

Flow Laws for the Lower Crust and Upper Mantle

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Constitutive Equations / Flow Laws

$$\dot{\varepsilon} = \dot{\varepsilon}(\sigma, T, P, f_{\text{H}_2\text{O}}, a_{\text{ox}}, \dots, S, d, \dots, \Phi, \phi, \dots)$$

$$\eta = \eta(\sigma, T, P, f_{\text{H}_2\text{O}}, a_{\text{ox}}, \dots, S, d, \dots, \Phi, \phi, \dots)$$

Flow Laws for Steady-State Deformation – Diffusion Creep

Grain-matrix diffusion

$$\dot{\epsilon}_{\text{NH}} = \alpha_{\text{NH}} \frac{\sigma V_{\text{m}}}{RT} \frac{D_{\text{gm}}}{d^2}$$

$$D_{\text{gm}} = D_{\text{gm}}^{\circ} \exp\left(-\frac{\Delta E_{\text{gm}} + P\Delta V_{\text{gm}}}{RT}\right)$$

$$= D_{\text{gm}}^{\circ} \exp\left(-\frac{\Delta H_{\text{gm}}}{RT}\right)$$

Grain-boundary diffusion

$$\dot{\epsilon}_{\text{C}} = \alpha_{\text{C}} \frac{\sigma V_{\text{m}}}{RT} \frac{\delta D_{\text{gb}}}{d^3}$$

$$D_{\text{gb}} = D_{\text{gb}}^{\circ} \exp\left(-\frac{\Delta E_{\text{gb}} + P\Delta V_{\text{gb}}}{RT}\right)$$

$$= D_{\text{gb}}^{\circ} \exp\left(-\frac{\Delta H_{\text{gb}}}{RT}\right)$$

Flow Laws for Steady-State Deformation – Diffusion Creep

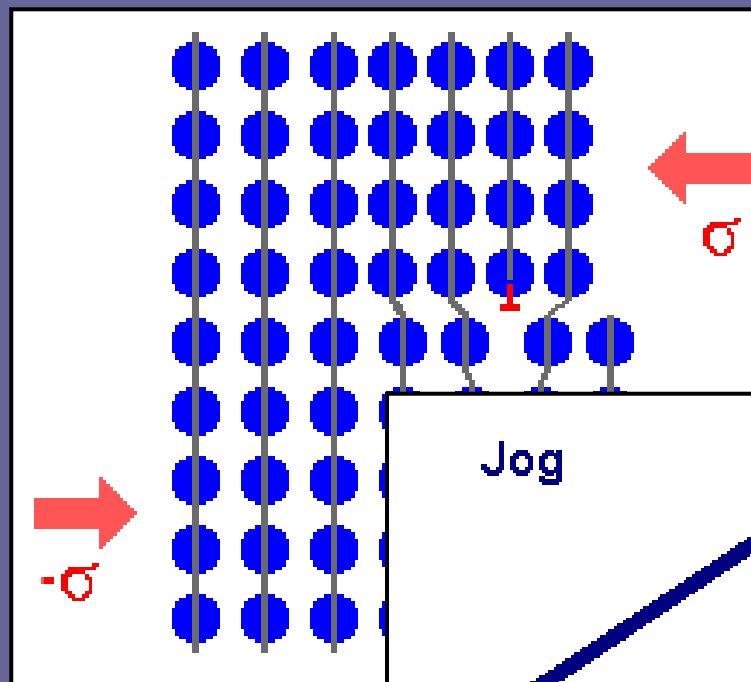
$$\dot{\epsilon}_{\text{diff}} = 14 \left(\frac{\sigma V_m}{RT} \right) \left(D_{\text{gm}} + \frac{\pi \delta D_{\text{gb}}}{d} \right) \left(\frac{1}{d^2} \right)$$

$$D = D^0 \exp \left(-\frac{\Delta H}{RT} \right)$$

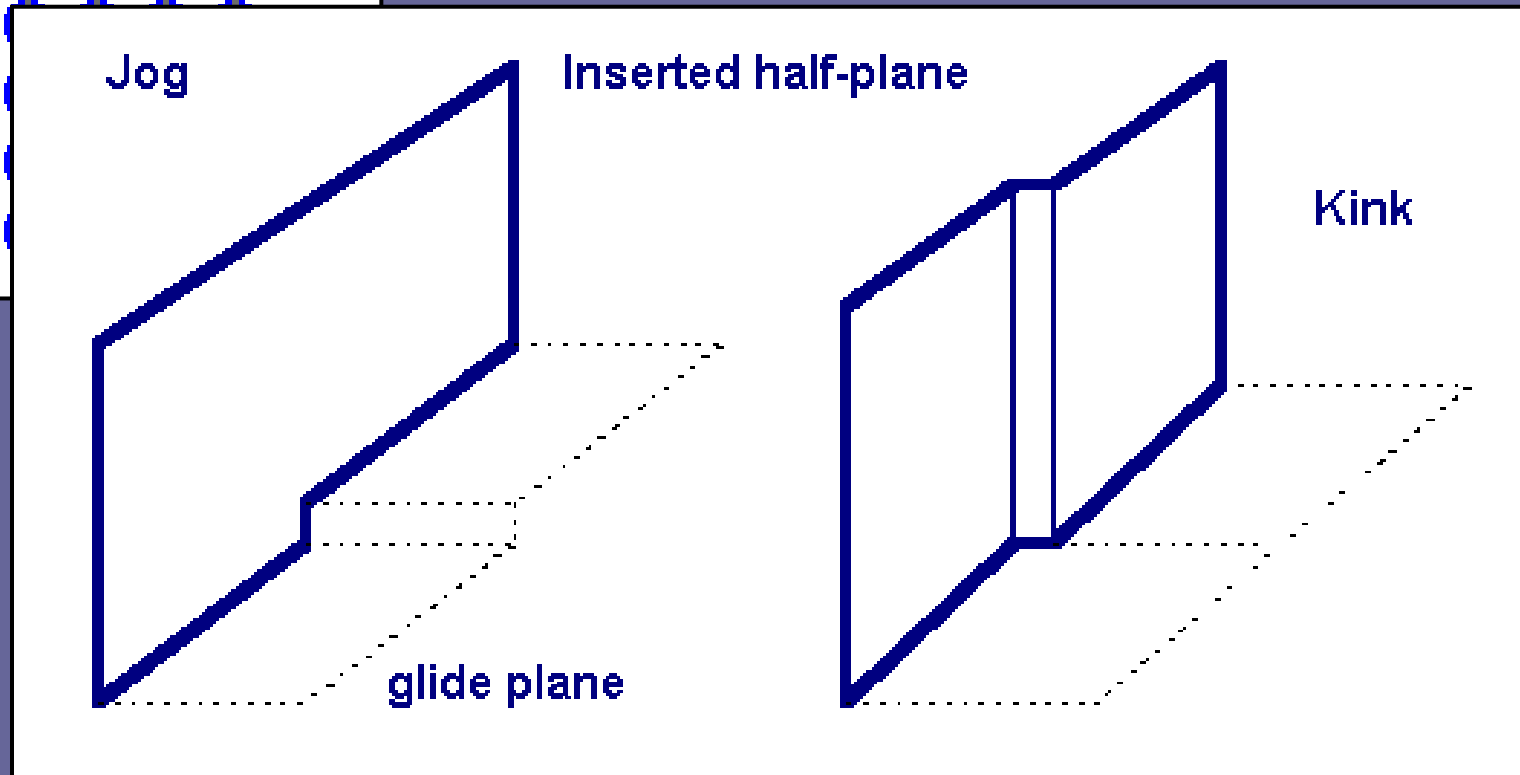
$$\Delta H_{\text{gb}} < \Delta H_{\text{gm}}$$

slowest ion along fastest path

Flow Laws for Steady-State Deformation – Dislocation Creep

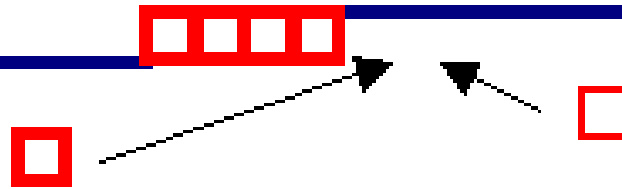


http://www.techfak.uni-kiel.de/matwis/amat/def_en/index.html

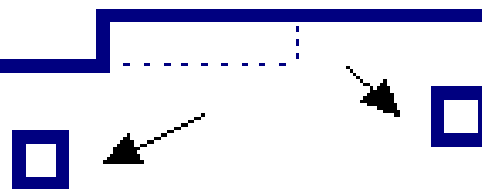


High-Temperature Deformation – Dislocation Climb

**Movement of a jog by
addition of interstitials**



**Movement of a jog by
absorption of vacancies
or emission of interstitials**



Flow Laws for Steady-State Deformation – High-Temperature Dislocation Creep

$$\dot{\varepsilon} = A \frac{\sigma^n}{d^m} f_{\text{O}_2}^p f_{\text{H}_2\text{O}}^q \exp\left(-\frac{Q_{\text{cr}}}{RT}\right)$$

$\dot{\varepsilon} = \rho b \bar{u}$

$\rho \approx \left(\frac{\sigma}{Gb}\right)^2$

$\bar{u} = \frac{l_g + l_c}{t_g + t_c} \approx \frac{l_g}{l_c} u_c$

$\dot{\varepsilon} = 2\pi \frac{GV_m}{RT} \left(\frac{\sigma}{G}\right)^3 \frac{D}{b^2} \frac{1}{\ln(G/\sigma)} \frac{l_g}{l_c}$

$u_c = 2\pi \frac{\sigma V_m}{RT} \frac{D}{b} \frac{1}{\ln(R_o/r_c)}$

Origin of Dependence of Viscosity on Fugacity

$$X_{\text{ion}} D_{\text{ion}} = X_{\text{V}} D_{\text{V}}$$

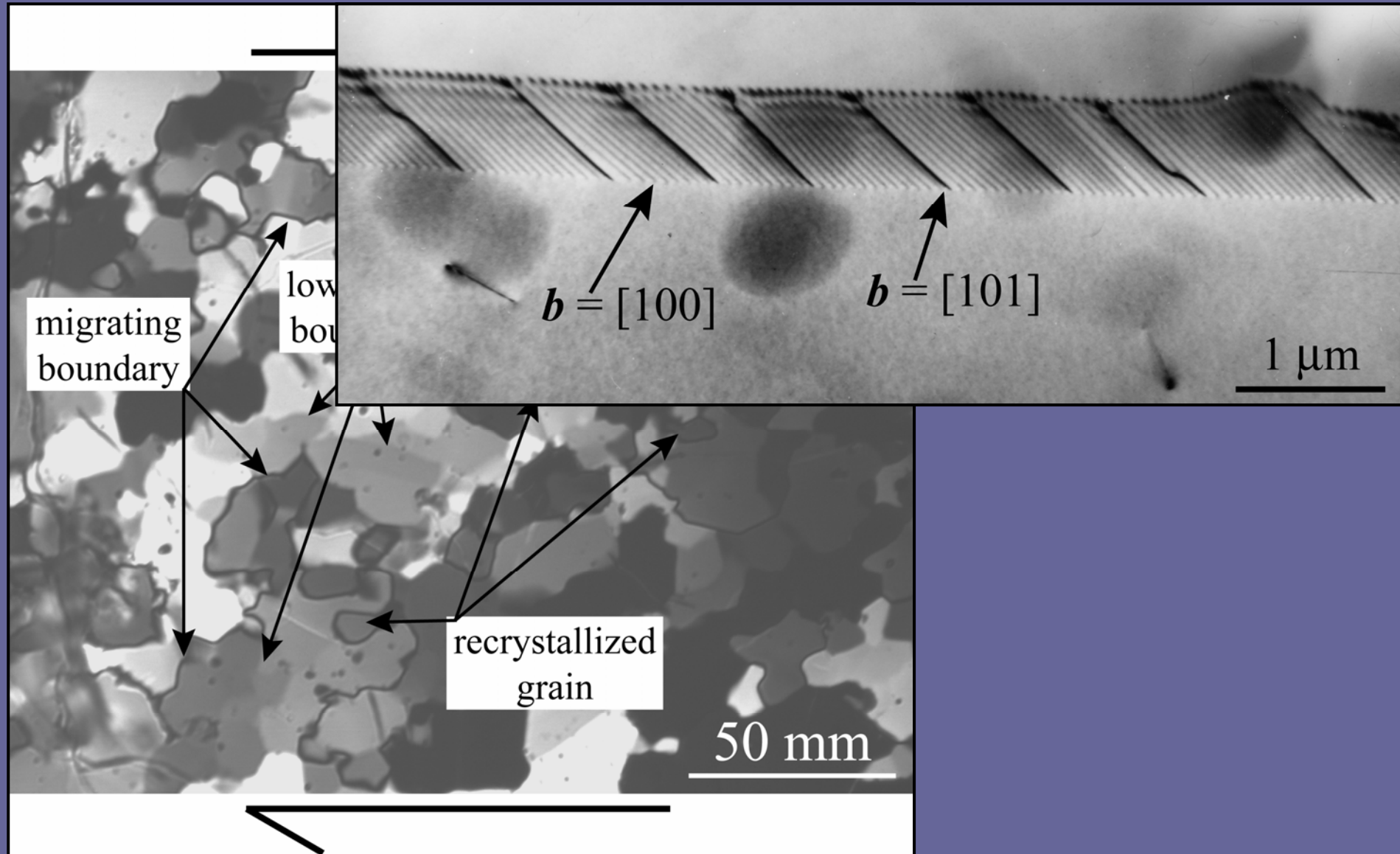
$$X_{\text{ion}} = (1 - X_{\text{V}}) \approx 1$$

$$D_{\text{ion}} \approx X_{\text{V}} D_{\text{V}}$$

$$D_{\text{ion}} \ll D_{\text{V}}$$

$$D_{\text{ion}} \propto X_{\text{V}} \propto f_{\text{O}_2}^p f_{\text{H}_2\text{O}}^q$$

Dislocation-Accommodated Grain Boundary Sliding

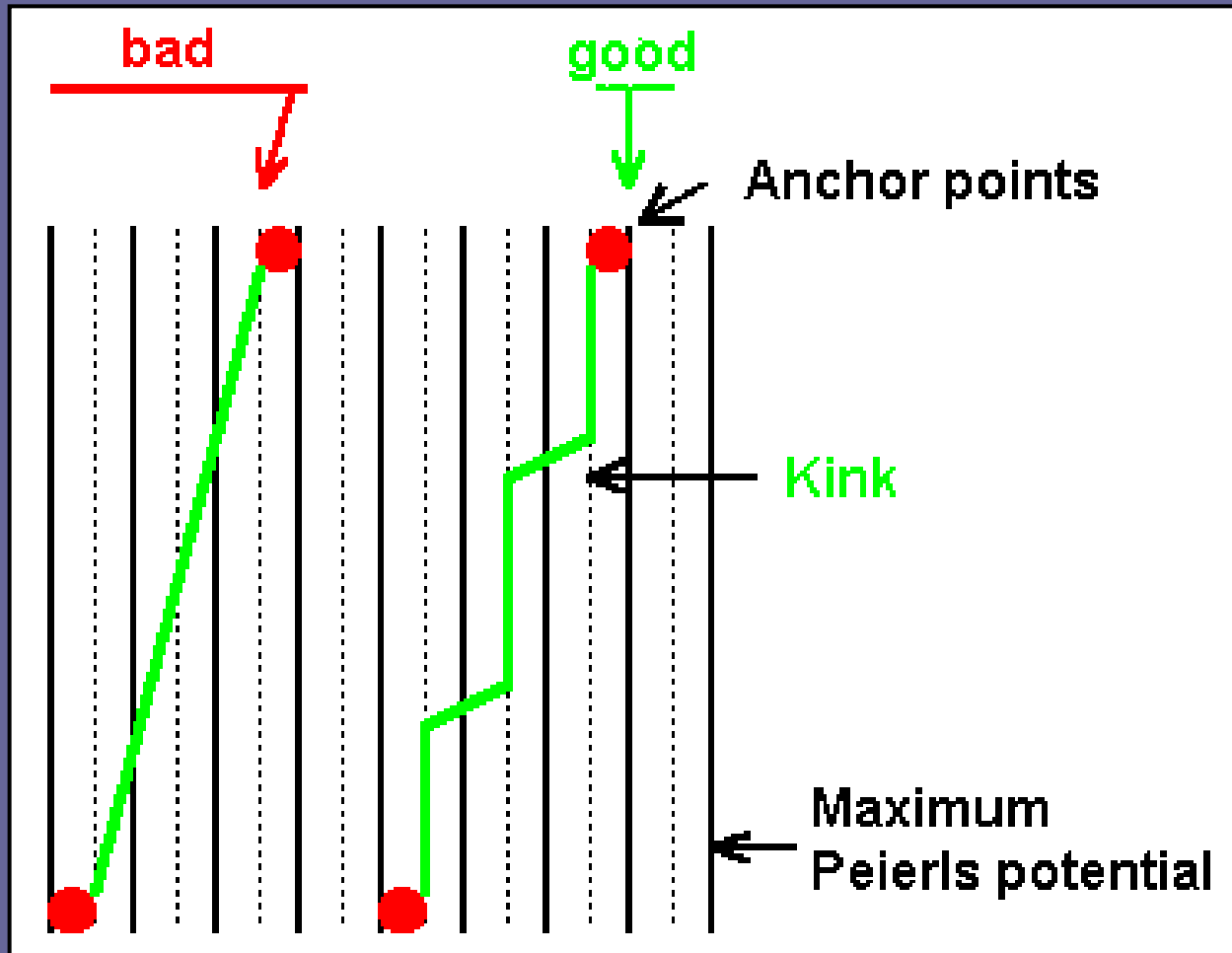


Flow Laws for Steady-State Deformation – Dislocation-Accommodated Grain Boundary Sliding

$$\dot{\varepsilon} = B_{\text{gbs}} \frac{D_{\text{gm}}}{d^1} \frac{GV_{\text{m}}}{RT} \left(\frac{\sigma}{G} \right)^3 \quad d_{\text{sgs}} < d$$

$$\dot{\varepsilon} = A_{\text{gbs}} \frac{D_{\text{gb}}}{d^2} \frac{GV_{\text{m}}}{RT} \left(\frac{\sigma}{G} \right)^2 \quad d_{\text{sgs}} > d$$

Low-Temperature Deformation – Dislocation Glide



Flow Laws for Steady-State Deformation – Low-Temperature Dislocation Creep

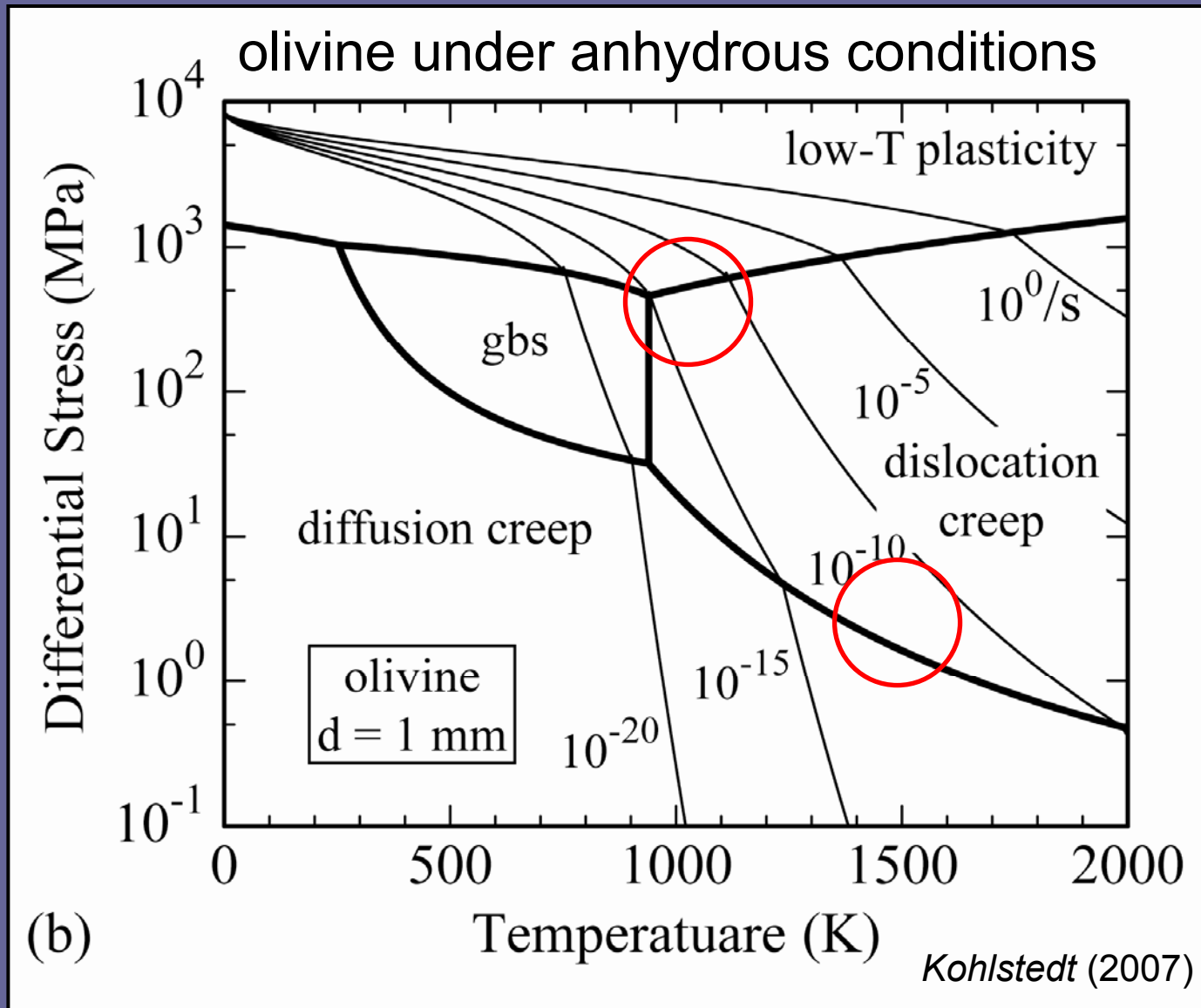
$$\bar{u} \approx u_g = c_k u_k$$

$$\Delta H_k(\sigma) = \Delta H_k^0 \left[1 - \left(\frac{\sigma}{\sigma_P} \right)^r \right]^s$$

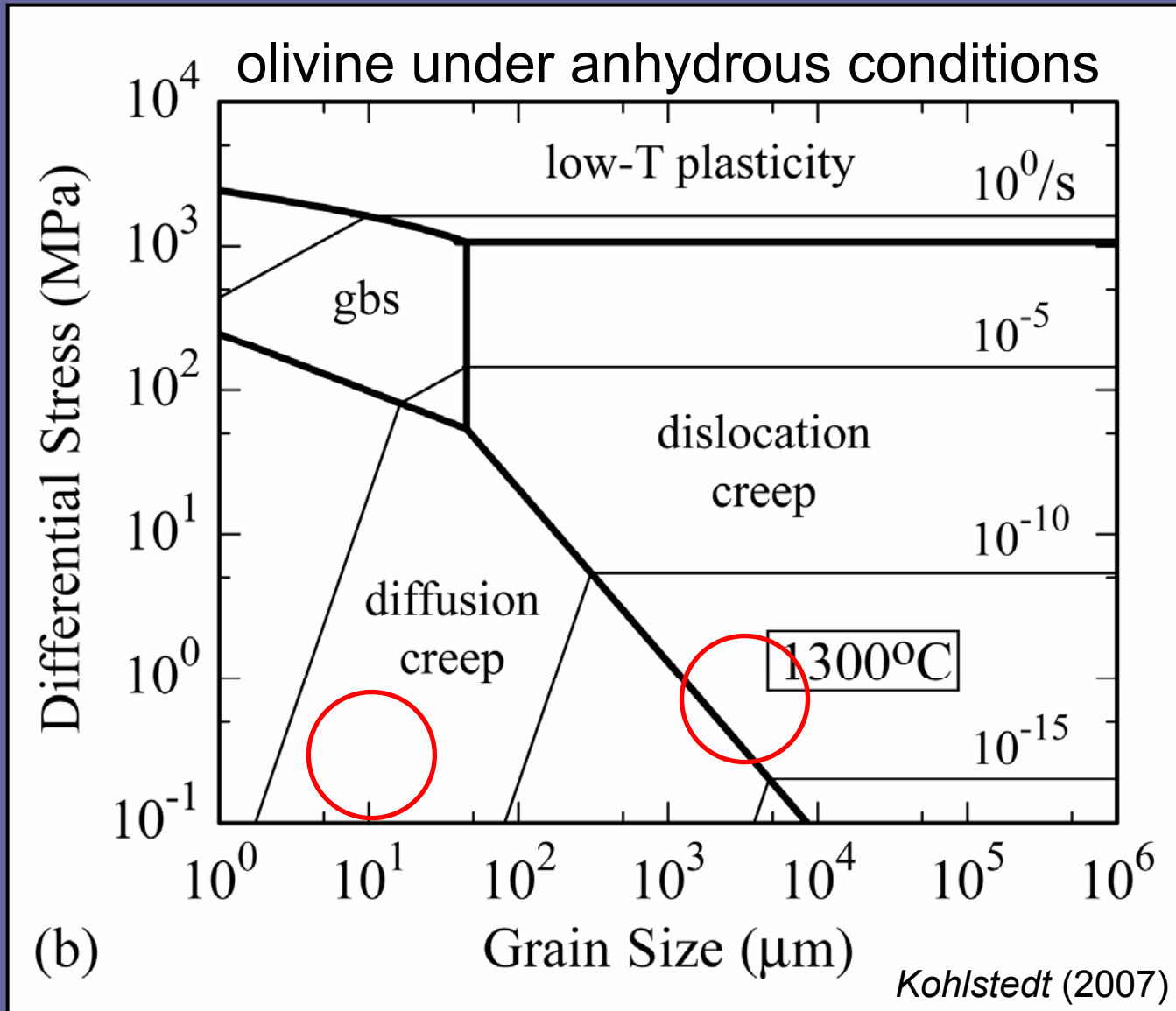
$$\dot{\epsilon} = \dot{\epsilon}_P \left(\frac{\sigma}{G} \right)^2 \exp \left\{ - \frac{\Delta H_k^0}{RT} \left[1 - \left(\frac{\sigma}{\sigma_P} \right)^r \right]^s \right\}$$

σ_P = Peierls stress, intrinsic lattice resistance to glide

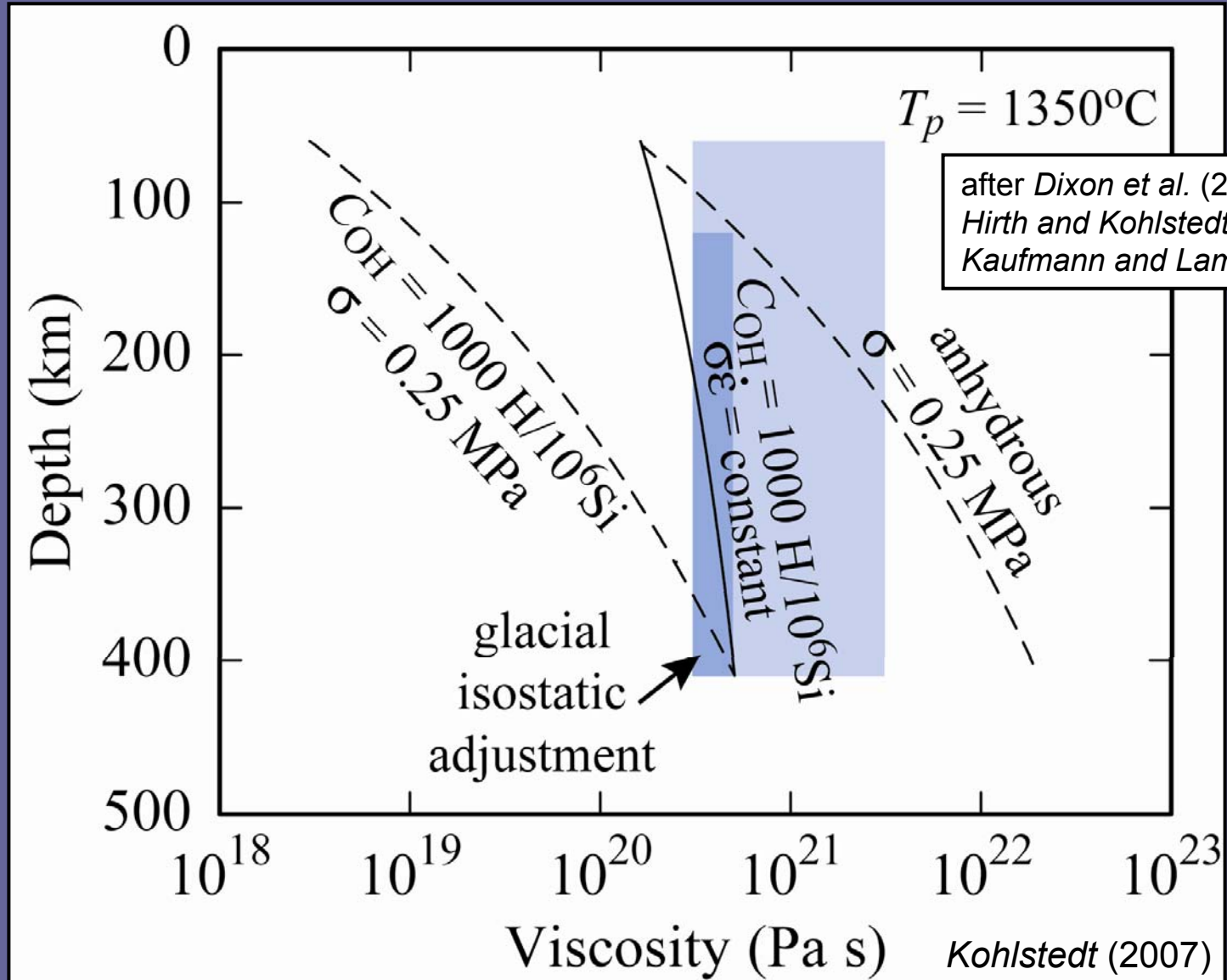
Deformation Mechanism Map – σ vs T



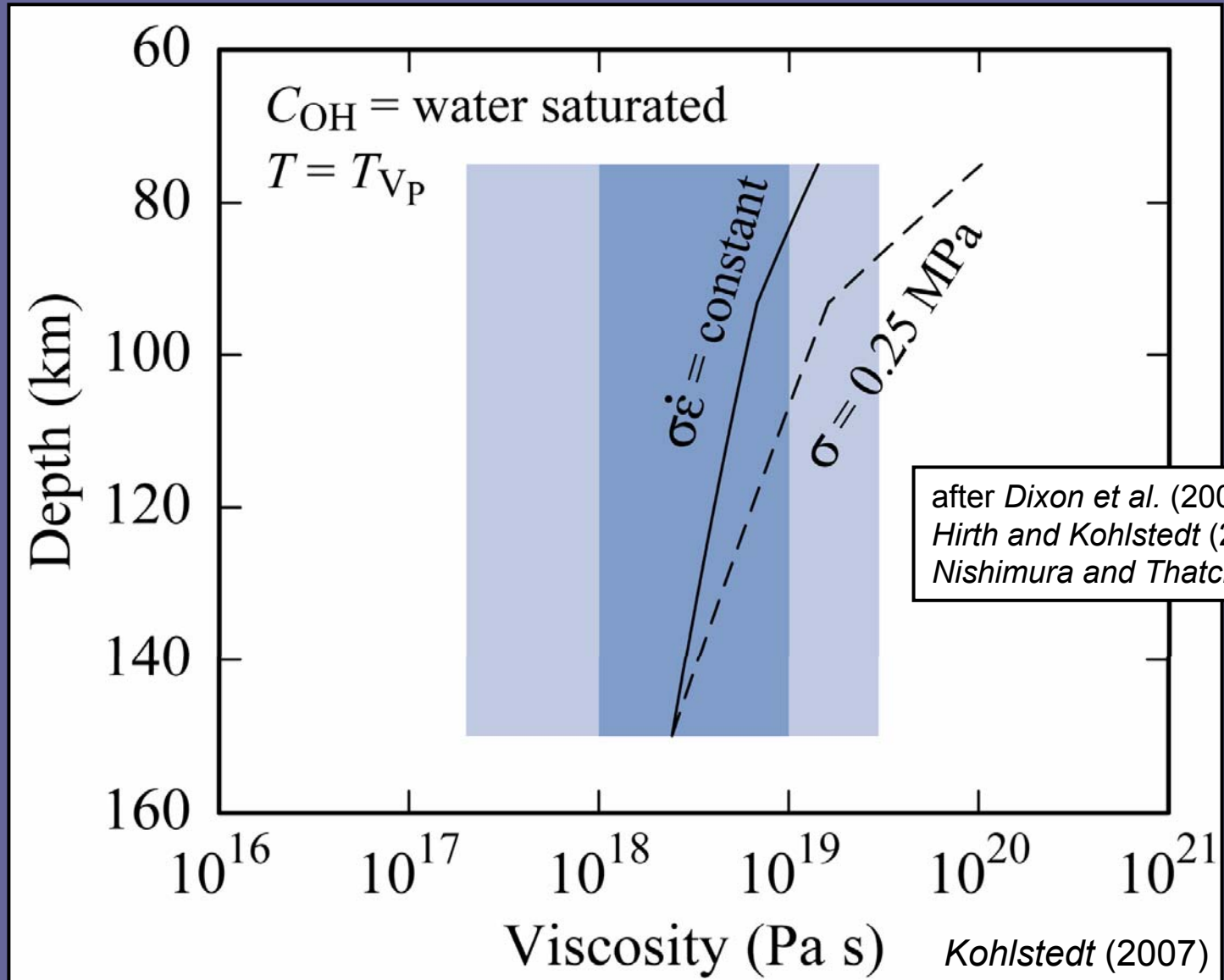
Deformation Mechanism Map – σ vs d



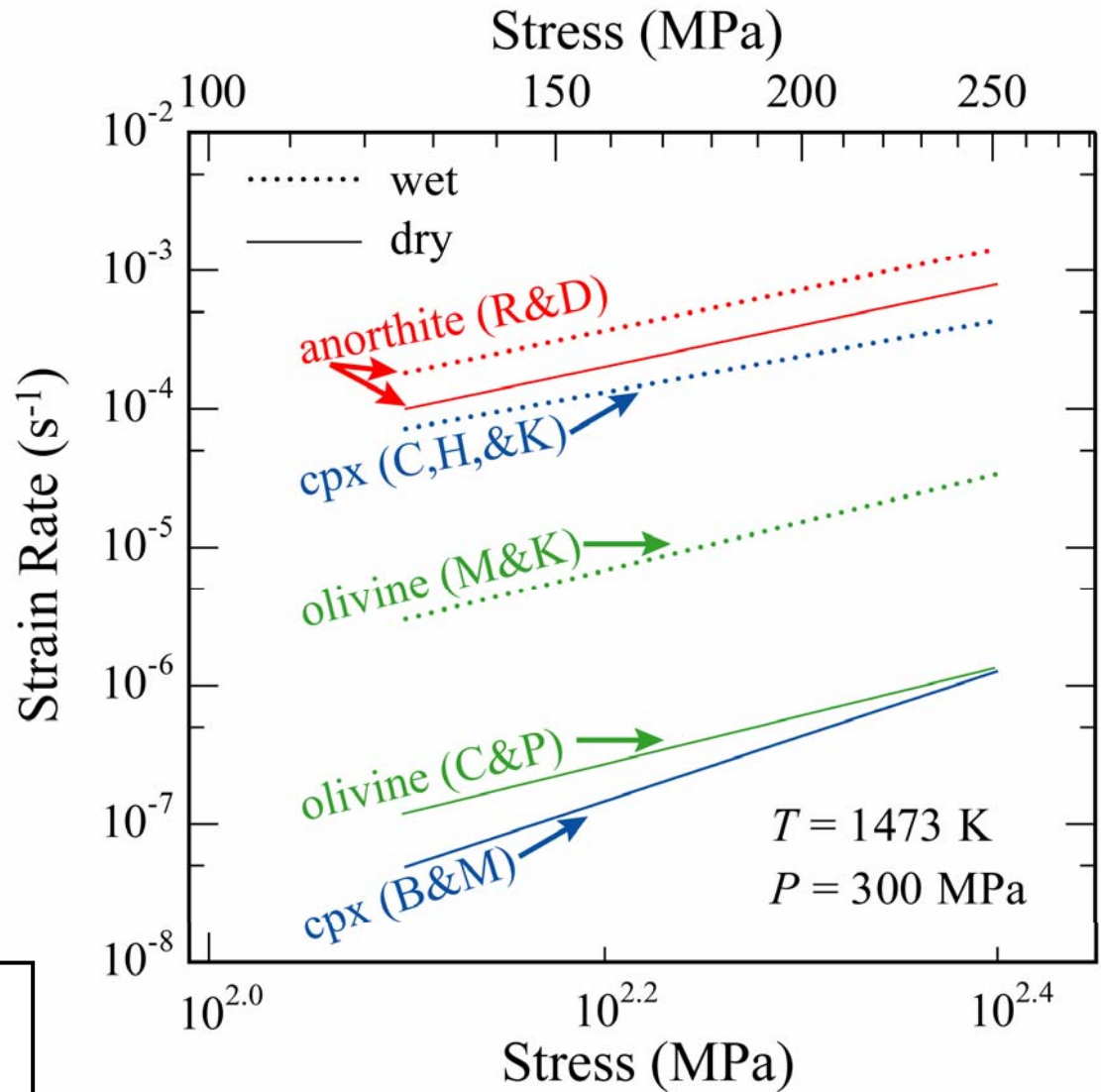
Viscosity Profiles vs Glacial Isostatic Adjustment Global Average



Viscosity Profiles vs Glacial Isostatic Adjustment Western U.S.



Comparison
of Flow
Behavior of
Several
Single-Phase
Rocks
Deformed
Under Wet
and Dry
Conditions



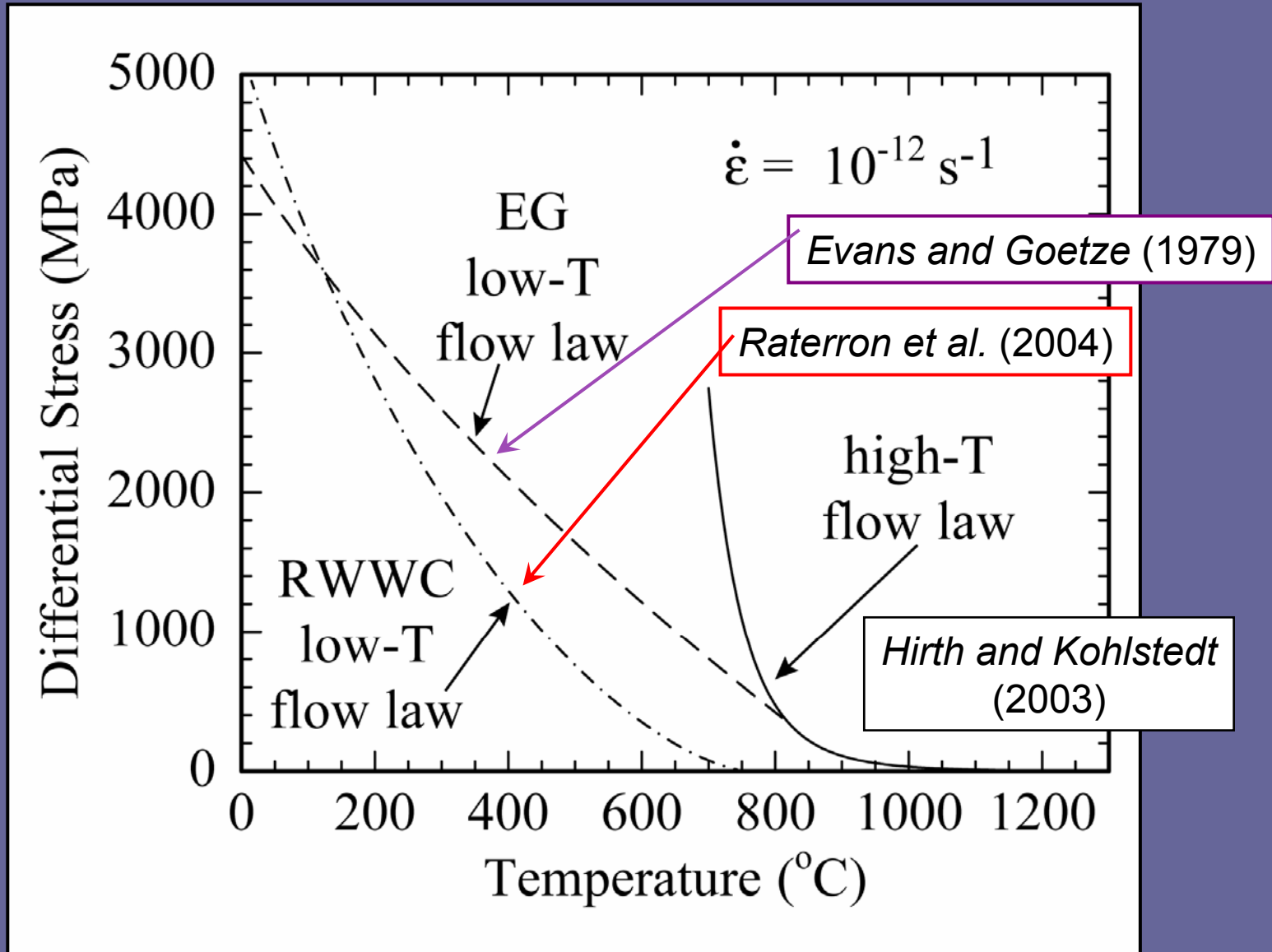
Chopra and Paterson (1981)
Rybacki and Dresen (2000)
Mei and Kohlstedt (2000)
Bystricky and Mackwell (2001)
Chen et al. (2006)

Kohlstedt (2007)

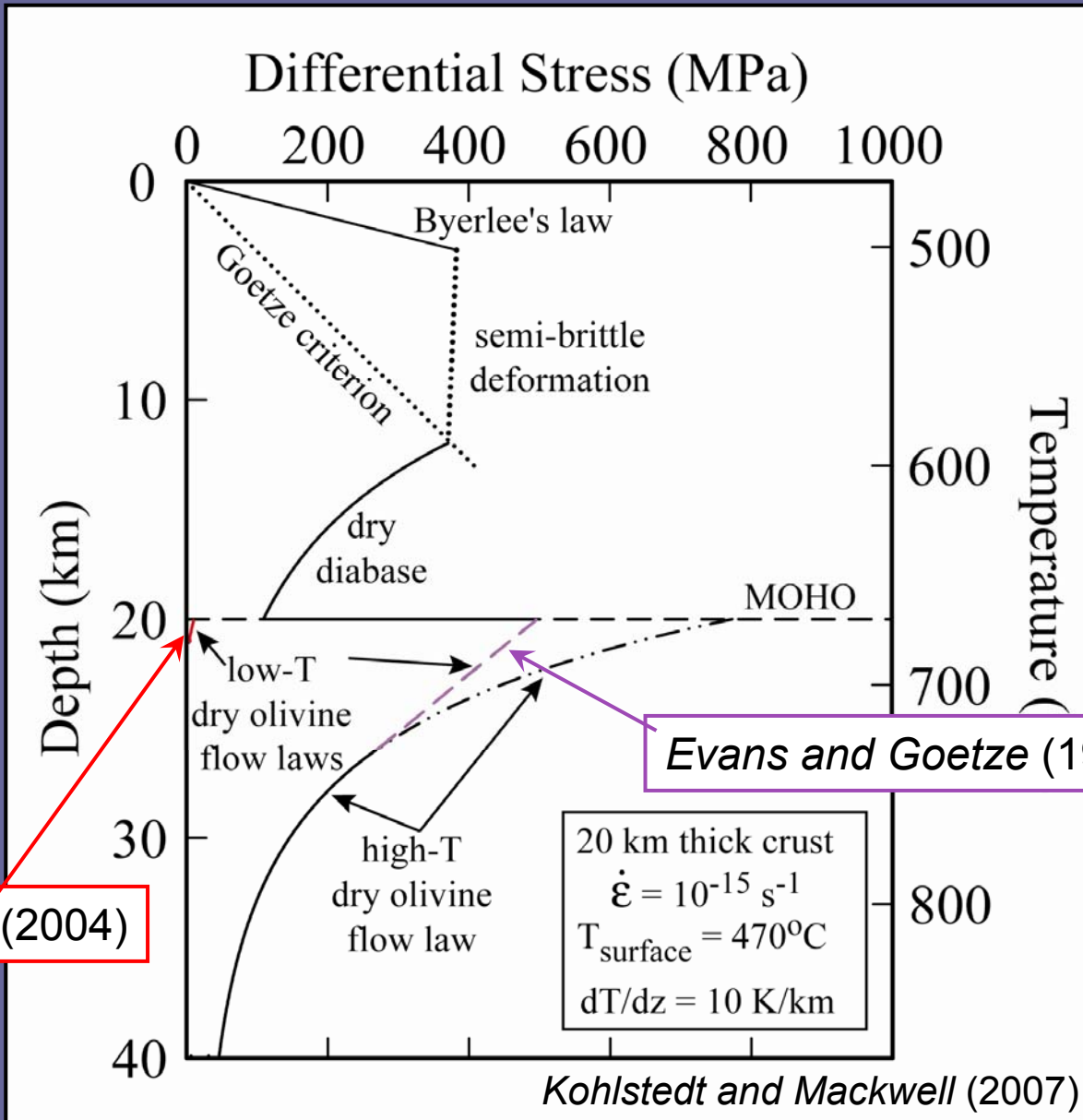
Why the Emphasis on Single-Phase Rather than Multi-Phase Rocks?

eutectic melting

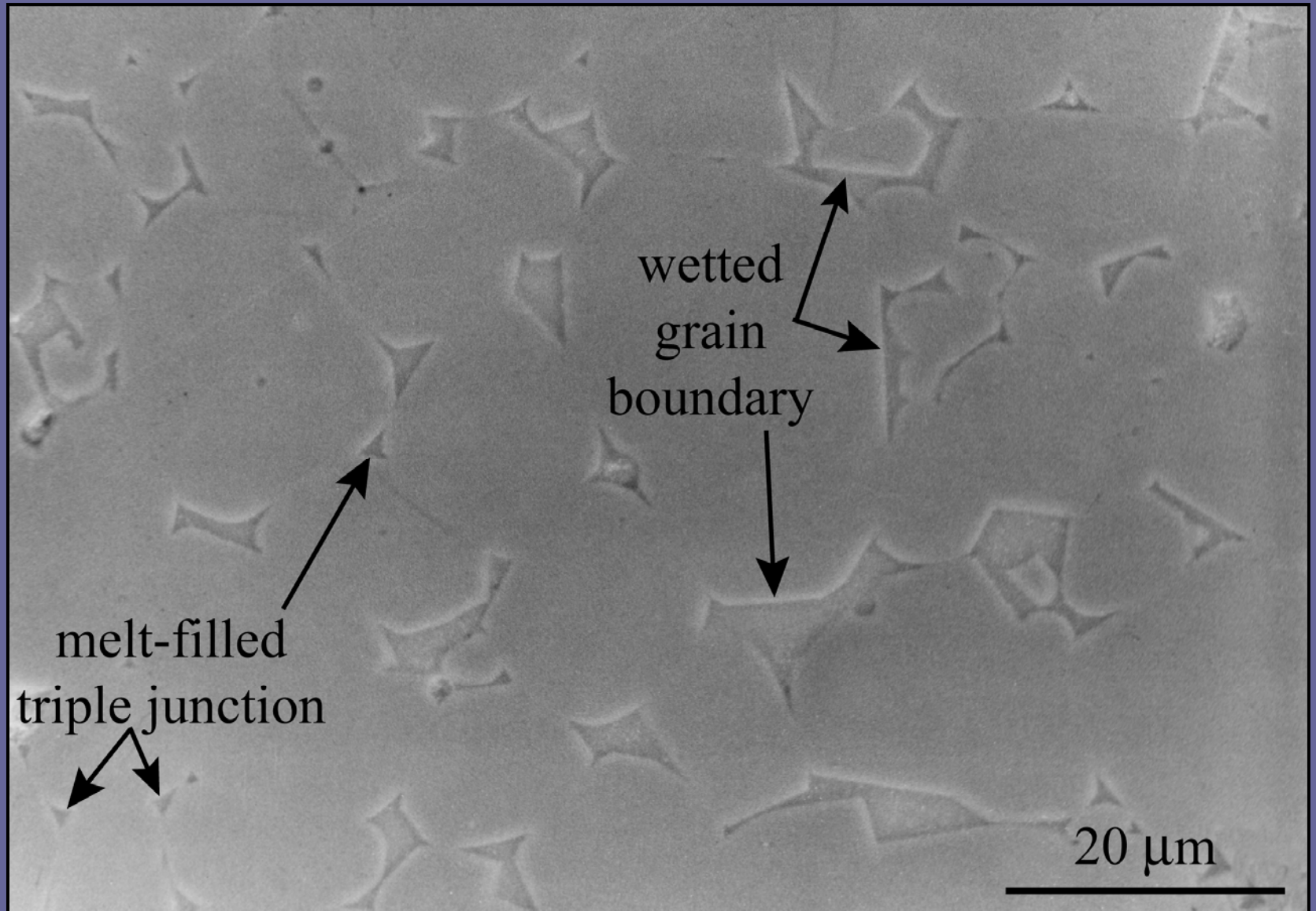
Low-Temperature Plasticity



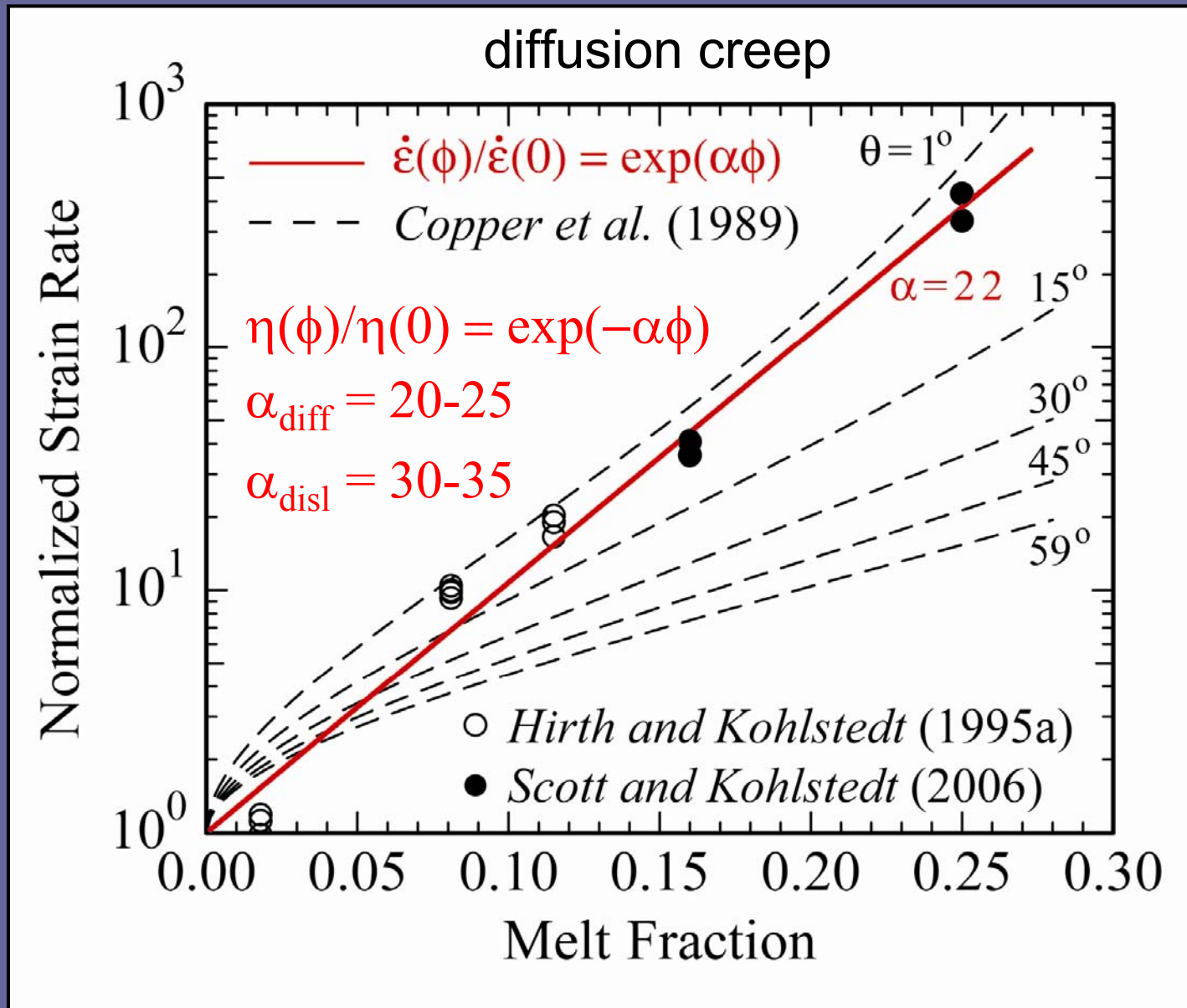
Low-Temperature Plasticity



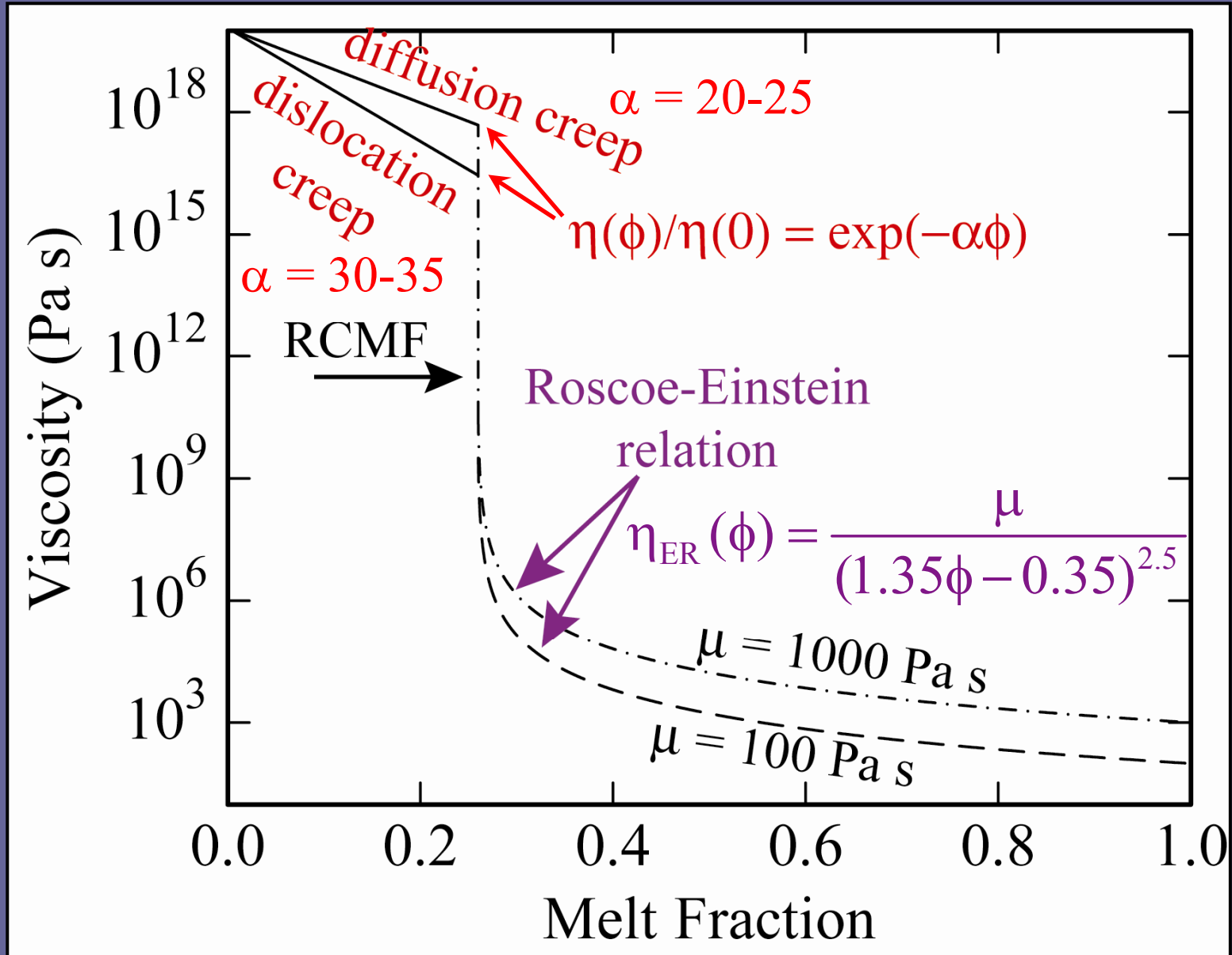
Melt Distribution in Partially Molten Rock



Flow Behavior of Partially Molten Rock



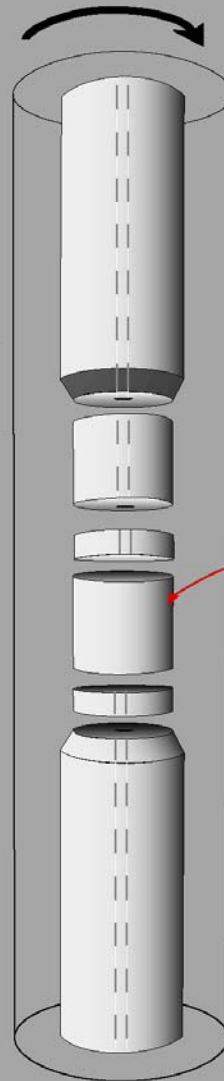
Flow Behavior of Partially Molten Rock



High-Strain Torsion Experiments

constant twist-rate or torque
applied from above

alumina
pistons
and
spacers



gas-
medium
pressure
vessel

torsion
experiments

$T = 1450 \text{ K}$
 $P = 300 \text{ MPa}$
constant
twist-rate

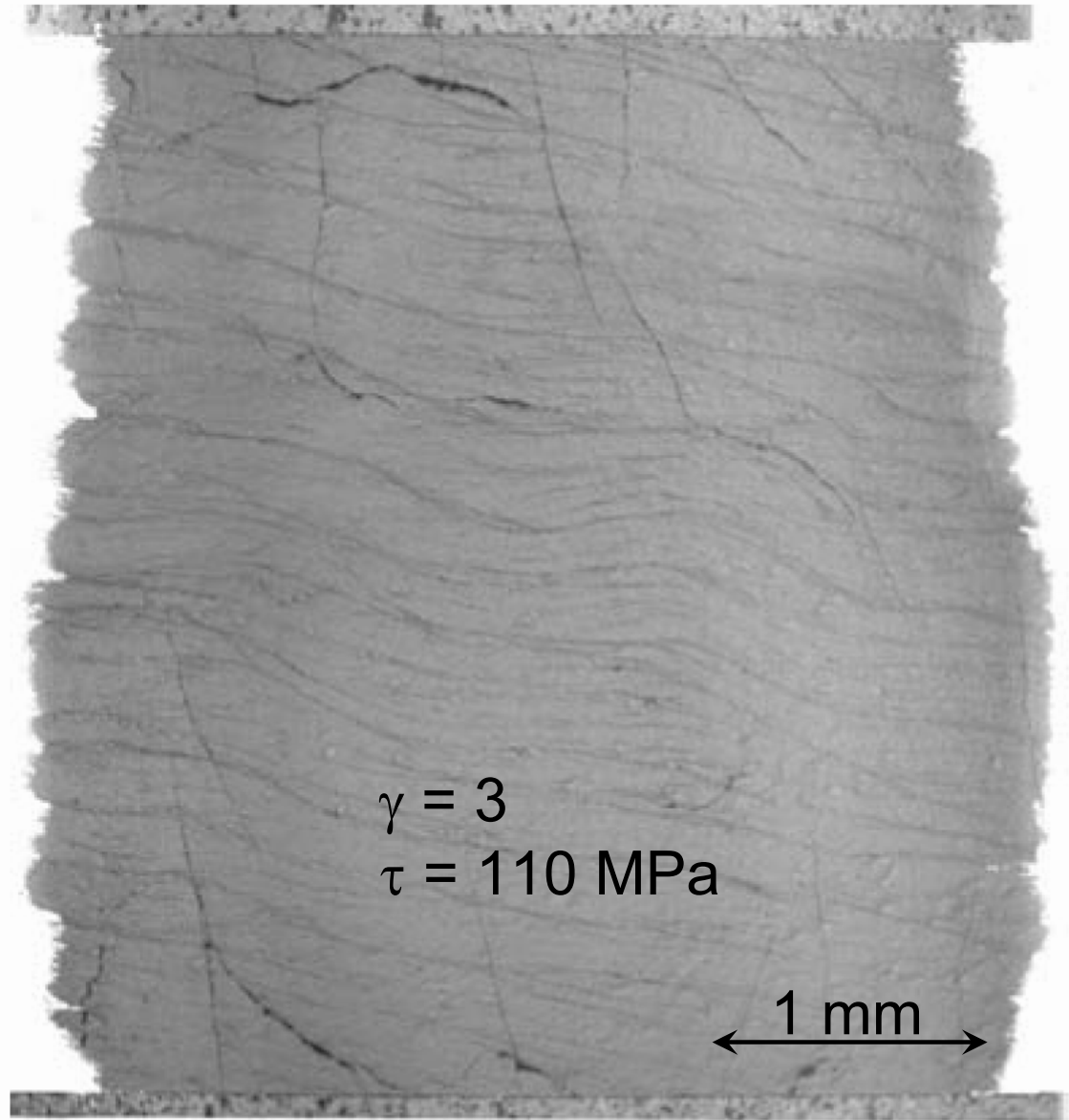
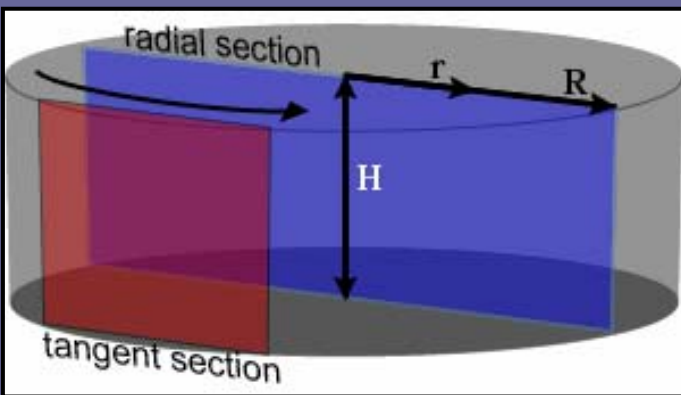
$\text{An}_{70} \pm \text{basalt}$
 $d = 3 \mu\text{m}$

sample



15 mm

Torsional Deformation of Partially Molten Rock



tangential section

Melt Segregation in Shear Partially Molten Rock

Transmitted light
plane polarized

An₇₀ $\phi = 0.07$

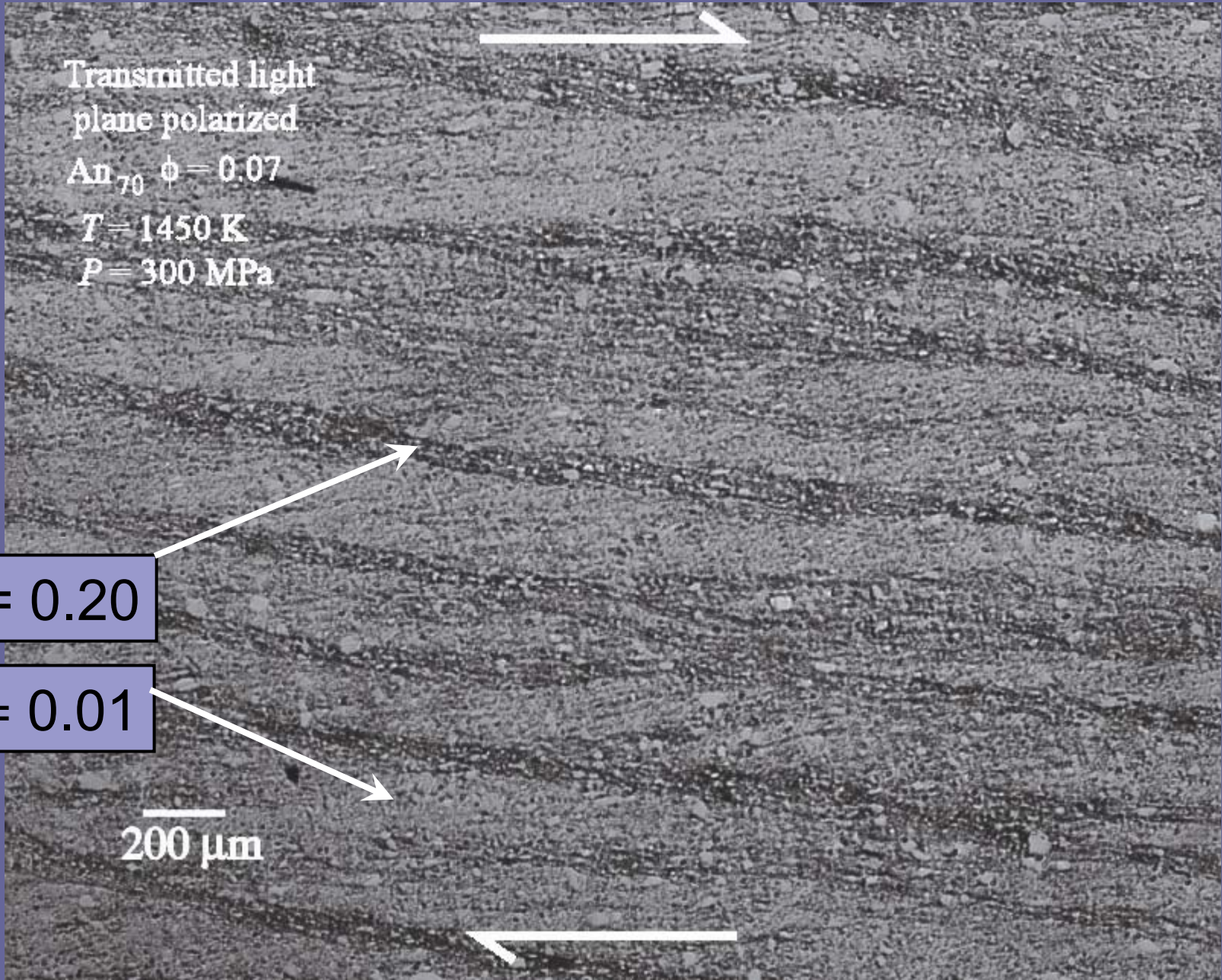
$T = 1450 \text{ K}$

$P = 300 \text{ MPa}$

$\phi = 0.20$

$\phi = 0.01$

200 μm

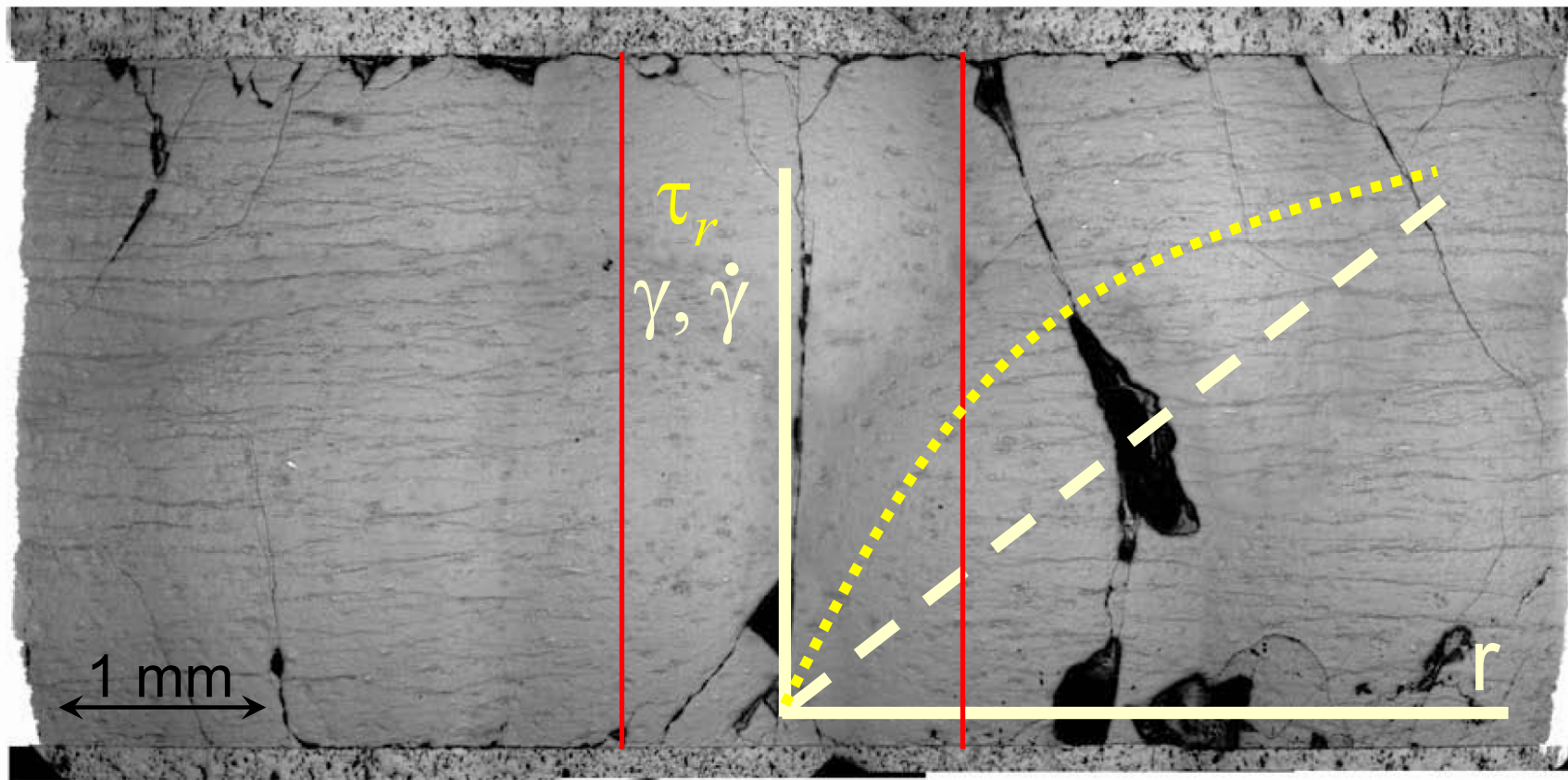
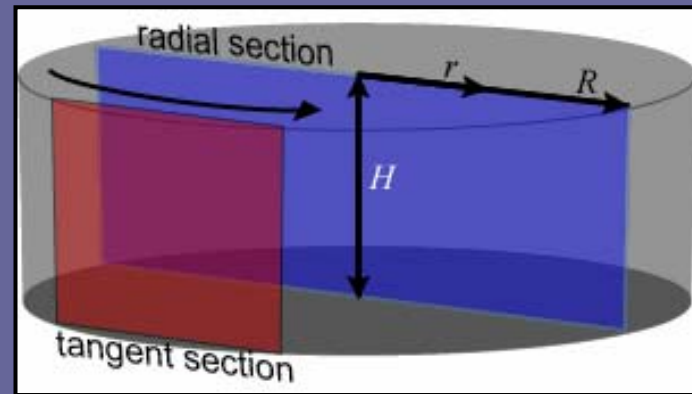


Torsional Deformation of Partially Molten Rock

$$\gamma_r = \theta r / H$$

$$\tau_r \propto r^{1/n}$$

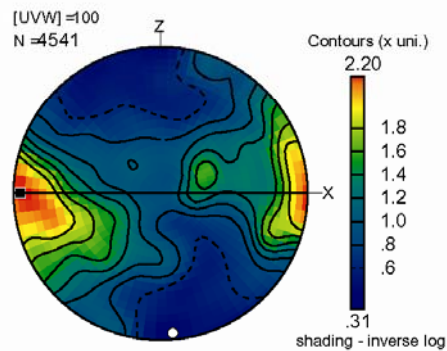
$$\dot{\gamma} \propto \tau^n$$



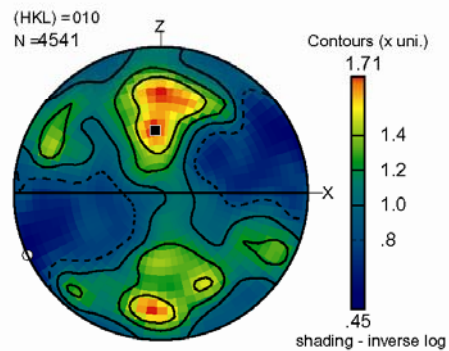
radial section

LPO from Shear Partially Molten Rock

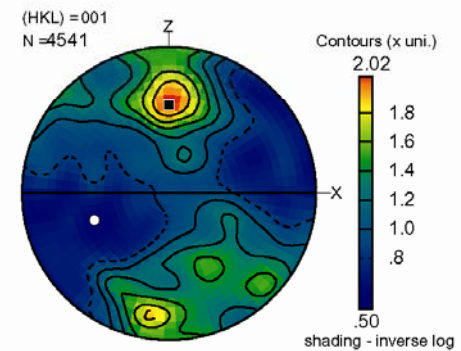
TR5, An₇₀ $\phi < 0.01$ (no MORB added), $\gamma = 4$



pfJ = 1.20

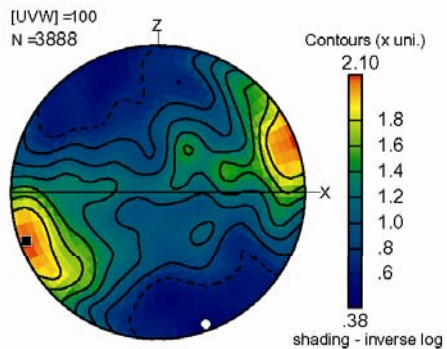


pfJ = 1.07

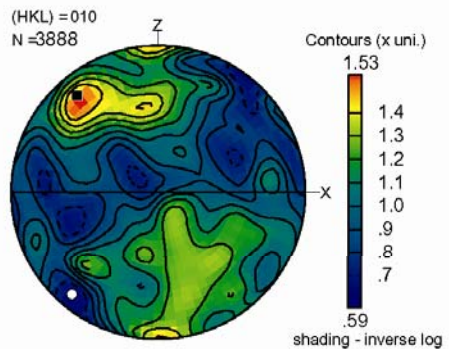


pfJ = 1.11

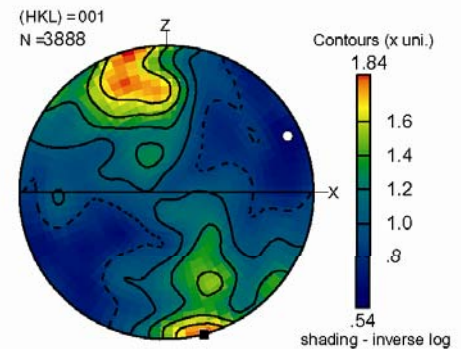
TR3, An₇₀ $\phi \approx 0.07$ (10 vol% MORB added), $\gamma = 5$



pfJ = 1.16

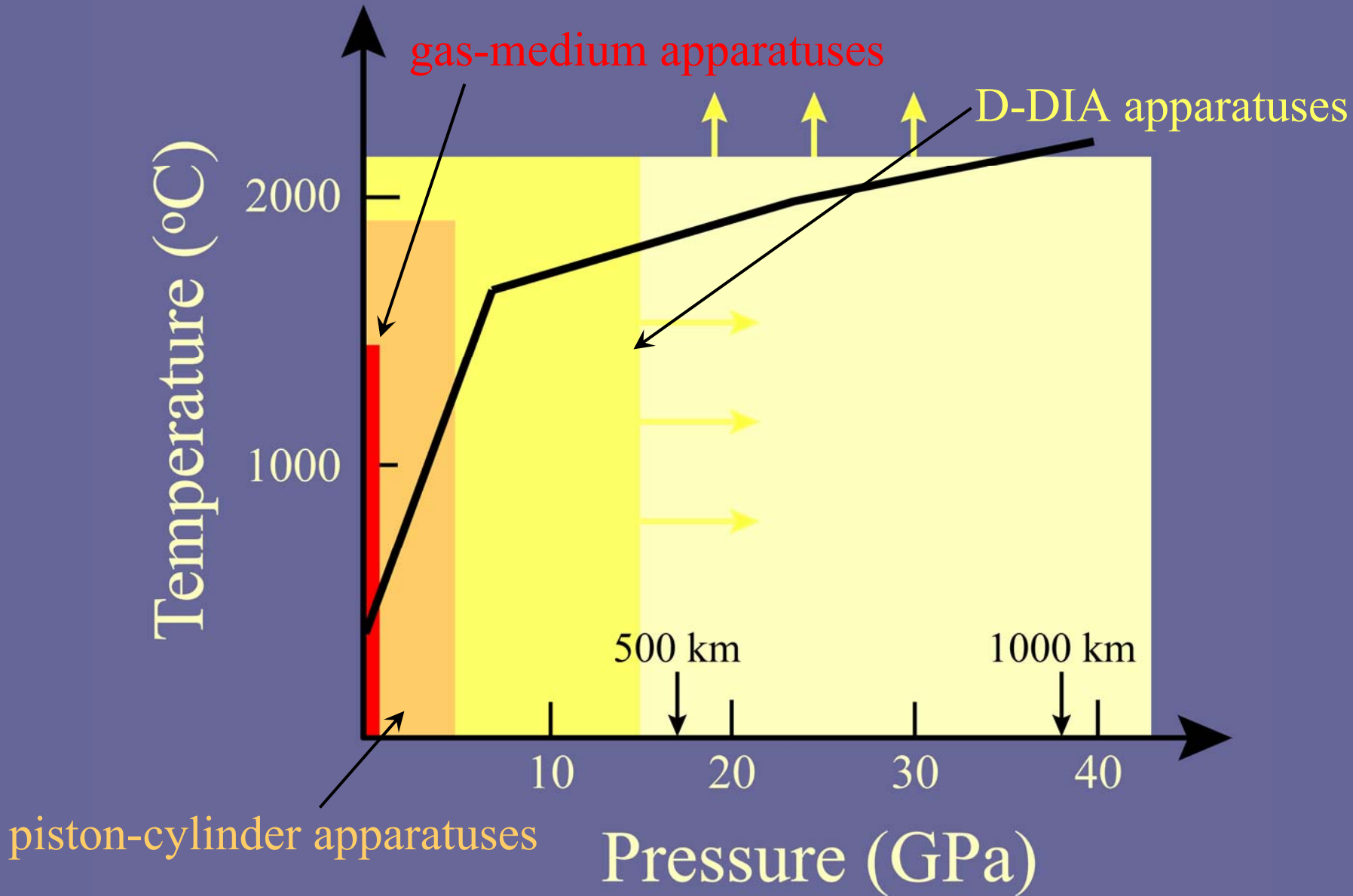


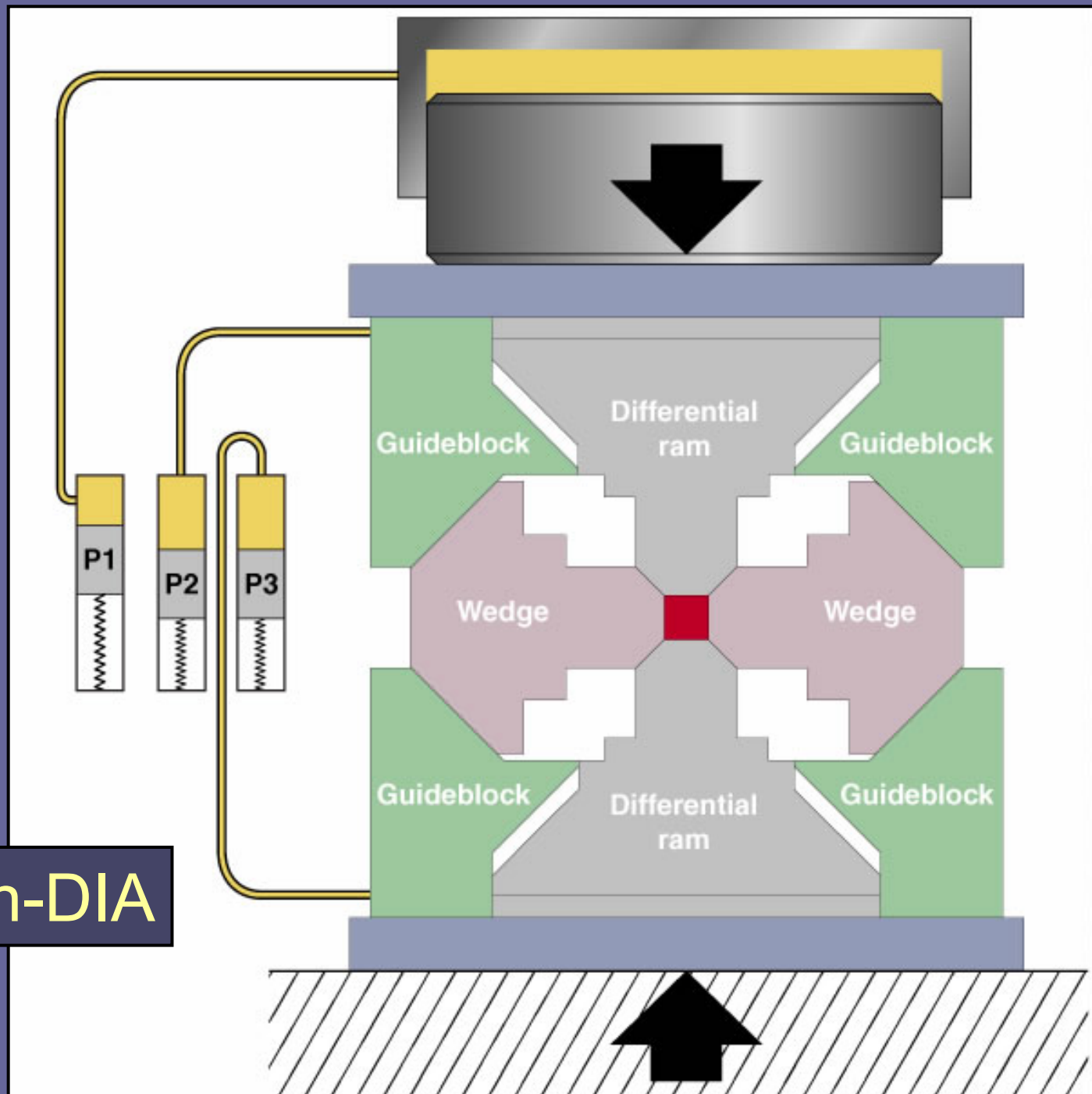
pfJ = 1.04



pfJ = 1.08

High-Pressure, High-Temperature Apparatuses

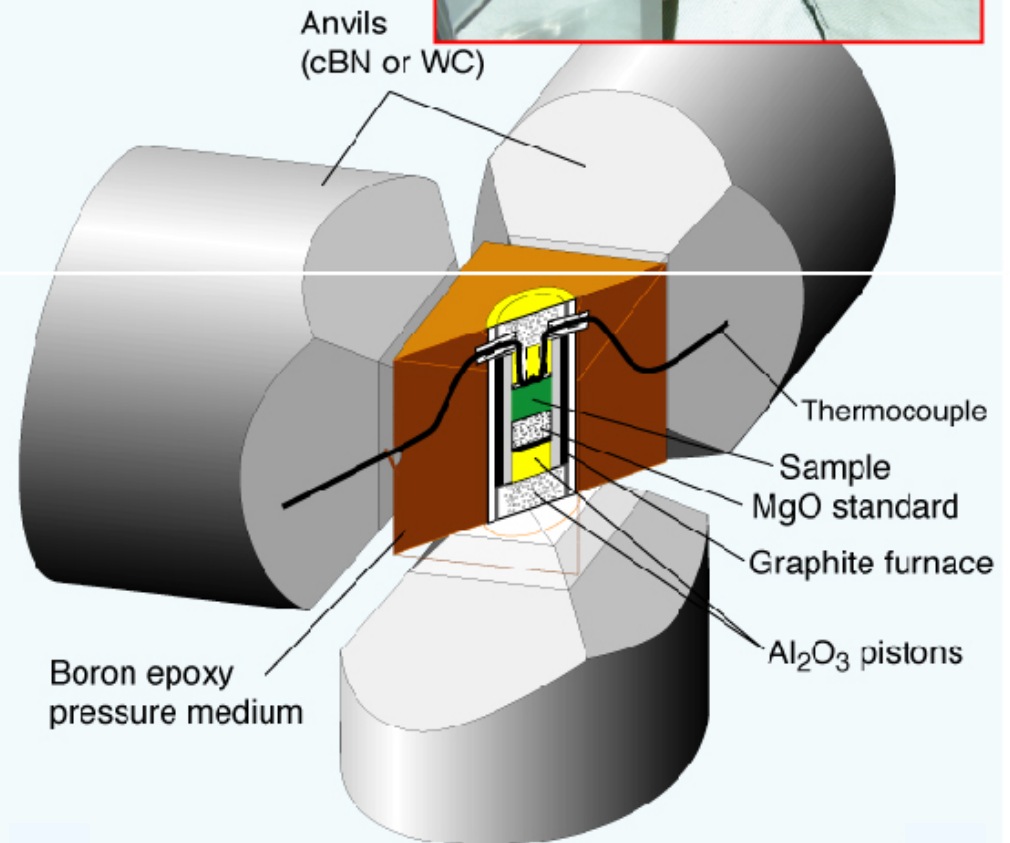




Deformation-DIA

Deformation-DIA Sample Assembly

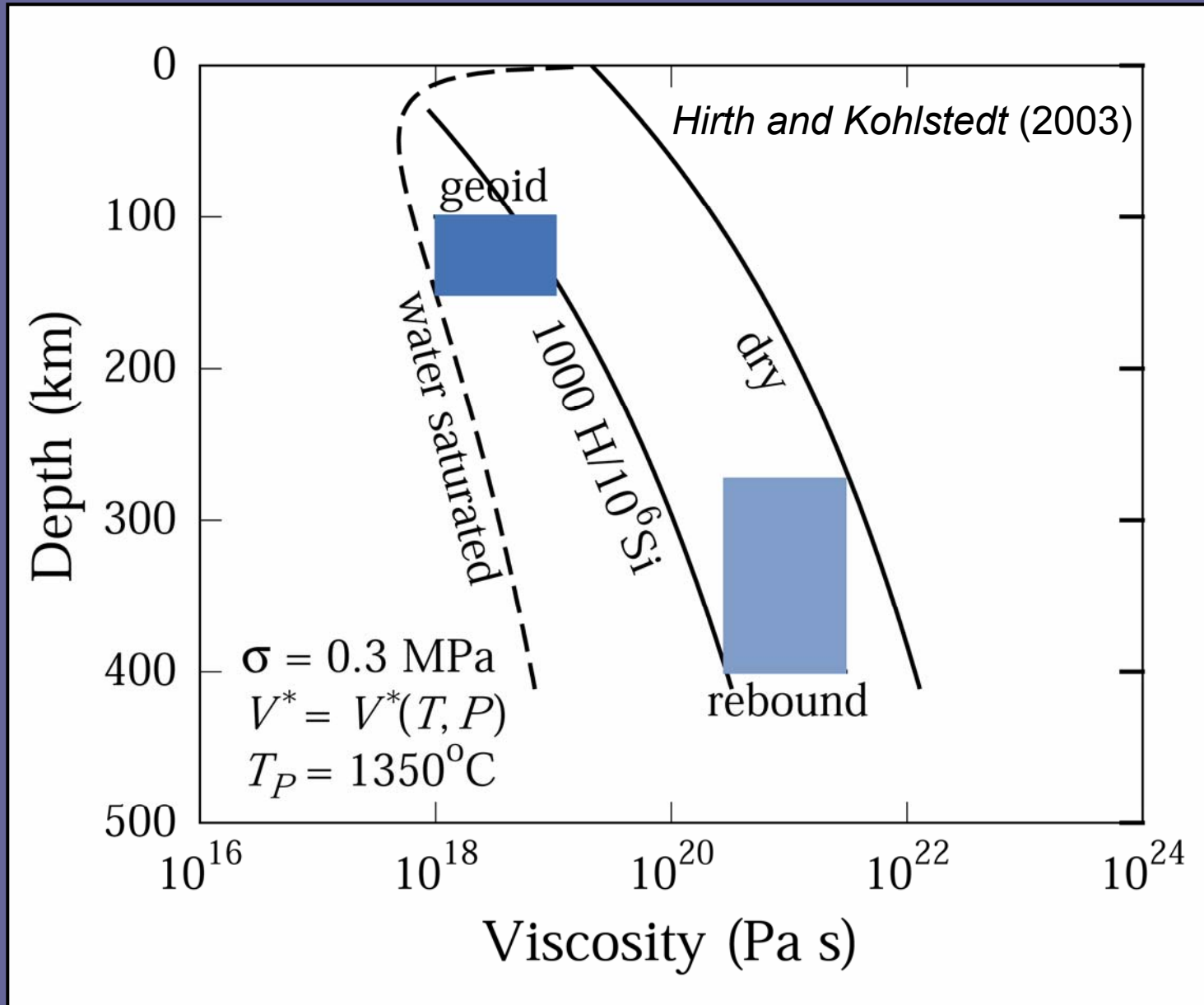
D-DIA sample
assembly



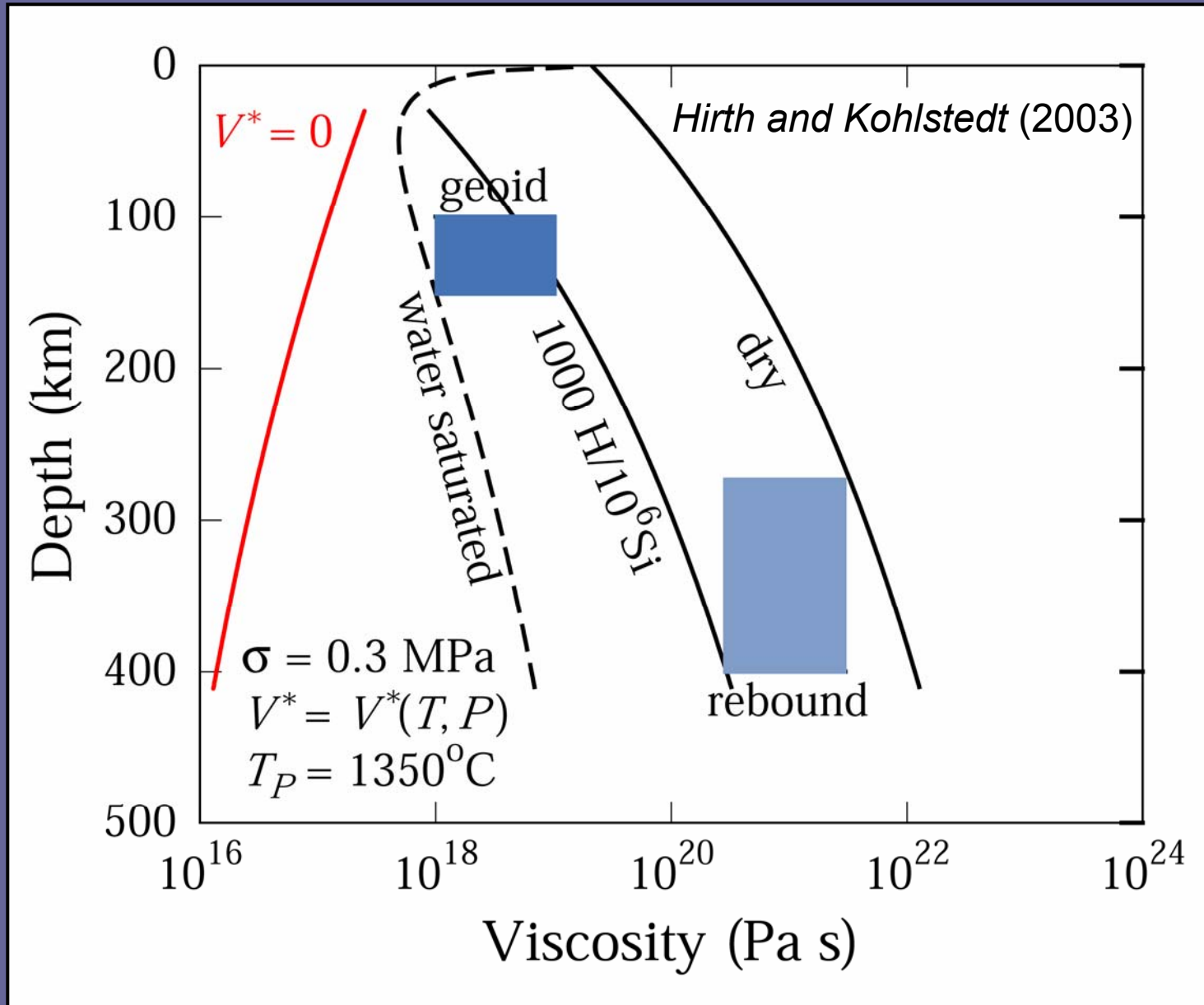
samples 1 mm in diameter

(adapted from M.Vaughan)

Viscosity Profiles vs Geoid and Rebound



Viscosity Profiles vs Geoid and Rebound



Constitutive Equations: Non-Steady State

$$\dot{\varepsilon} = \dot{\varepsilon}(\sigma, T, P, f_{\text{H}_2\text{O}}, a_{\text{ox}}, \dots, S, d, \dots, \Phi, \phi, \dots)$$

~~$$\dot{\varepsilon} = \dot{\varepsilon}(\sigma, T, P, f_{\text{H}_2\text{O}}, a_{\text{ox}}, \dots, S, d, \dots, \Phi, \phi, \dots, \varepsilon, t)$$~~

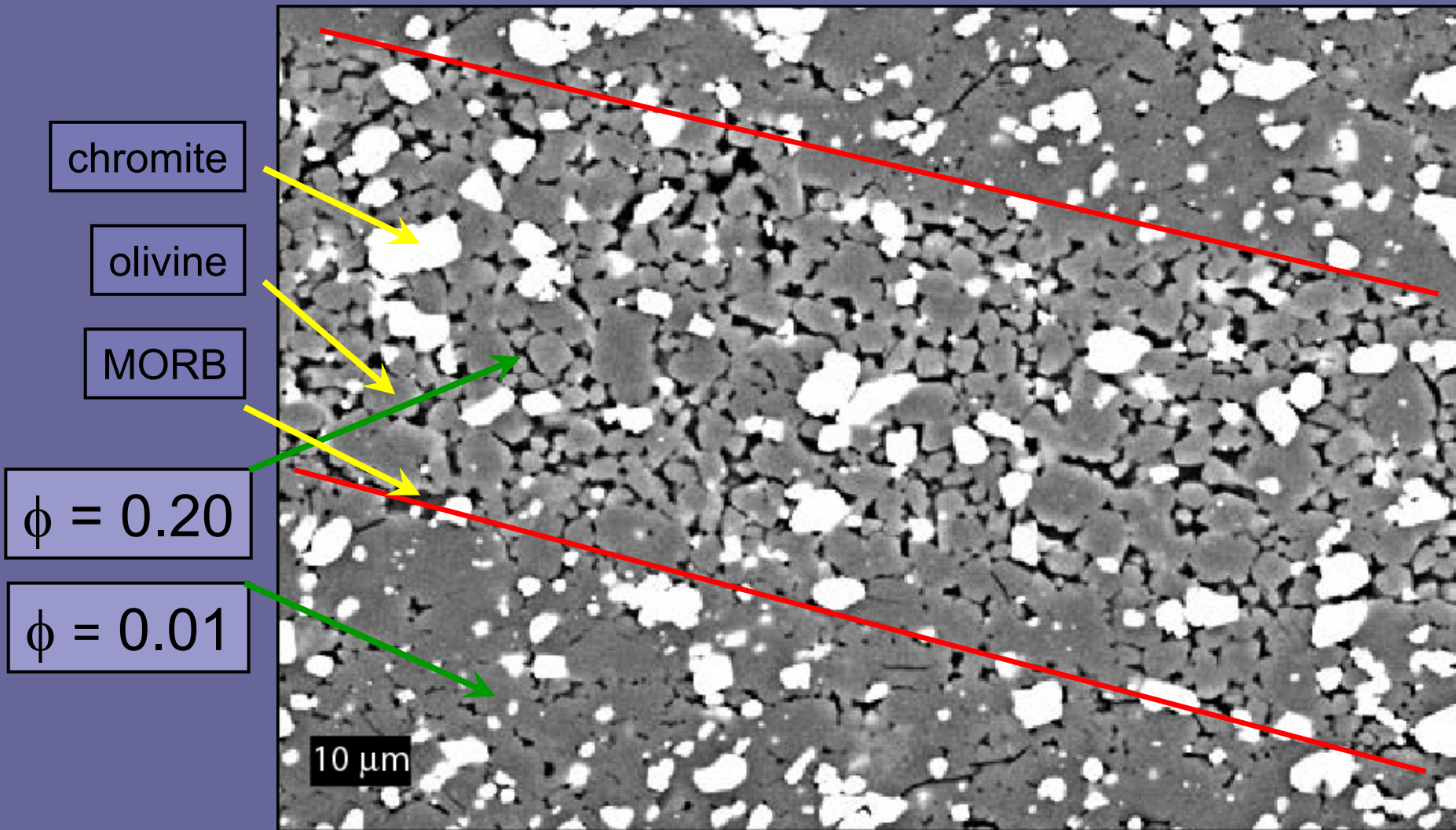
$$\dot{\varepsilon} = \dot{\varepsilon}(\sigma, T, P, f_{\text{H}_2\text{O}}, a_{\text{ox}}, \dots, S, d, \dots, \Phi, \phi, \dots, \sigma^*, d\sigma^*/dt)$$

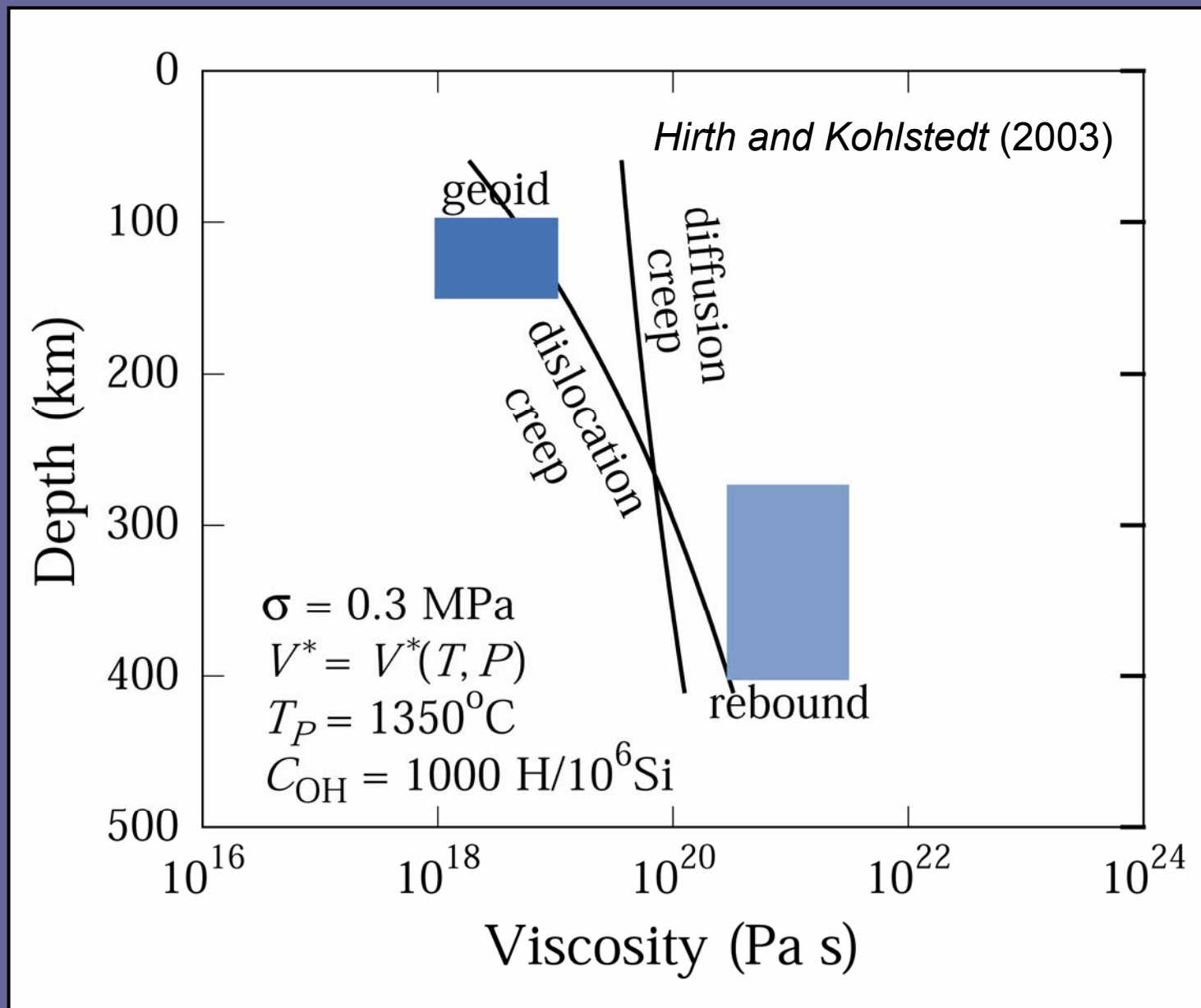
σ^* = hardness parameter, *Hart* (1970)

measure of resistance of grains to dislocation movement
possibly correlates with sub-grain size, *Stone et al.* (2004)

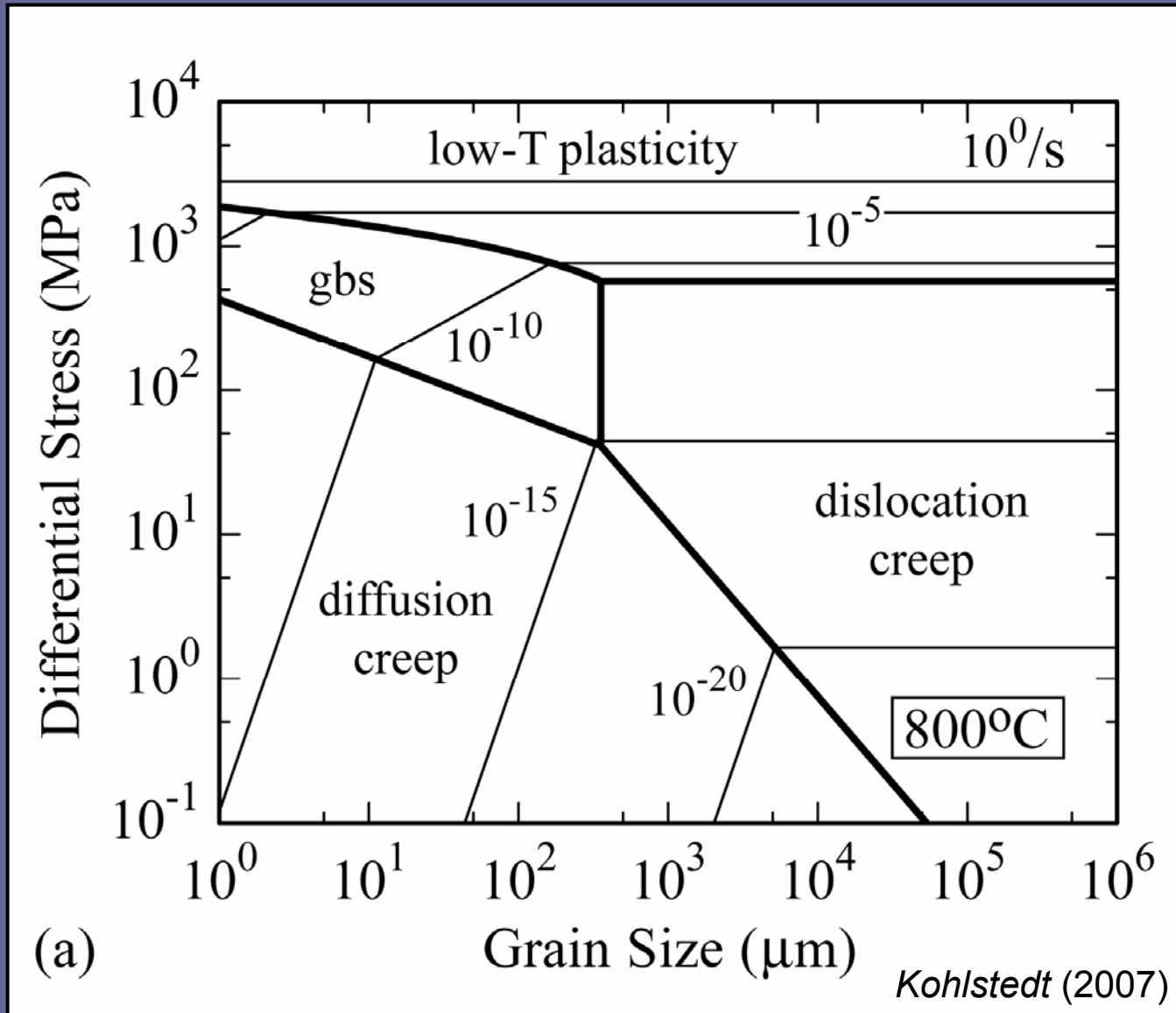
Detailed View of a Melt-Rich Band

olivine + 25 vol % chromite + 6% MORB, $\gamma = 3$, $\sigma = 100$ MPa

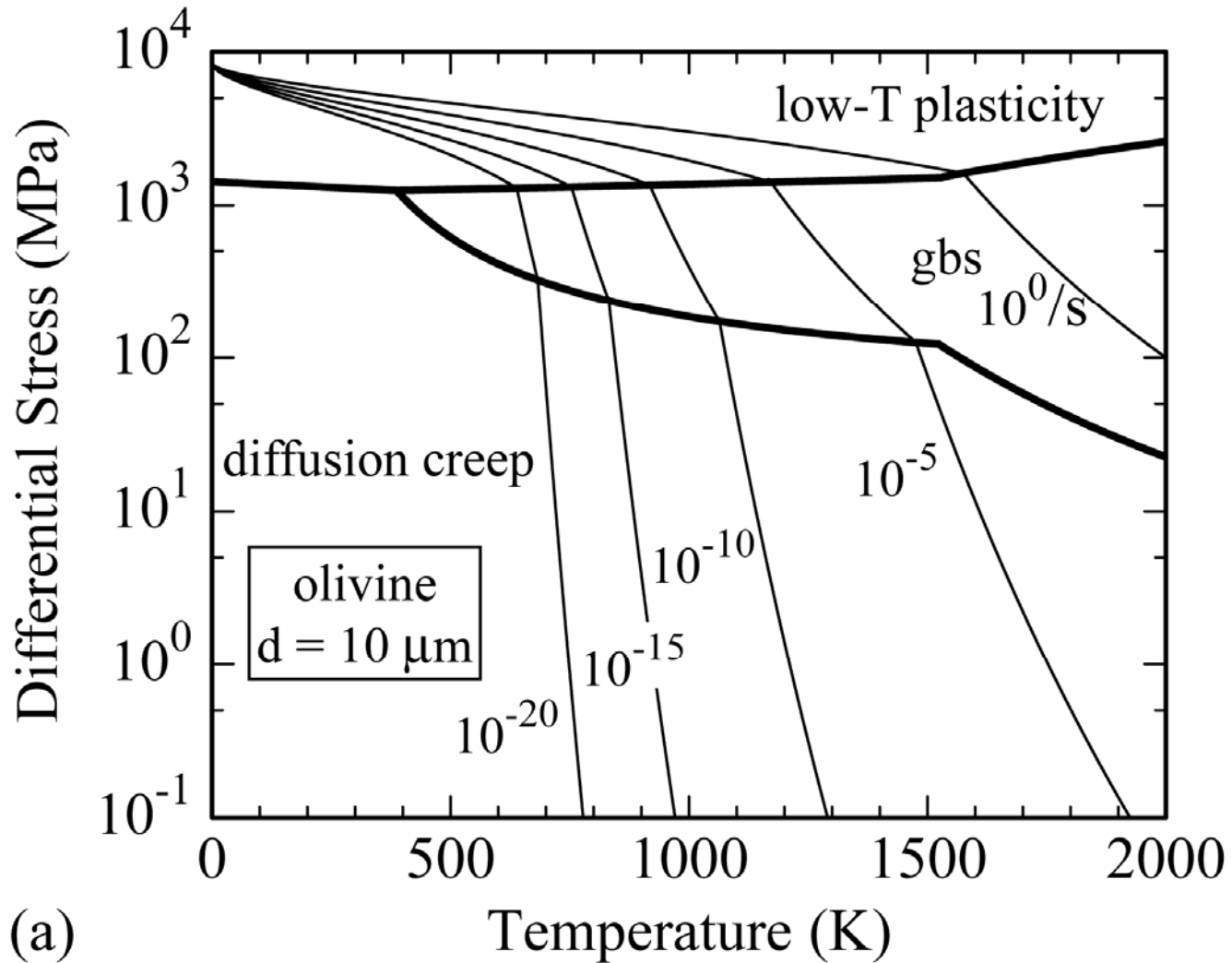




Deformation Mechanism Map – σ vs d



Deformation Mechanism Map – σ vs T



(a)

Kohlstedt (2007)