

# Three-dimensional numerical simulations of thermo-chemical multiphase convection in Earth's mantle

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# Outline

1. Introduction: Previous studies on numerical models of thermo-chemical mantle convection.
2. Numerical models: Getting in sight to methods briefly.
3. Examples: Single-component phase changes system and multi-component phase change system.
4. Examples: Single-component phase changes with the post-perovskite phase change.
5. Summary.

## Previous 3D studies

- Flow-scale and formation of thermo-chemical plumes in a rectangular box [Tackley, 1998; McNamara and Zhong, 2004 EPSL] and spherical shells [McNamara and Zhong, 2004 JGR].
- Heat flow scaling in a spherical shell [Oldham and Davies, 2004].
- All those studies have been done by a layered initial condition for the compositional field.
  - No melt-induced material differentiation with a uniform initial condition for the compositional field.

## **Previous studies on the melt-induced material differentiation in mantle convection models**

- Formation of compositional stratification in the uppermost lower mantle due to the density crossover between olivine phase system and pyroxene phase system [Tackley and Xie, 2003; Ogawa, 2004].
- Those studies have been done in only two-dimensional cases.
- Investigating three-dimensional effects is essential for understanding thermo-chemical structure in Earth's mantle.

## Post-perovskite: Newly discovered phase change near the core-mantle boundary

- Discovered by high pressure mineral physics [Murakami et al., 2004; Oganov and Ono, 2004; Tsuchiya et al., 2004].
- Large positive Clapyeron slope (8 to 10MPa/K)
- Destabilize thermo-chemical boundary layer [Nakagawa and Tackley, 2005].
- Still worth to investigate a stability of thermo-chemical boundary layer above the CMB in a three-dimensional geometry.

## In this study

- Checking the timing information on various types of CPU.
- Three-dimensional numerical simulations of thermo-chemical multiphase convection to investigate thermo-chemical structure in the uppermost lower mantle and mass transport across 660km depth.
- Adding newly discovered phase change near the core-mantle boundary and see what happens to chemically-dense layering above the CMB.

# Numerical Method

- STAG3D: Finite Volume multigrid code for thermo-chemical multiphase convection in a compressible anelastic fluid as a modeled mantle.
- 128x128x64 resolution with 16 tracers in each cell to track the chemical composition and degree of melting.
- Melting tracers are instantaneously moved to the surface when it is erupted.
- Grid refinements are used in the surface, CMB and 660km to resolve thermal boundary layer.
- Simple core-cooling condition is applied to bottom thermal boundary condition.

# CPU's that is using in this study

- Timing information: Using Intel P4 3.2GHz x 16 CPUs. (Integrating over 4.5Gyrs)
  - Approximately 5 days with using g77 compiler.
  - Approximately 4 days with using Intel Fortran Compiler.
  - Seems to be reasonable time to study three-dimensional cases systematically.
- Challenging the optimization to the Earth Simulator
  - Fails to vectorize STAG3D because special multigrid techniques that is required to simulate large viscosity variation due to temperature and yield stress has not been suitable for the vectorization.
  - Timing is the similar to Pentium 4.

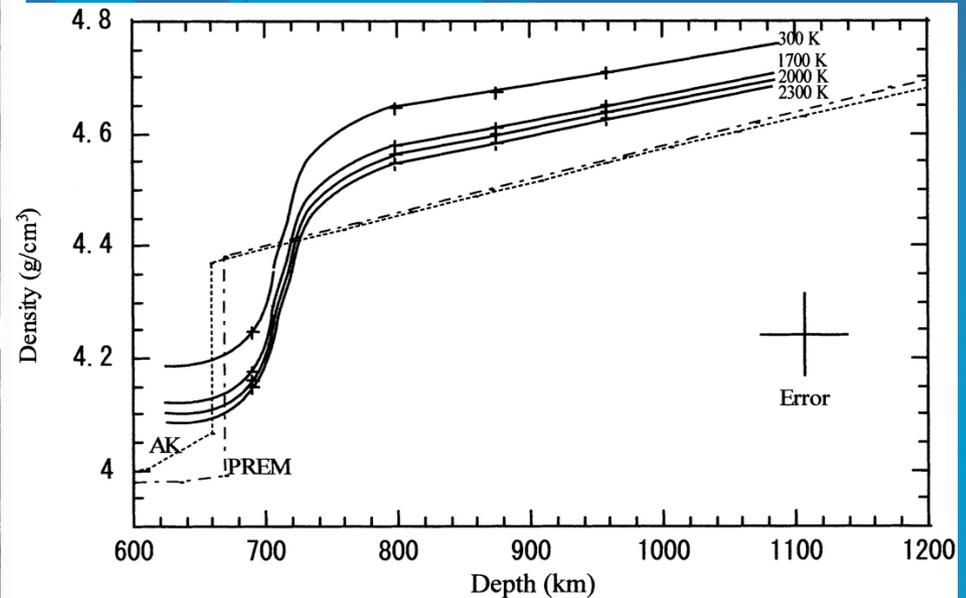
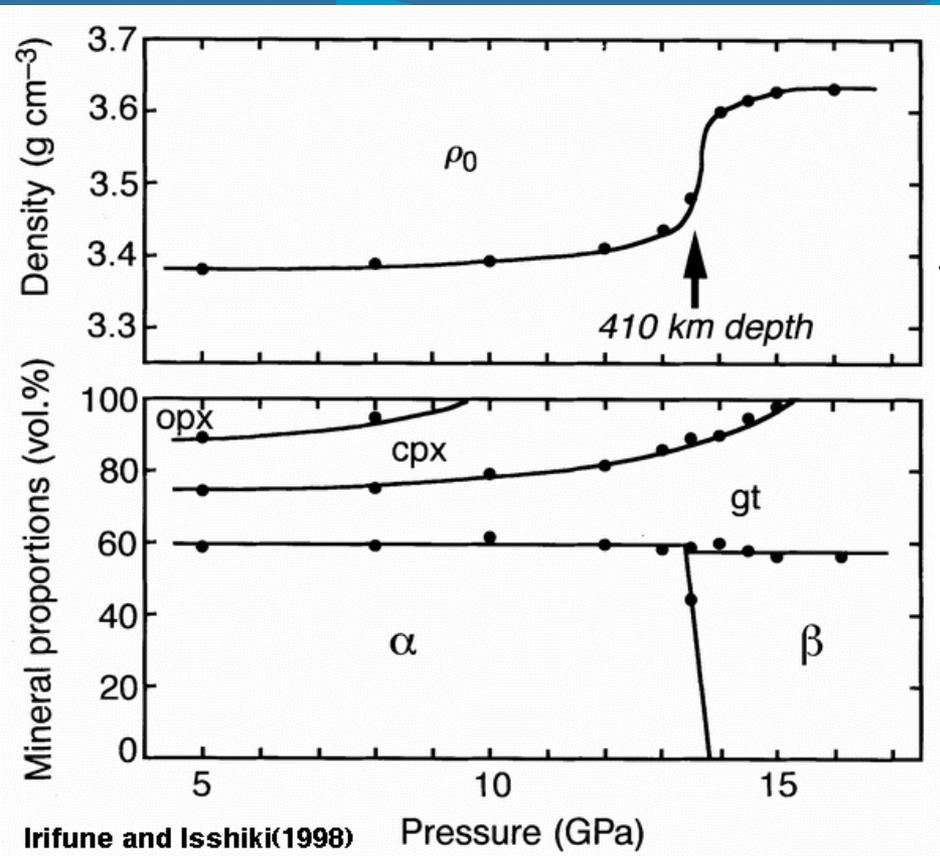
## Important material properties

- Viscosity: Temperature- (an order of six), depth (an order of two) and yield stress-dependent viscosity under Bingham fluid approximation [Trampert and Hansen, 1998; Stein et al., 2004].



- Phase changes: olivine-spinel-perovskite-post perovskite and pyroxene-garnet-perovskite-post perovskite systems are assumed.

# Phase diagram and density profiles from experimental studies

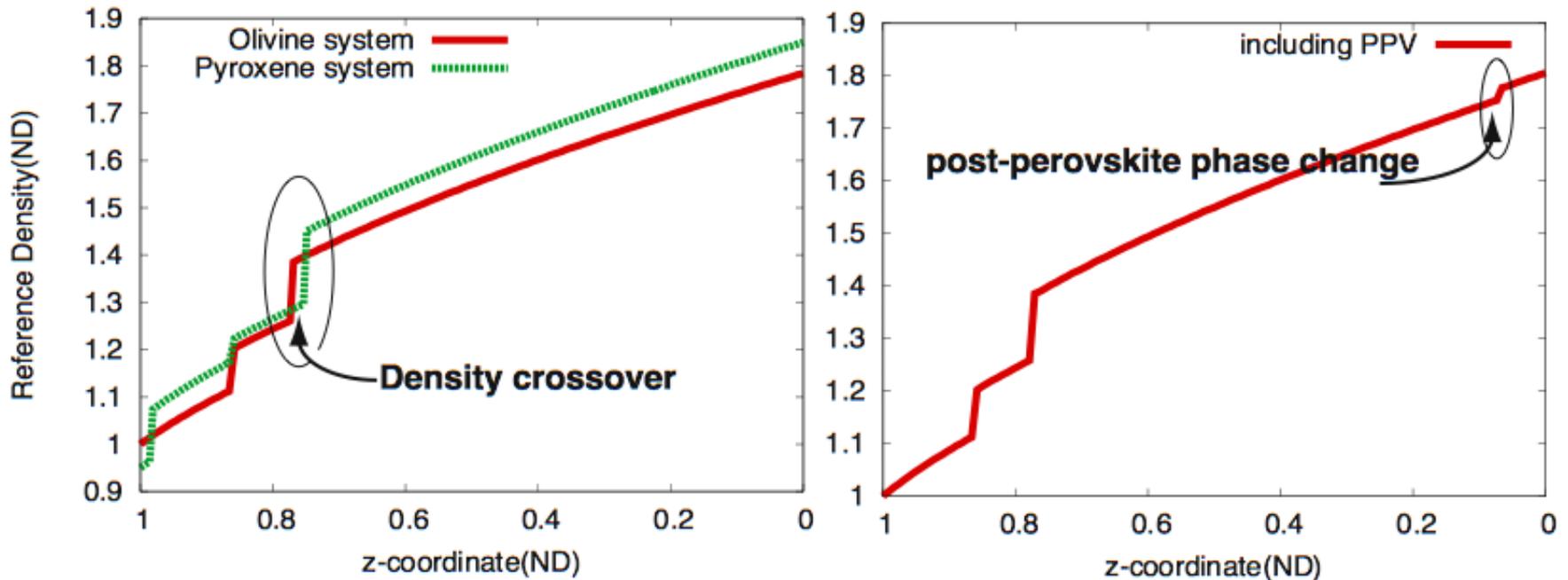


Lower mantle density for MORB  
(Ono et al., 2001)

Phase diagram and upper mantle density  
(Irifune and Isshiki, 1998)

Realistic mantle: 60% olivine system and 40% pyroxene system with the density crossover between 660km and 720km depths.

# Reference density profiles

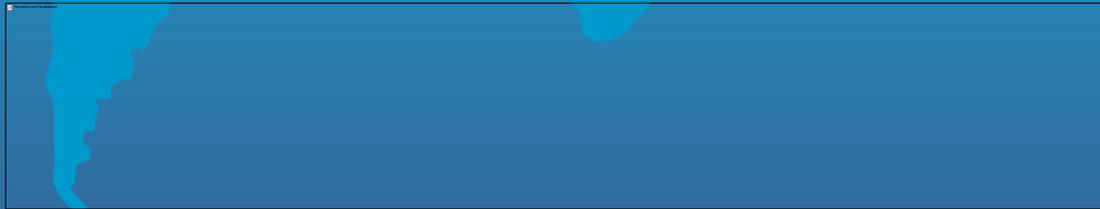


# Implementations for phase changes



$\Gamma$ : Sheet mass anomalies for  
olivine-spinel-perovskite phase changes  
pyroxene-garnet-perovskite phase changes

For the post-perovskite phase change:



## Two series of numerical simulations

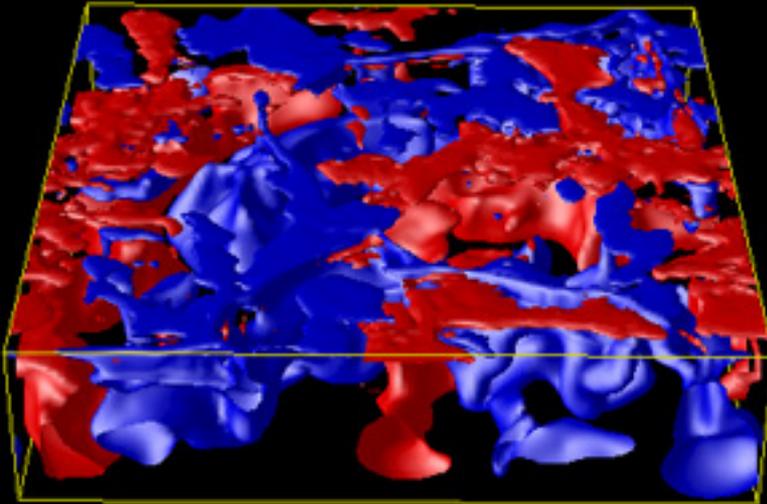
- Multiphase thermo-chemical convection:
  - No post-perovskite effects
  - 60% olivine and 40% pyroxene .
  - 100% olivine cases are done for comparing with realistic composition cases.
- Effects of post-perovskite phase change:
  - No pyroxene phase change system.
  - Two cases (+16MPa/K with 2% density difference between basaltic material and harzburgite, +0MPa/K with 2% density anomalies).

# Example 1: Non-olivine effects

100% Olivine

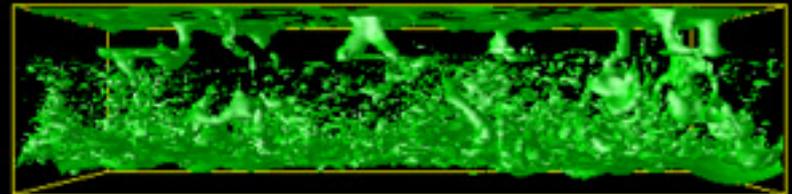
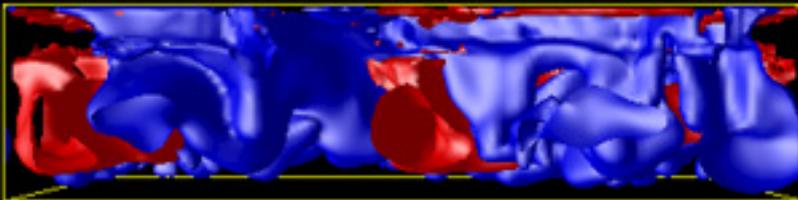
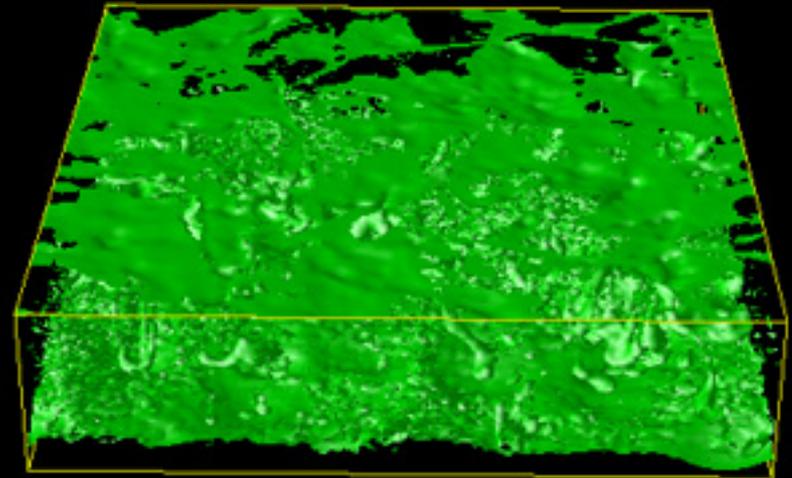
(a) Temperature residual

(red: +0.1; blue: -0.1)

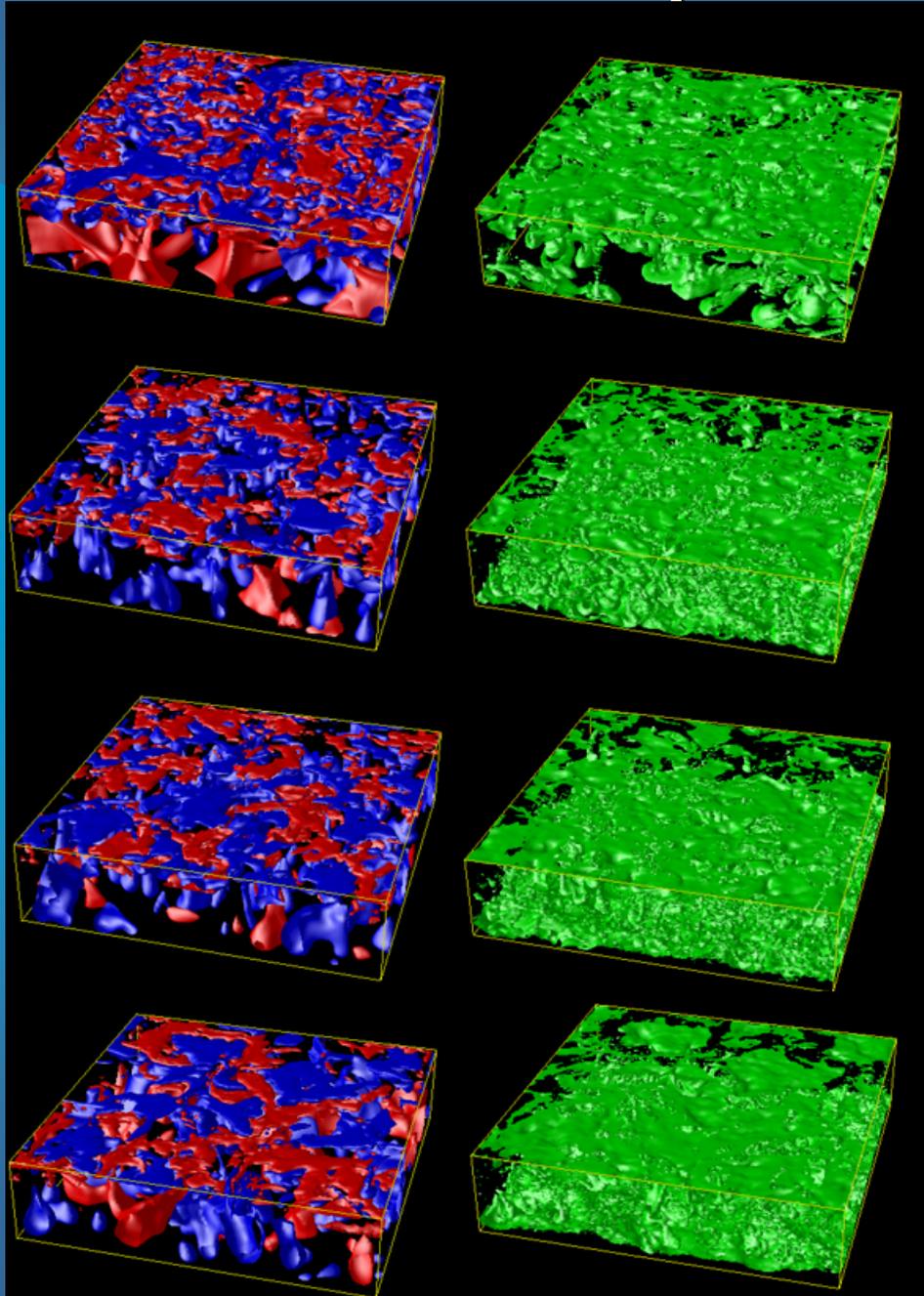


(b) Compositional isosurface

(C=0.4)



# Time variation of temperature and composition



$t = 0.00236$

$t = 0.00472$

$t = 0.00708$

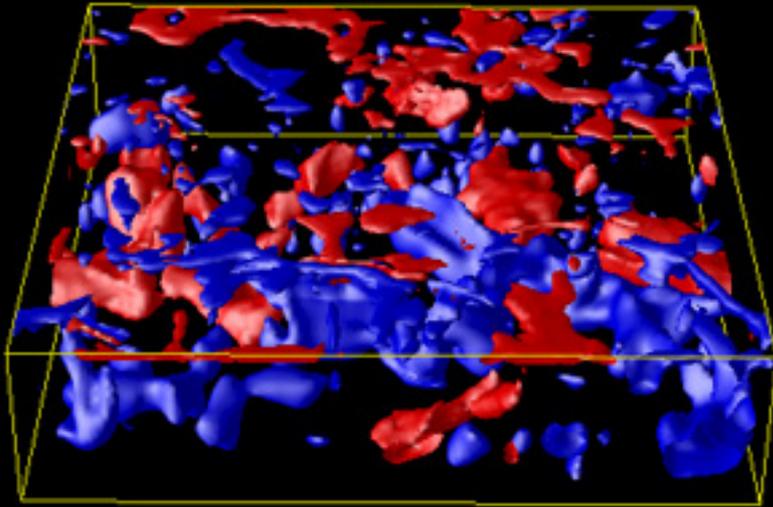
$t = 0.00944$

# Example 1: Non-olivine effects

60% Olivine + 40% Pyroxene (referred as 60/40)

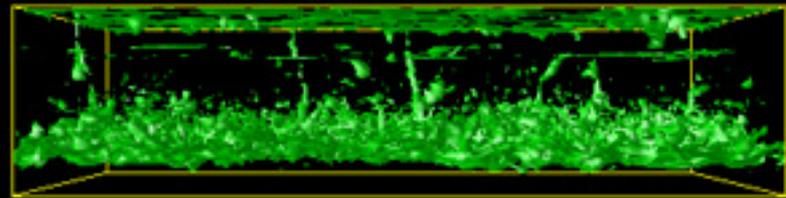
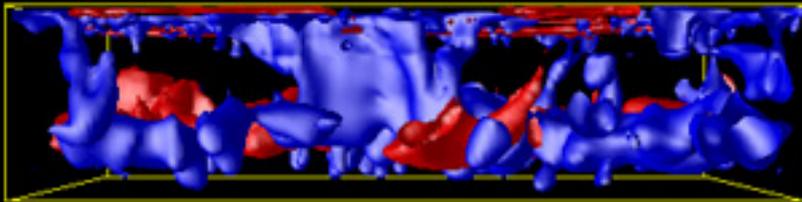
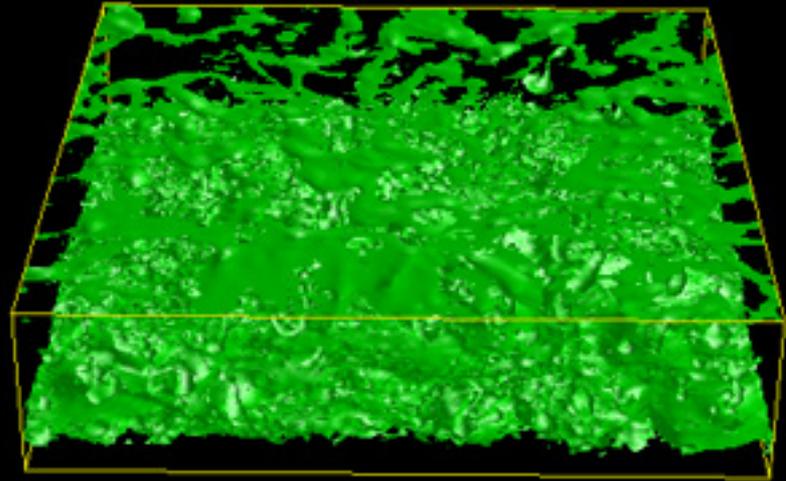
(a) Temperature residual

(red: +0.1; blue: -0.1)

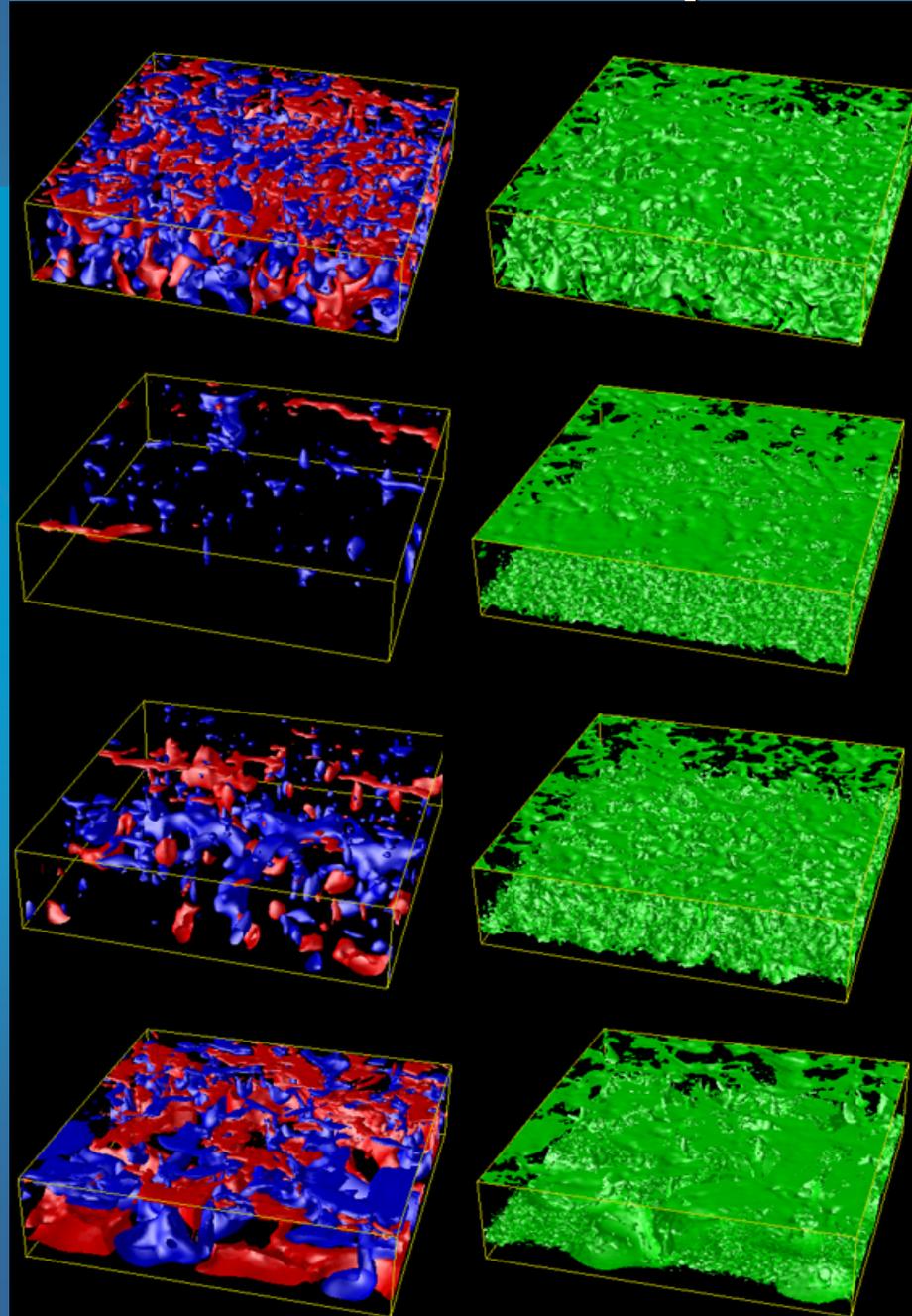


(b) Compositional isosurface

(C=0.4)



# Time variation of temperature and composition



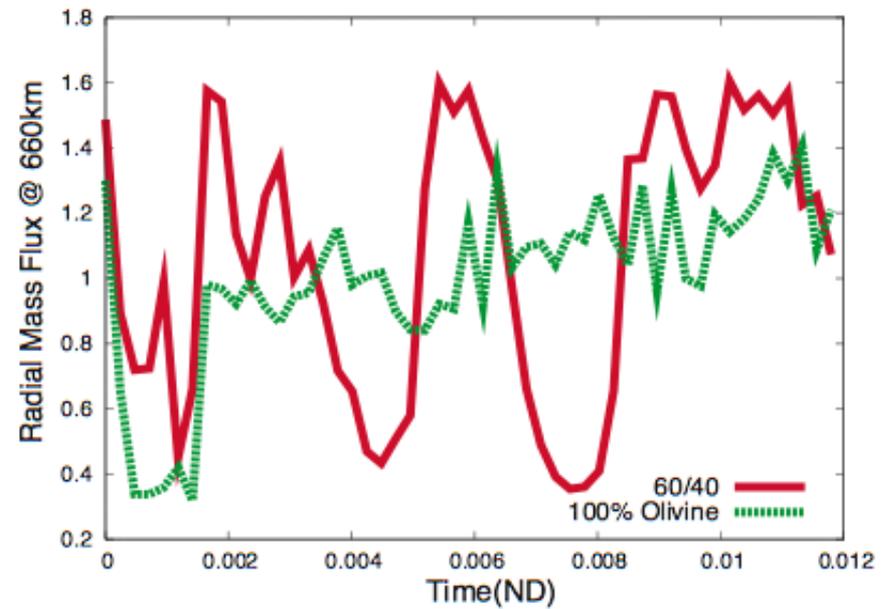
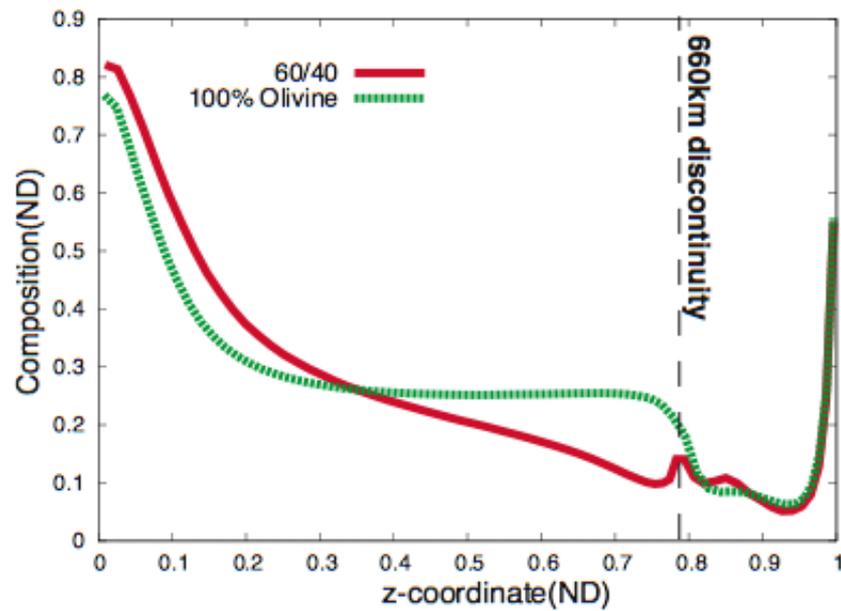
$t = 0.00236$

$t = 0.00472$

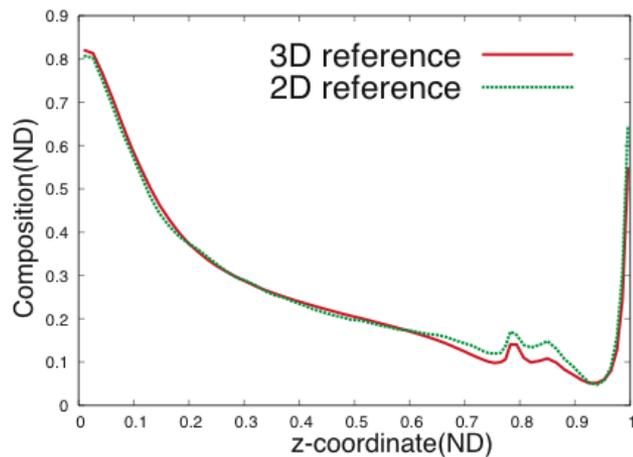
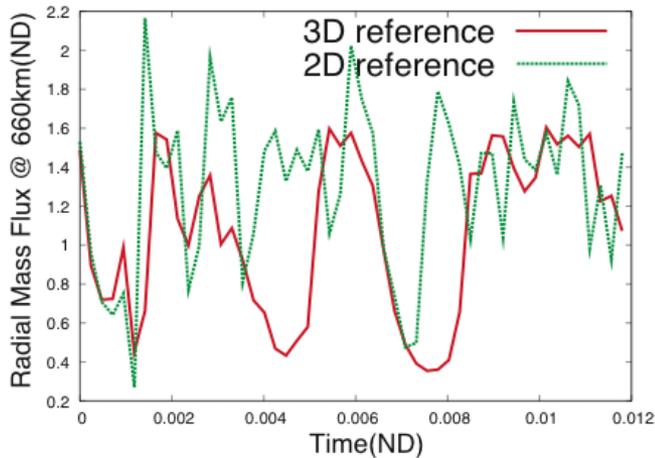
$t = 0.00708$

$t = 0.00924$

# Comparison of diagnostics



# Diagnostics (Compositional structures and mass flux) compared to two-dimensional cases



Top: Radial Mass Flux @ 660km  
Bottom: Horizontally-averaged C  
Both plots are 60/40 cases.

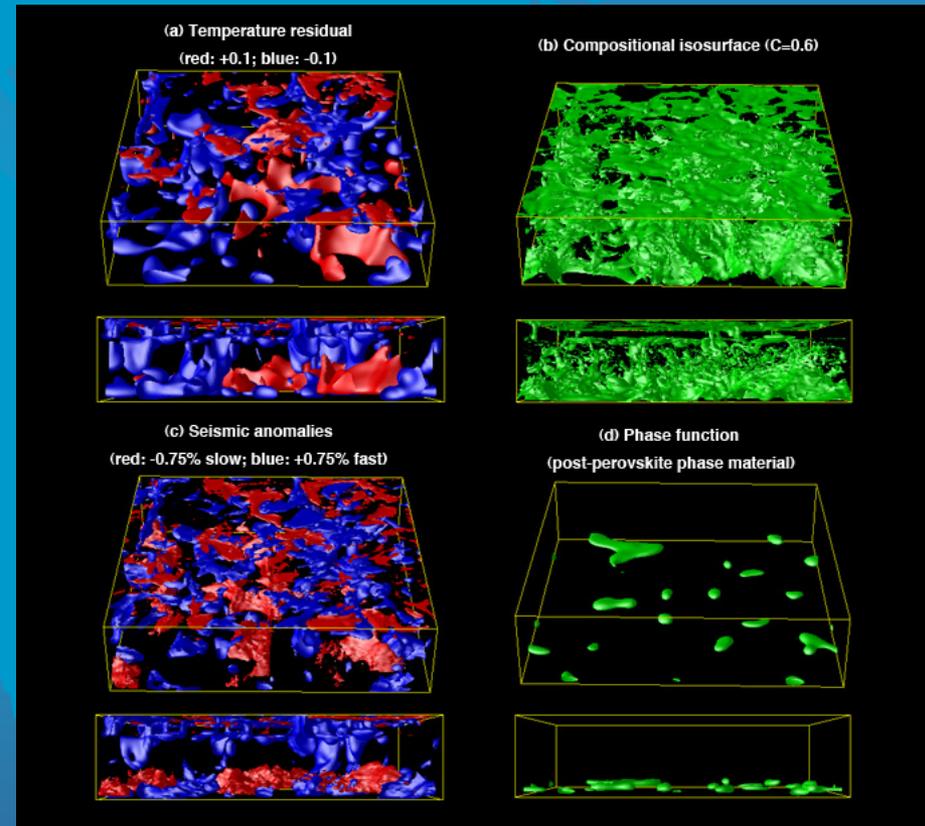
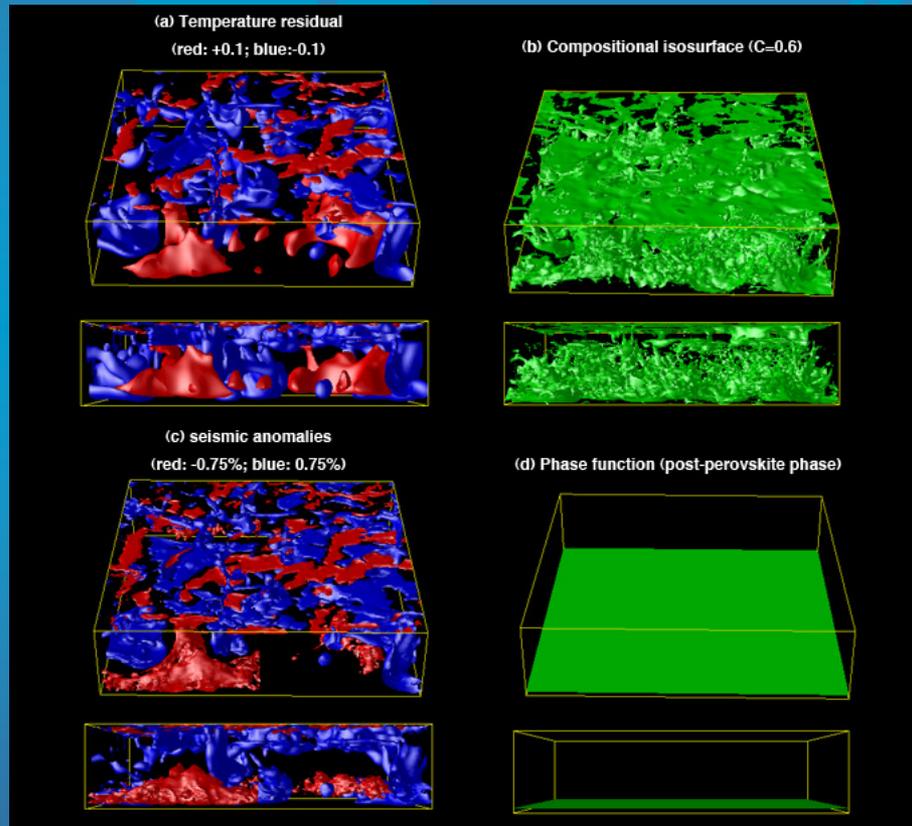
Time variation for mass flux:  
3D case has longer term variation

Vertical structure of C:  
Not quite different between both cases

# Example 2: Effects of post-perovskite phase change near the core-mantle boundary

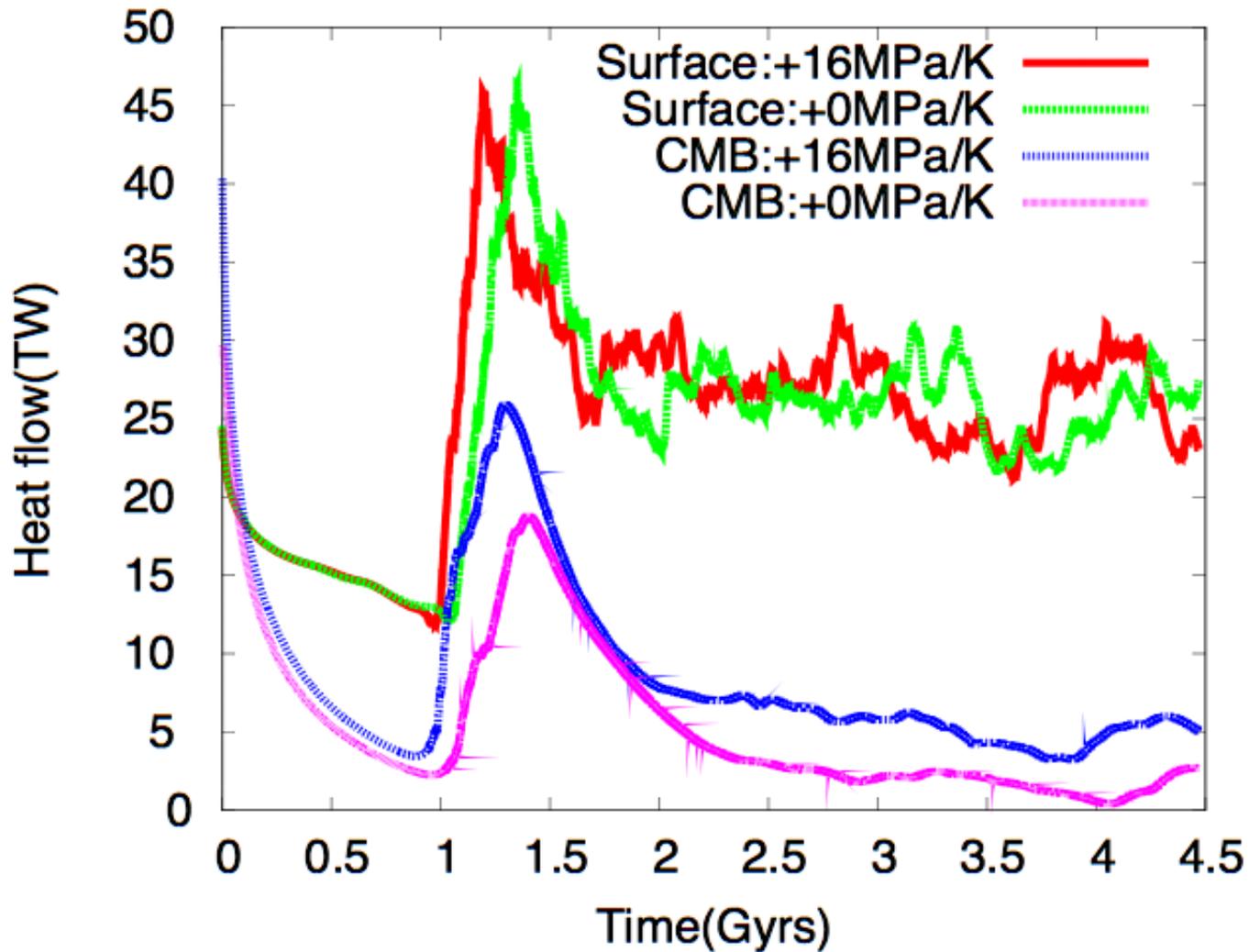
Zero Clapeyron slope

Exaggerated Clapeyron slope  
(+16MPa/K)



Both cases are 2% density difference between basaltic and harzburgite

# Heat flow as a function of time



# Summary

- Three-dimensional numerical simulations of thermo-chemical multiphase convection in Earth's mantle including the melt-induced differentiation can be done for a realistic computational time.
- Two examples of complicated phase changes and newly discovered phase change are shown using three-dimensional cases.
- Comparing between two- and three-dimensional results, flux across a certain boundary is longer time scale for 3D cases than for 2D cases.
- Higher heat flow across the CMB is expected when the post-perovskite phase change is taken into account, which is consistent with preliminary two-dimensional results [Nakagawa and Tackley, 2004].
- Necessary to expand a spherical geometry for STAG3D and investigate more realistic thermo-chemical evolution in Earth's mantle in the future.

## Next direction (due date: AGU Fall meeting ?)

- Non-olivine effects cases: Change the compressibility in the lower mantle and comparing to 2D cases for the formation of the dense piles.
- Post-perovskite phase change cases: See what happens the cases of realistic Clapeyron slope and doing various cases (0% density difference and more denser cases).
  - Comparing with 2D cases as well.

A faint, light blue world map is centered in the background of the slide. The map shows the outlines of the continents: North America, South America, Europe, Africa, Asia, and Australia. The map is semi-transparent and serves as a decorative element.

おしまい!

The End !