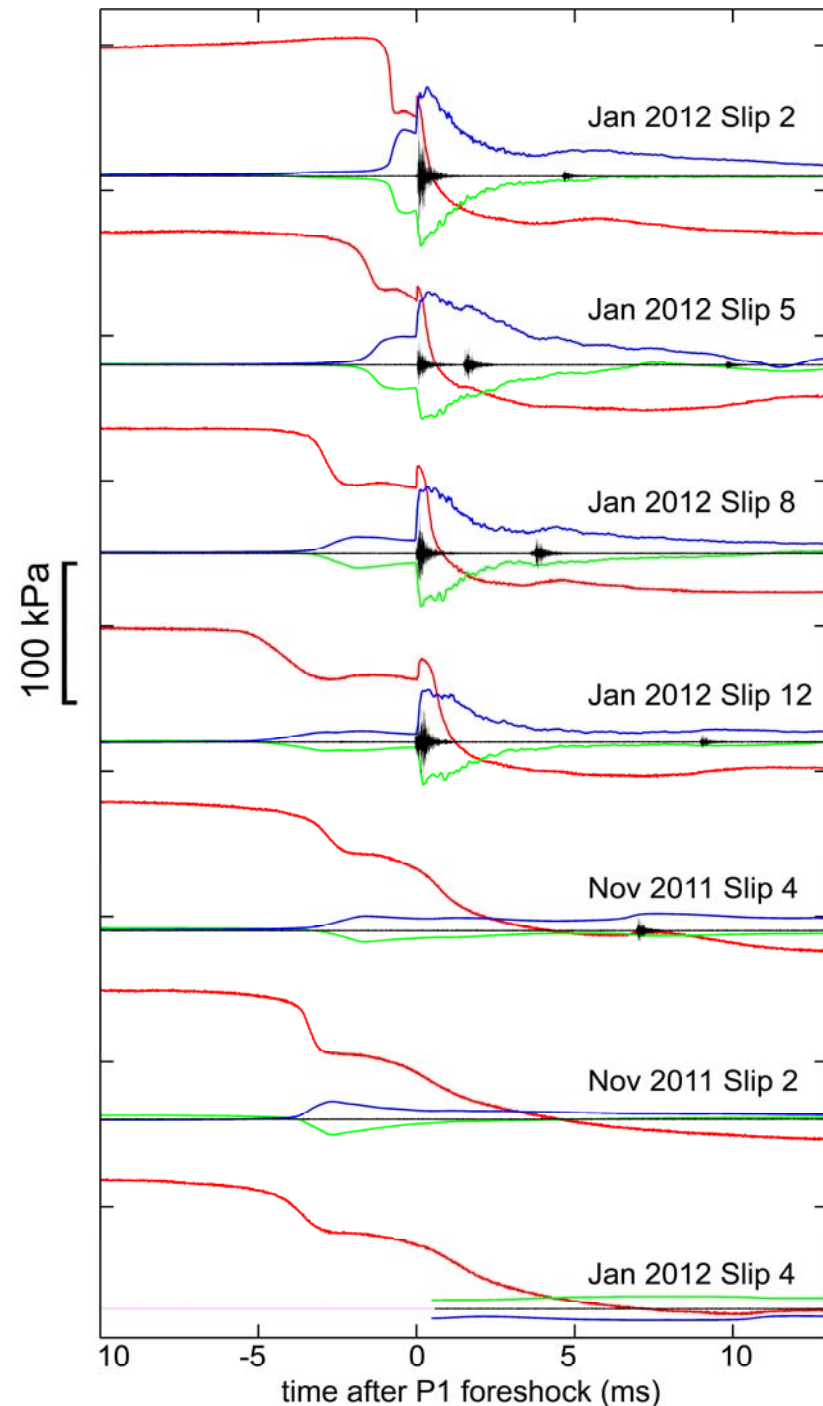
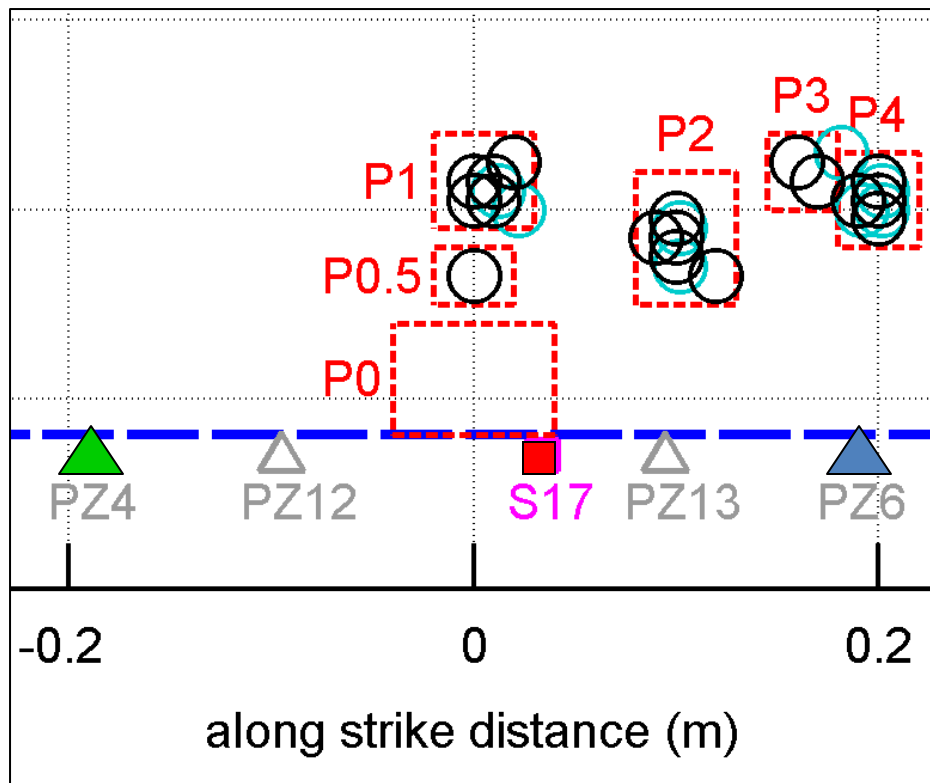


Aseismic slip rapidly loads foreshock patch

Slip transient is slowed or “hung up” by foreshock patch

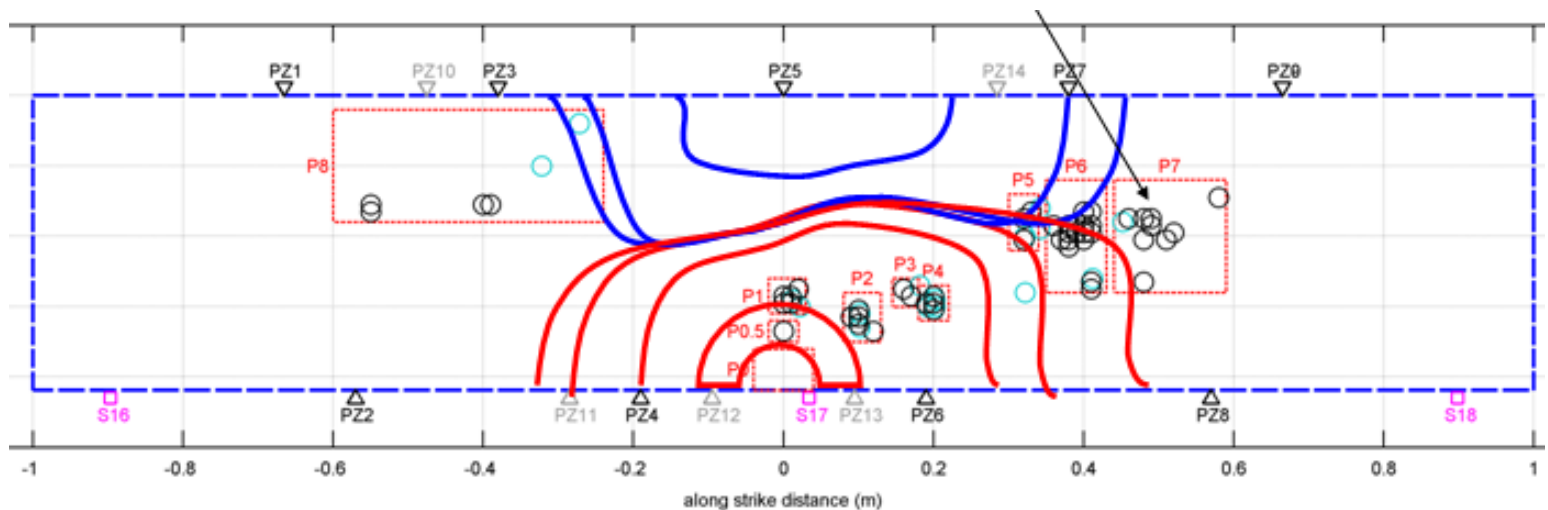
The foreshock accelerates the aseismic slip transient

If the foreshock patch is loaded more slowly, it will slip aseismically.



Summary of Observations

- Foreshocks represent local instabilities within a larger, quasi-stable slipping region consistent with nucleation models.
- Much of the fault is aseismic
- Repeating foreshocks means strength heterogeneity is required for foreshocks
- Transition from aseismic to seismic is very abrupt



Scaling of lab results?

- If we assume no scaling at all...
 - A sequence of M_w -4 to -6 foreshocks 20 ms prior to a $M_w \sim 0$ event.
- Lab versus Field:
 - Same --- elastic moduli, density, wave velocity
 - Different ---mineralogy, temp, pressure, fluids?, roughness
- Scaling of D_c ?

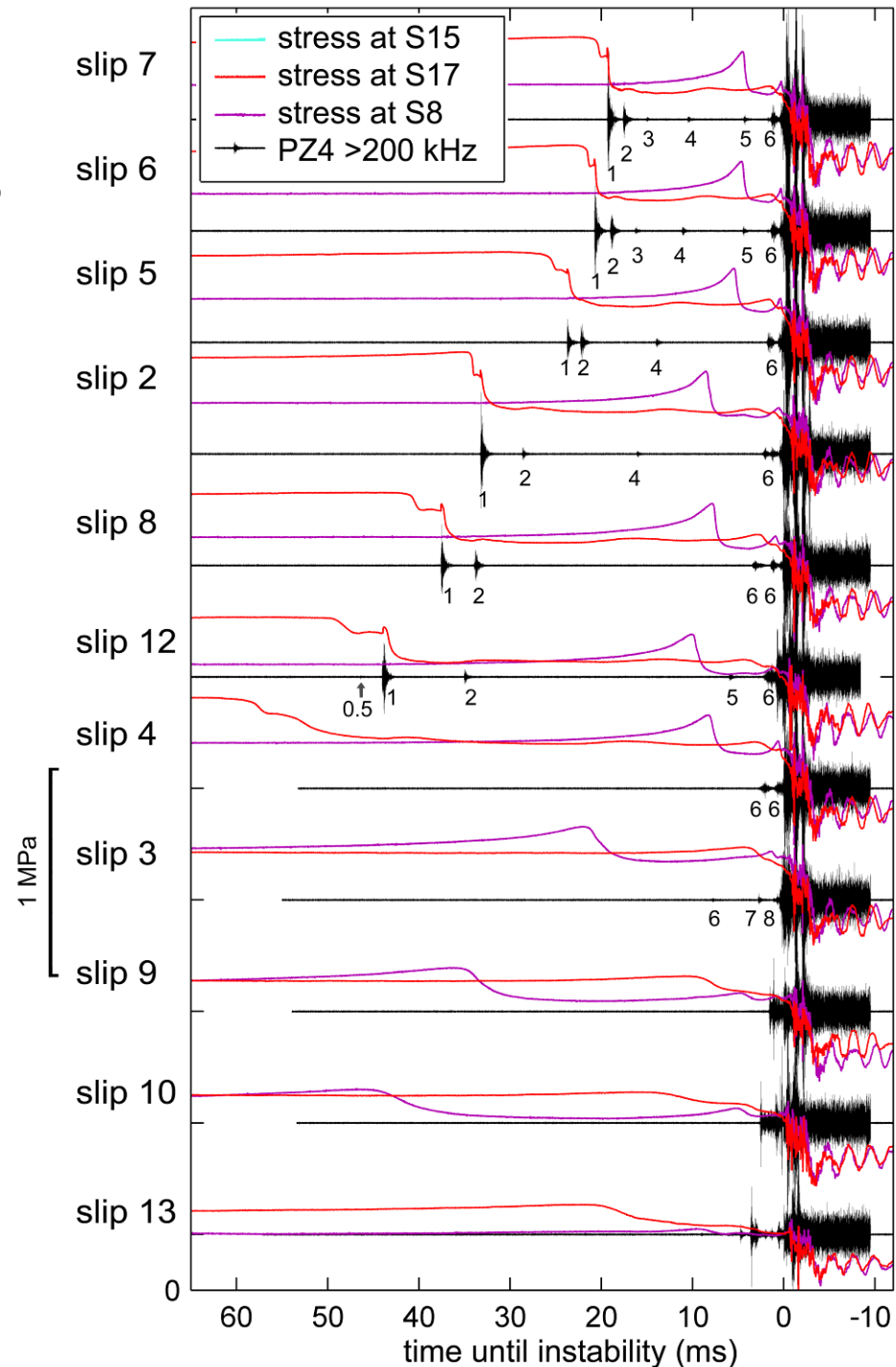
Scaling of D_c ?

- Natural faults may be rougher and have a thicker layer of fault gouge
 - D_c ($\sim \mu\text{m}$ lab, mm field?)
- Nucleation zone size, slip, and duration increases with D_c .

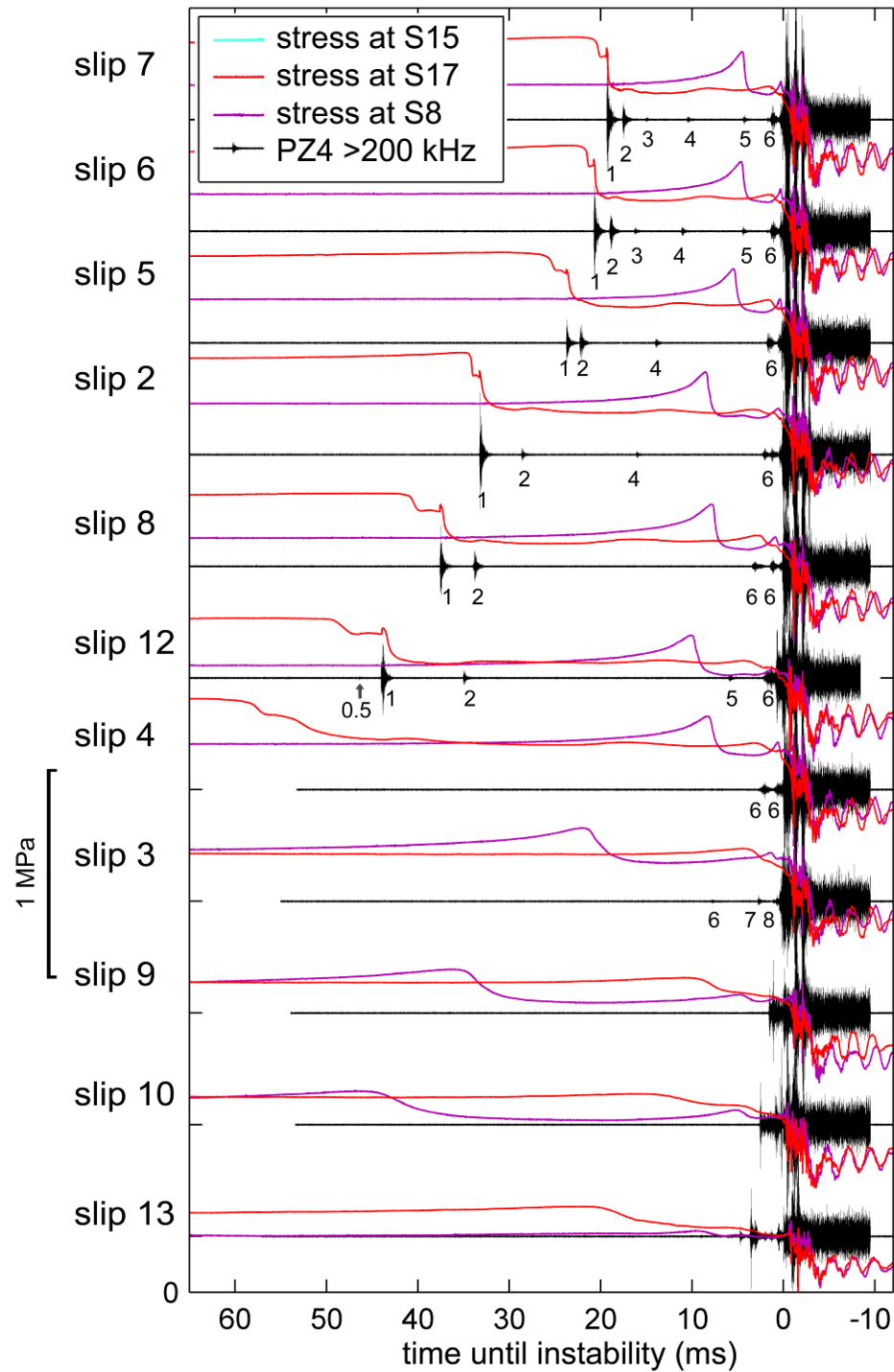


Foreshock Models

- ETAS versus “Preslip” foreshock model?
 - Foreshocks are certainly driven by underlying aseismic slip
 - Foreshocks also accelerate the aseismic slip, and therefore indirectly trigger more foreshocks



Thanks!



- What causes the variation in nucleation rates?
 - Initial stress distribution (left by the last event)
 - Healing!

