Low velocities in the oceanic upper mantle and their relations to plumes: insights from SEM-based waveform tomography

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Outline

- Historical background and methods

 > from normal modes to SEM
- Strategy and model development
- SEMum2 model salient features and the challenge of validation
- Outlook

Time domain waveform inversion in global seismology

Woodhouse and Dziewonski (1984)



- Normal mode theory
- Path AVerage Approximation (PAVA)->
 1D sensitivity kernels
- Later, complement with body wave travel times (ray theory) to access lower mantle structure



Toroidal modes : $n \mathbf{T} \mathbf{l}$



l: angular order, horizontal nodesn: overtone number, vertical nodes

4) Non-linear asymptotic coupling theory (NACT)-> 2D kernels in the vertical plane (Li and Romanowicz, 1995)



Full Waveform Tomography of the whole mantle



- NACT: Surface waves, overtones (T>80s), body waves (T>32 s)

-Misfit function: waveform difference + Windowing and weighing, in order to equalize amplitudes.

Several generations of whole mantle shear velocity models,
Including radial anisotropy, attenuation
→ (Li and Romanowicz, 1996->Panning and Romanowicz, 2006)





Shear velocity Depth = 2800 km

"LLSVP" "Superplumes"



Lekic et al, 2012, EPSL

The road forward: Full Waveform Tomography using SEM:



Replace mode synthetics by numerical synthetics computed using the Spectral Element Method (SEM)

UC Berkeley Global Seismology Group The 1-D Reference Earth

The View so far

Our immediate goal



Strategy:

- Take "one step at a time" and in the process hope to learn something new about the earth:
- I "Hybrid" inversion approach:
 - Compute forward wavefield precisely using C-SEM
 - Keep same framework for time domain inversion as in previous model developments (3 component, wavepackets)
 - Compute inverse kernels using NACT
- II -Use "homogenized", smooth crustal model appropriate for the period range considered
- III Start at long periods (T>60s) and progress towards shorter periods, "assimilating" more data

I-HYBRID INVERSION APPROACH

At each iteration:

1-Forward modeling step

Use coupled spectral element method of Capdeville et al. (2003) to accurately forward model wave propagation through the 3D Earth



 Γ 1= Normal modes in 1D Γ 2 = Spectral element method

2-Inverse step

Use approximate Hessian calculated in NACT. Much faster than adjoint!



Li and Romanowicz, 1995

II-Smooth homogeneized crustal model

 Construct an equivalent smooth anisotropic crustal model (Backus, 1962 – effective medium theory)

- Crustal thickness saturated at 30 km

 Start with filtered Crust2.0

- Fit global dataset of dispersion maps (*Ritzwoller et al.,* 2002) using Monte Carlo approach



Top: Smooth Moho of SEMum2 (km); **Bottom**: SEM (*C=0.4*) time step (s)

=> SEM time step prolonged ~4 times



Long period synthetics filtered to 60s



Long period synthetics filtered to 60s



Long period synthetics filtered to 60s

SEMum and SEMum2

- Full waveforms, T> 60 s, 204 events
- Radially anisotropic models
 - Vs (isotropic shear velocity)
 - $\xi = (Vsh/Vsv)^2$
- Invert top 800 km of the mantle:
 - Lower mantle from existing tomographic model SAW24B16
- 1st generation model: uniform crustal thickness of 60 km
 - Start with 1D model
 - Progressively add waveforms and refine parametrization
 - -> 10 iterations: SEMum (Lekic and Romanowicz, 2011 GJI)
- 2nd generation model: SEMum2 variable thickness crust, finer scale parametrization, 2+ iterations (French et al., 2013)

SEMum2: 204 events ~ 100,000 Wavepackets

(Z,R,T) waveforms - Surface waves - Body waves

~5M indep. data constraints

+Surface wave dispersion 25-60s to constrain crust (Shapiro and Ritzwoller, 2002)















SEMum2 validation using RegSEM



SEMum2 validation using RegSEM











Isosurface levels: -1-> -3%%











Geographic extent of oceanic region OR2 in clustering analysis of SEMum with N=6, and the location of major hotspots



SEMum2 at 250 km depth



SEMum2 at 250 km depth







Bootstrapping Test: 20 resampled data realizations





SEMum2 at 250km depth:

Directional wavelet analysis of geoid:

Azimuth of dominant direction

60

70

80 90 100 110 120 130 140 150 160 170 180

"Richter Rolls"

Versus...

Richter and Parsons, 1975

Viscous fingering fed by plume-like conduits

Weeraratne and Parmentier, in prep.

 $(\underbrace{=}_{\mu 2} \underbrace{=}_{\mu 1}) \ddagger B$

Outlook

- Complete geodynamic interpretation of the low-velocity finger structures
- Extend to shorter periods (30-40s) and include body wavepackets
- Attenuation and azimuthal anisotropy
- Combine gradient computation (adjoint) with approx. Hessian
- Implement summed event (encoded source) approach

Source stacking

Input model (saw6) 75km depth

-0.014 0.000 0.014

2800km depth

Output model (iteration 3)

-0.014587 0.000000 0.014587

2800km depth

Capdeville, Gung, Romanowicz, GJI, 2005

Summed seismograms at station XAN

XAN

Black: observed trace (filtered between 60 and 250 s) *Red*: RegSEM synthetic in the 3D N-Born starting model

